

# *The Global Climate Thermostat: Fact, fiction, and computer models*

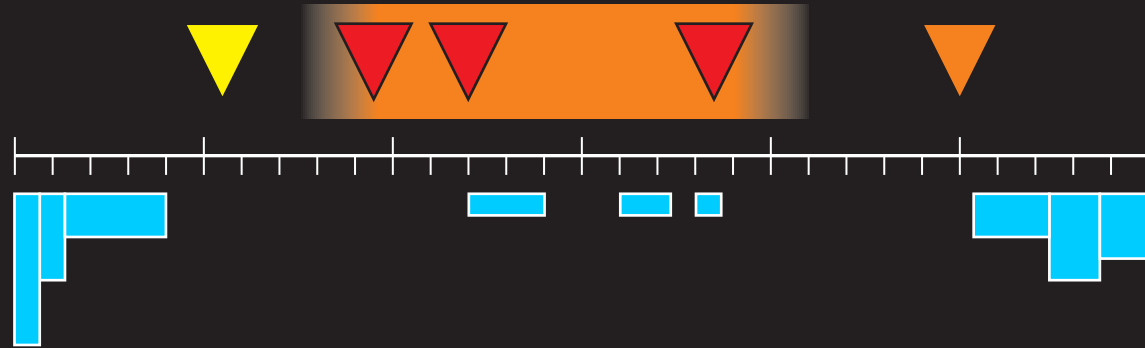
Andy Ridgwell



# Regulation of global climate



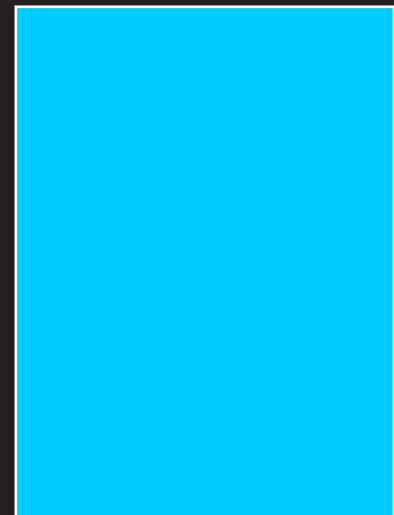
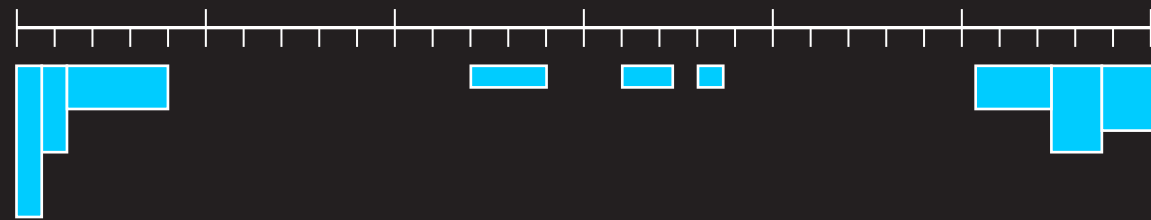
# Regulation of global climate



Occurrence of ice ages (relative intensity)  
source; Crowell [1999]



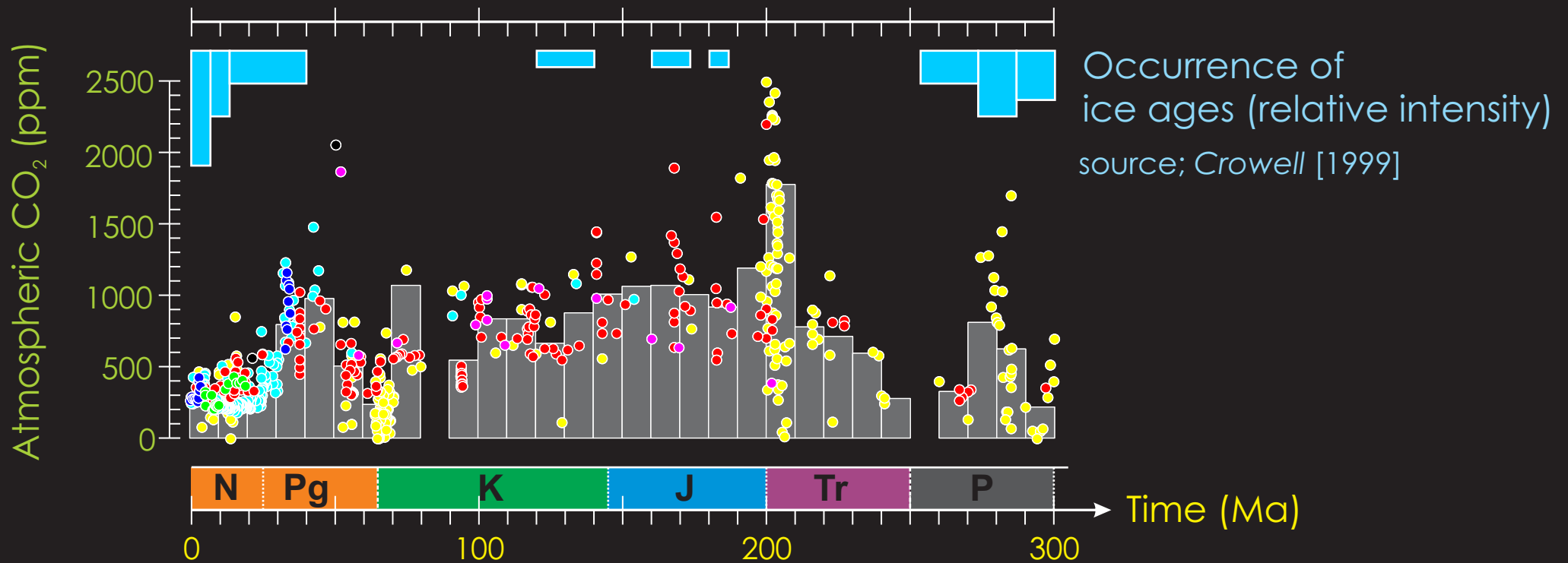
# Regulation of global climate



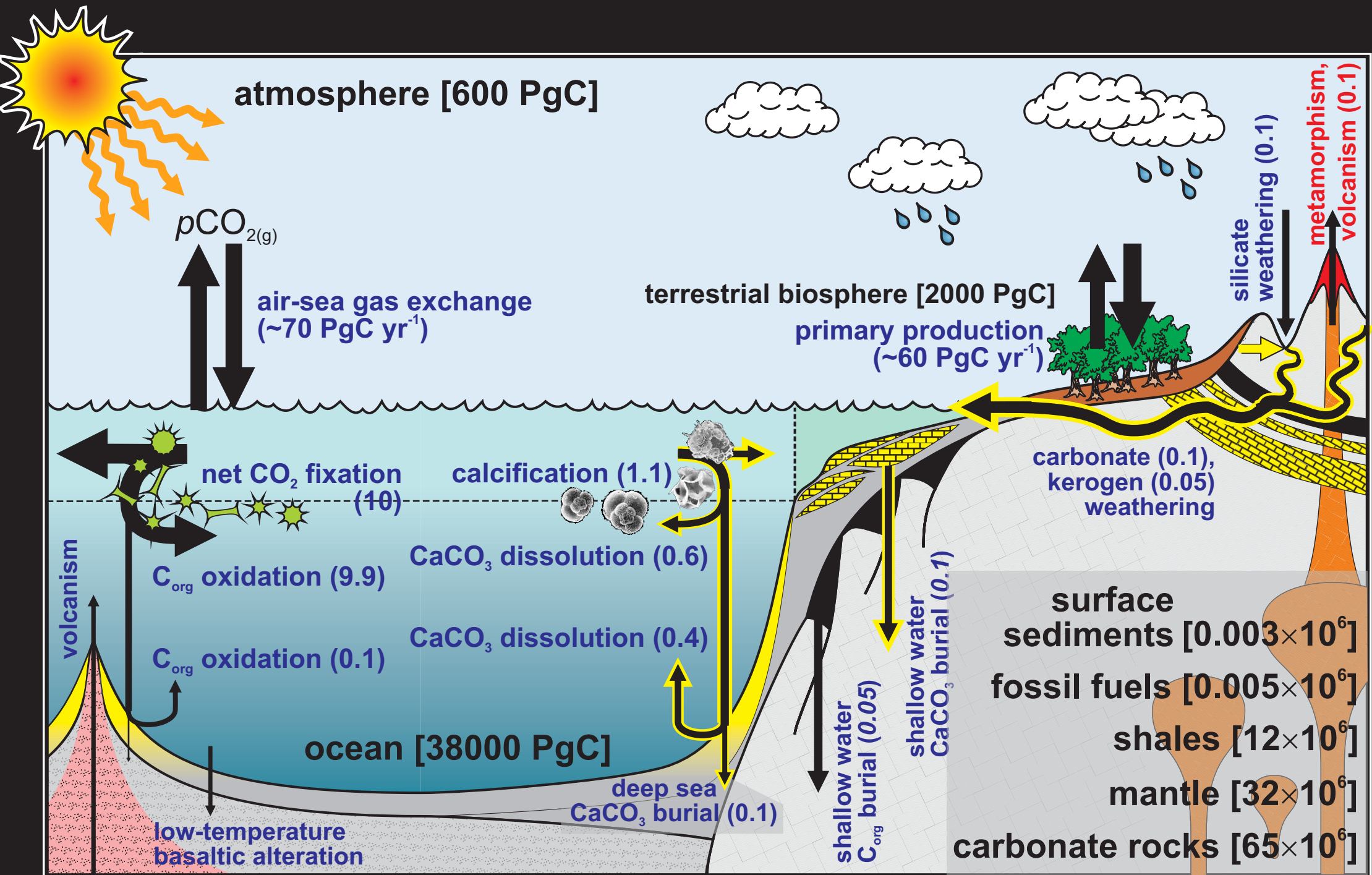
# Regulation of global ~~climate~~ carbon cycling



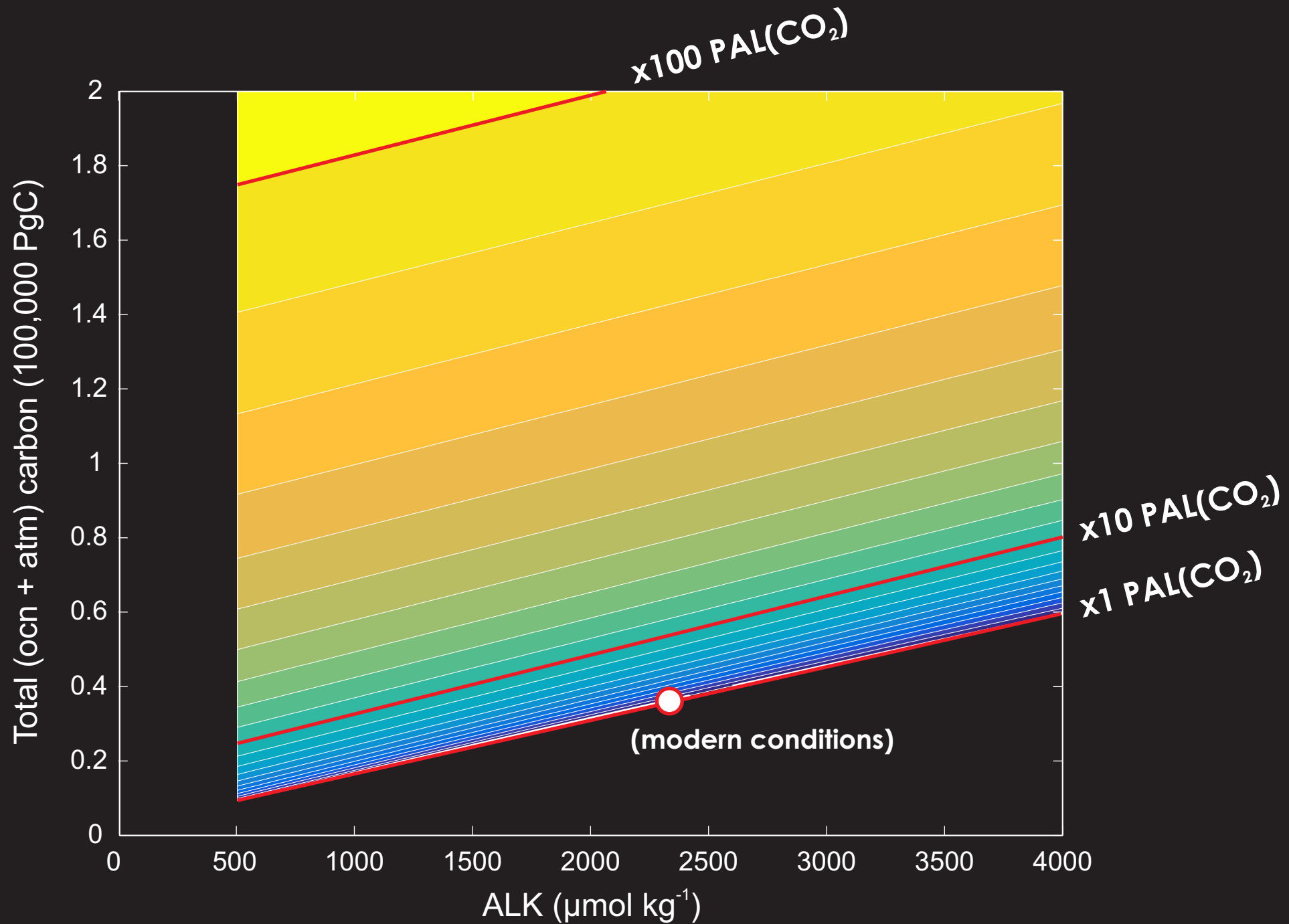
From: *Hönisch et al. [2012]*



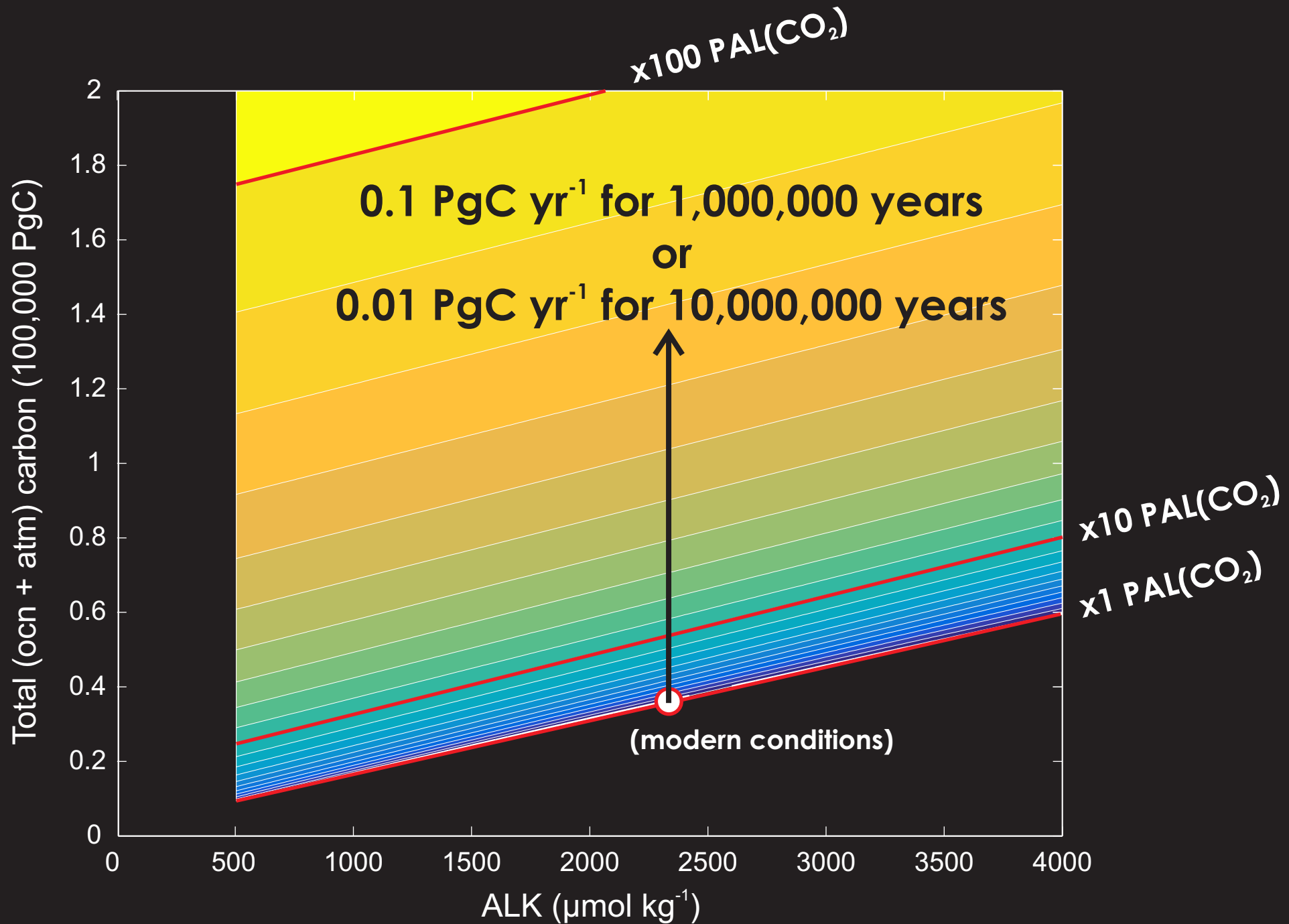
# Regulation of global ~~climate~~ carbon cycling



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# Regulation of global ~~climate~~ carbon cycling





Terrestrial weathering can be (approximately equally) divided into carbonate ( $\text{CaCO}_3$ ) and calcium-silicate (' $\text{CaSiO}_3$ ') weathering:



Ultimately, the (alkalinity:  $\text{Ca}^{2+}$ ) weathering products must be removed through carbonate precipitation and burial in marine sediments:



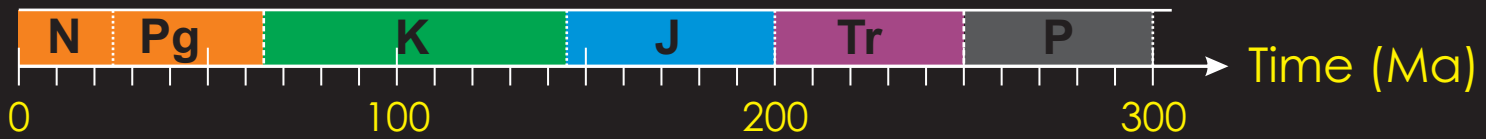
It can be seen that in (2) + (3), that the  $\text{CO}_2$  removed (from the atmosphere) during weathering, is returned upon carbonate precipitation (and burial). In (1) + (3) (silicate weathering)  $\text{CO}_2$  is permanently removed to the geological reservoir. This  $\text{CO}_2$  must be balanced by mantle (/volcanic) out-gassing on the very long term.

# Regulation of global ~~climate~~ carbon cycling



Furthermore, the rate of silicate weathering should scale with climate. Hence the **silicate weathering feedback** is formed:

higher  $p\text{CO}_2$   $\rightarrow$  higher temperatures (& rainfall)  $\rightarrow$  higher weathering rates  $\rightarrow$  lower  $p\text{CO}_2$



# lies, damn lies, and computer models



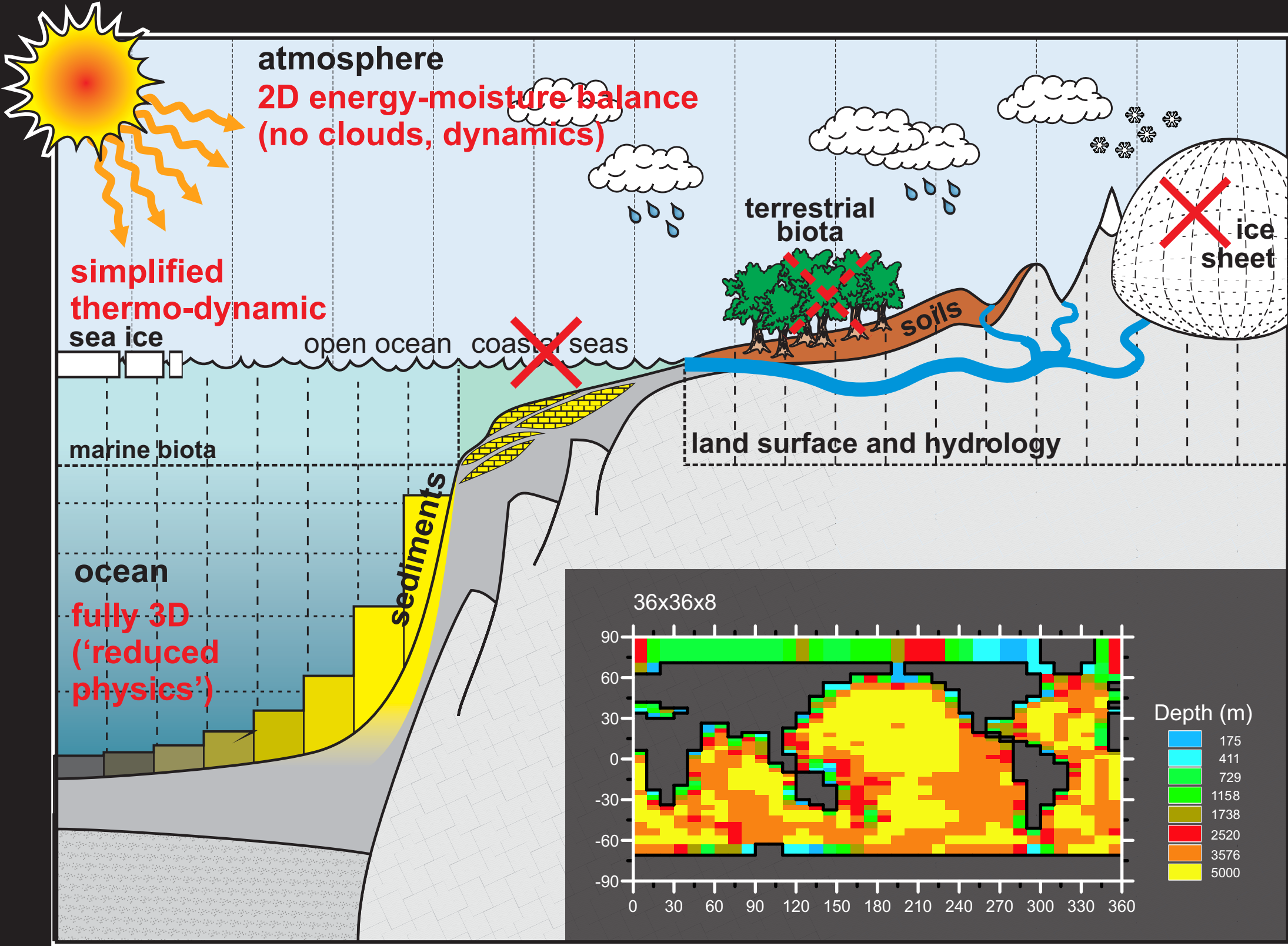
```
! calculate carbonate alkalinity
loc_ALK_DIC = dum_ALK &
& - loc_H4BO4 - loc_OH - loc_HPO4 - 2.0*loc_PO4 - loc_H3SiO4 - loc_NH3 - loc_HS &
& + loc_H + loc_HSO4 + loc_HF + loc_H3PO4

! estimate the partitioning between the aqueous carbonate species
loc_zed = ( &
& (4.0*loc_ALK_DIC + dum_DIC*dum_carbconst(icc_k) -
loc_ALK_DIC*dum_carbconst(icc_k))**2 + &
& 4.0*(dum_carbconst(icc_k) - 4.0)*loc_ALK_DIC**2 &
& )**0.5      loc_conc_HCO3 = (dum_DIC*dum_carbconst(icc_k) -
loc_zed)/(dum_carbconst(icc_k) - 4.0)

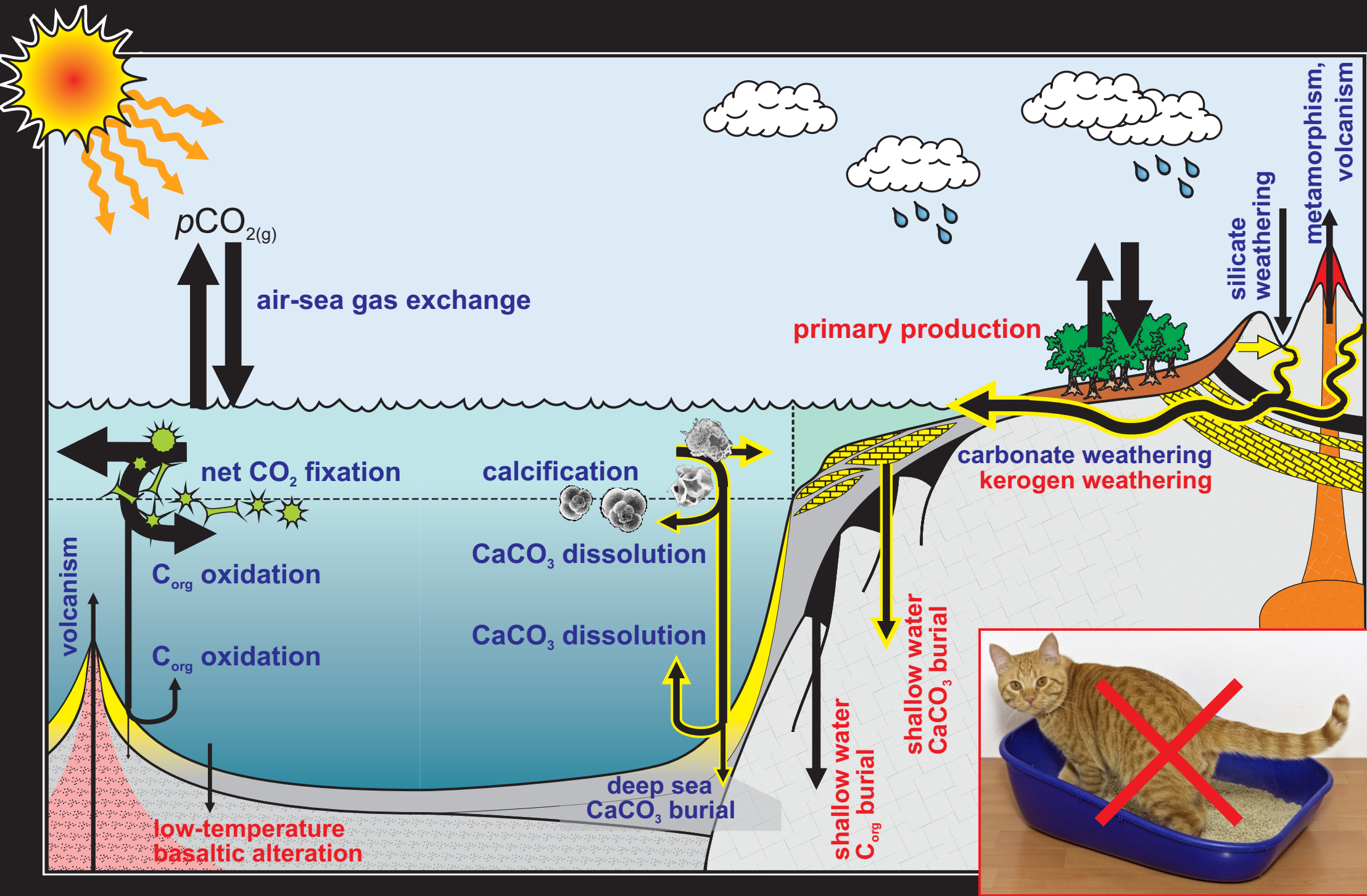
loc_conc_CO3 = &
& ( &
& loc_ALK_DIC*dum_carbconst(icc_k) - dum_DIC*dum_carbconst(icc_k) - &
& 4.0*loc_ALK_DIC + loc_zed &
& ) &
& /(2.0*(dum_carbconst(icc_k) - 4.0))

loc_conc_CO2 = dum_DIC - loc_ALK_DIC + &
& ( &
& loc_ALK_DIC*dum_carbconst(icc_k) - dum_DIC*dum_carbconst(icc_k) - &
& 4.0*loc_ALK_DIC + loc_zed &
& ) &
& /(2.0*(dum_carbconst(icc_k) - 4.0))

loc_H1 = dum_carbconst(icc_k1)*loc_conc_CO2/loc_conc_HCO3
loc_H2 = dum_carbconst(icc_k2)*loc_conc_HCO3/loc_conc_CO3
```



# lies, damn lies, and computer models

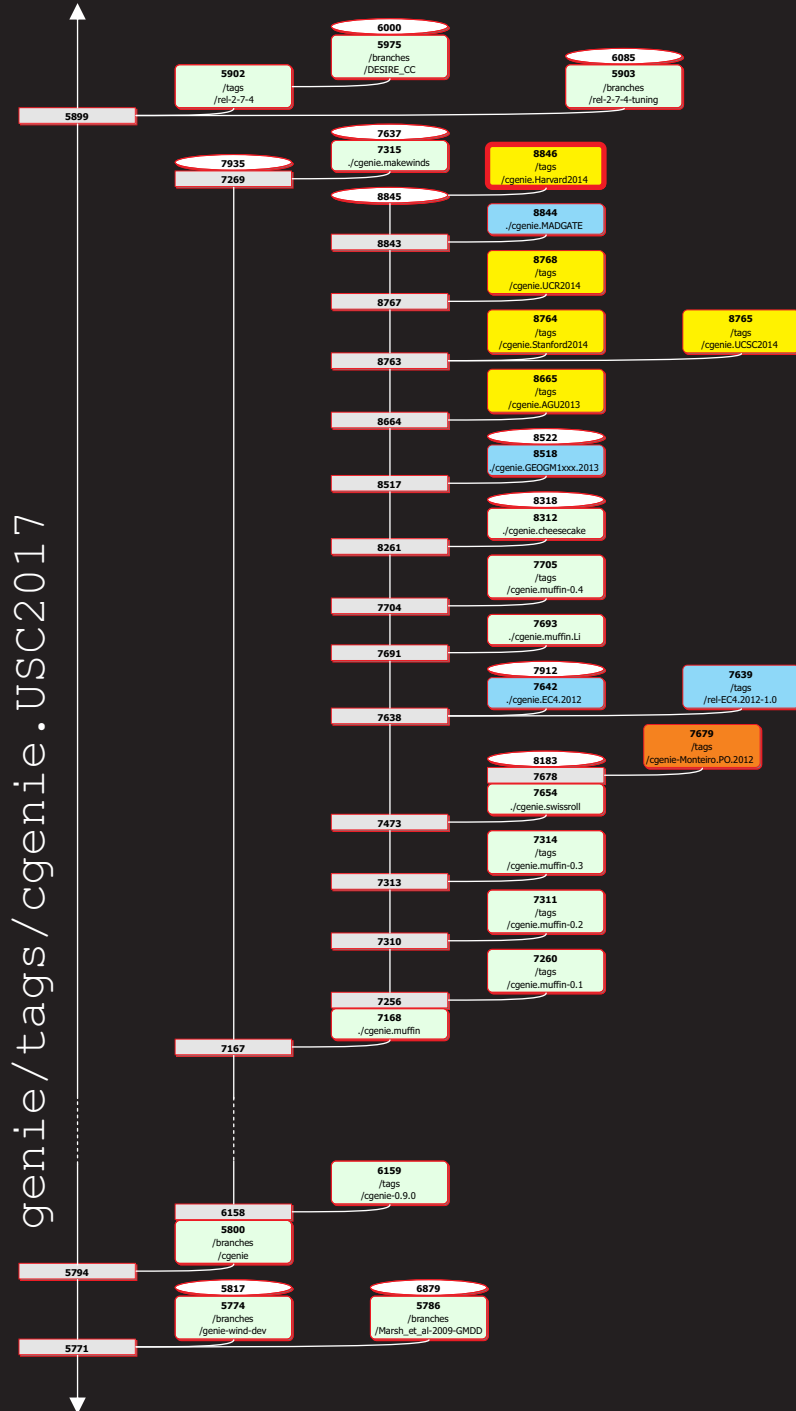


# lies, damn lies, and computer models



[http://www.seao2.info/misc\\_usc2017.html](http://www.seao2.info/misc_usc2017.html)

<https://svn.ggy.bris.ac.uk/subversion/genie/tags/cgenie.USC2017>

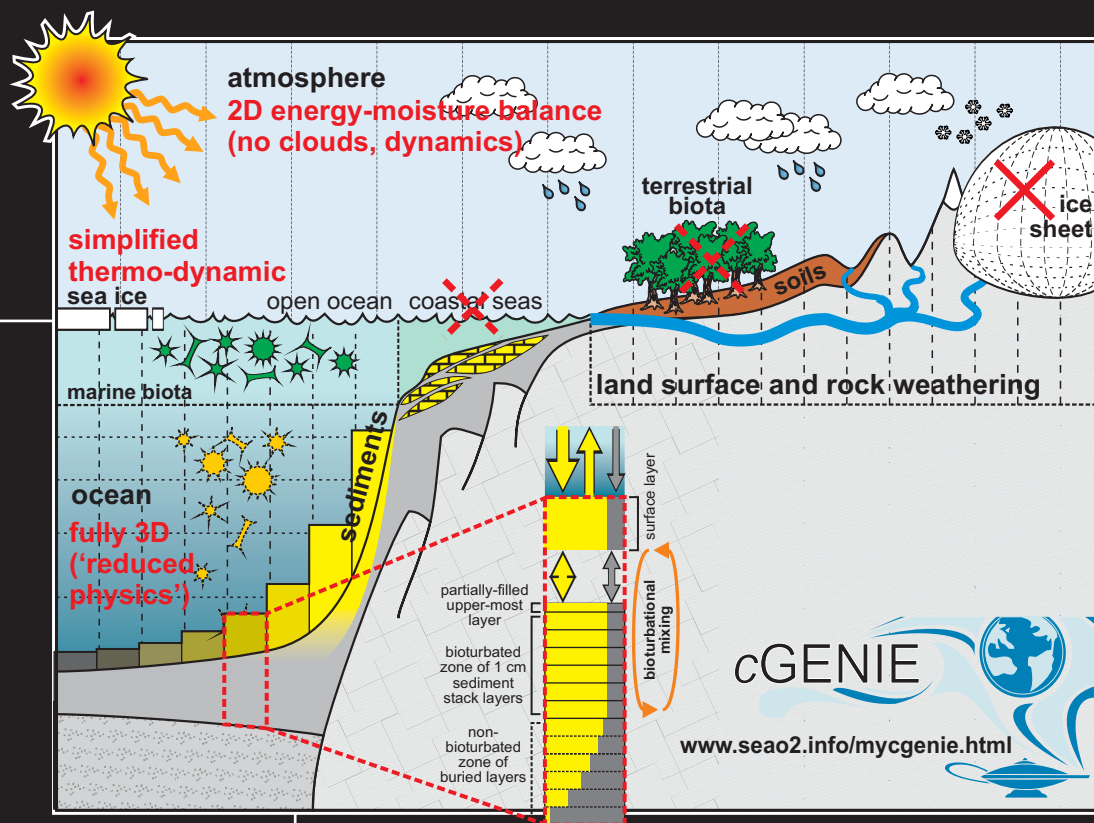
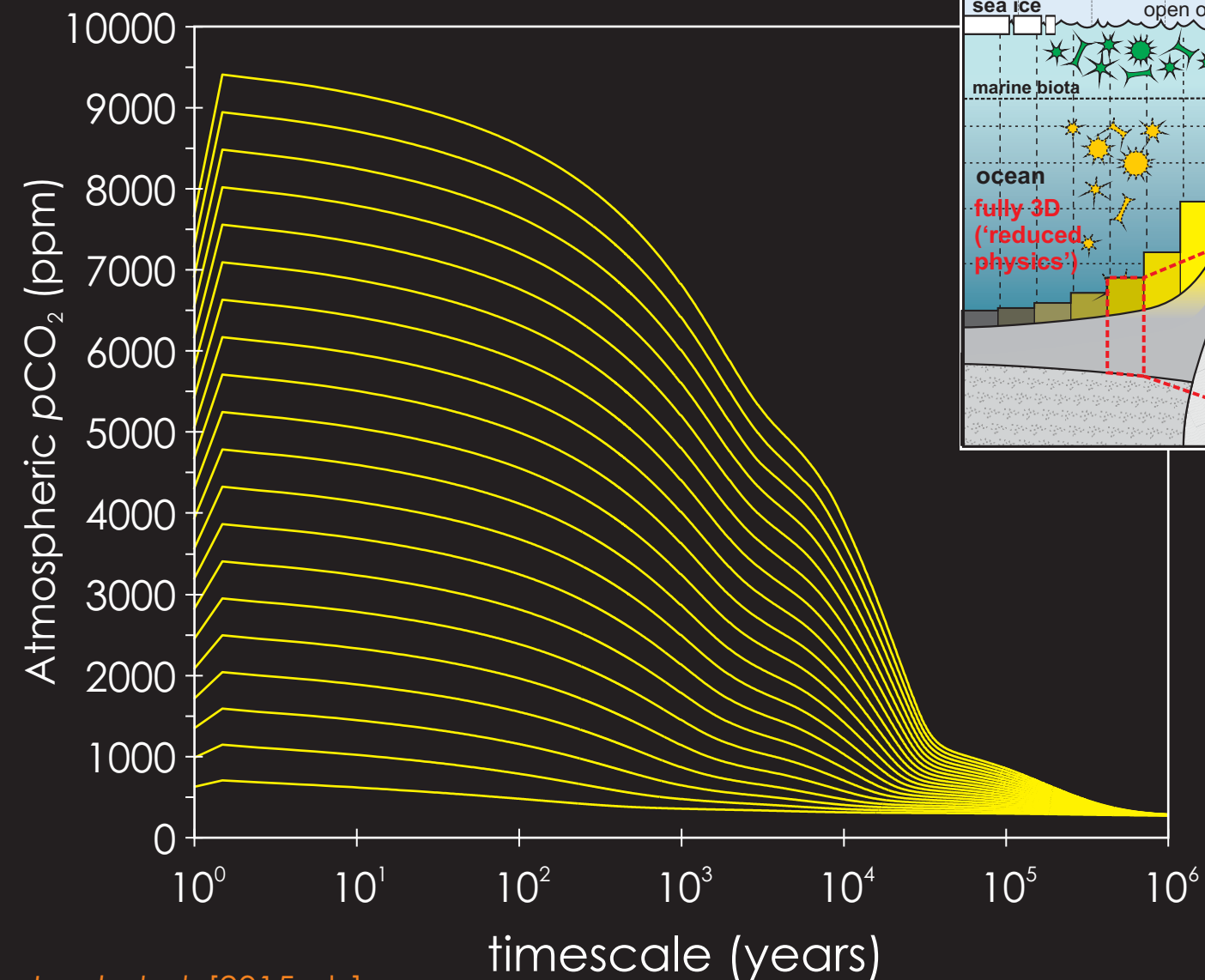




# Impulse response function analysis of the 'long tail' of CO<sub>2(excess)</sub>



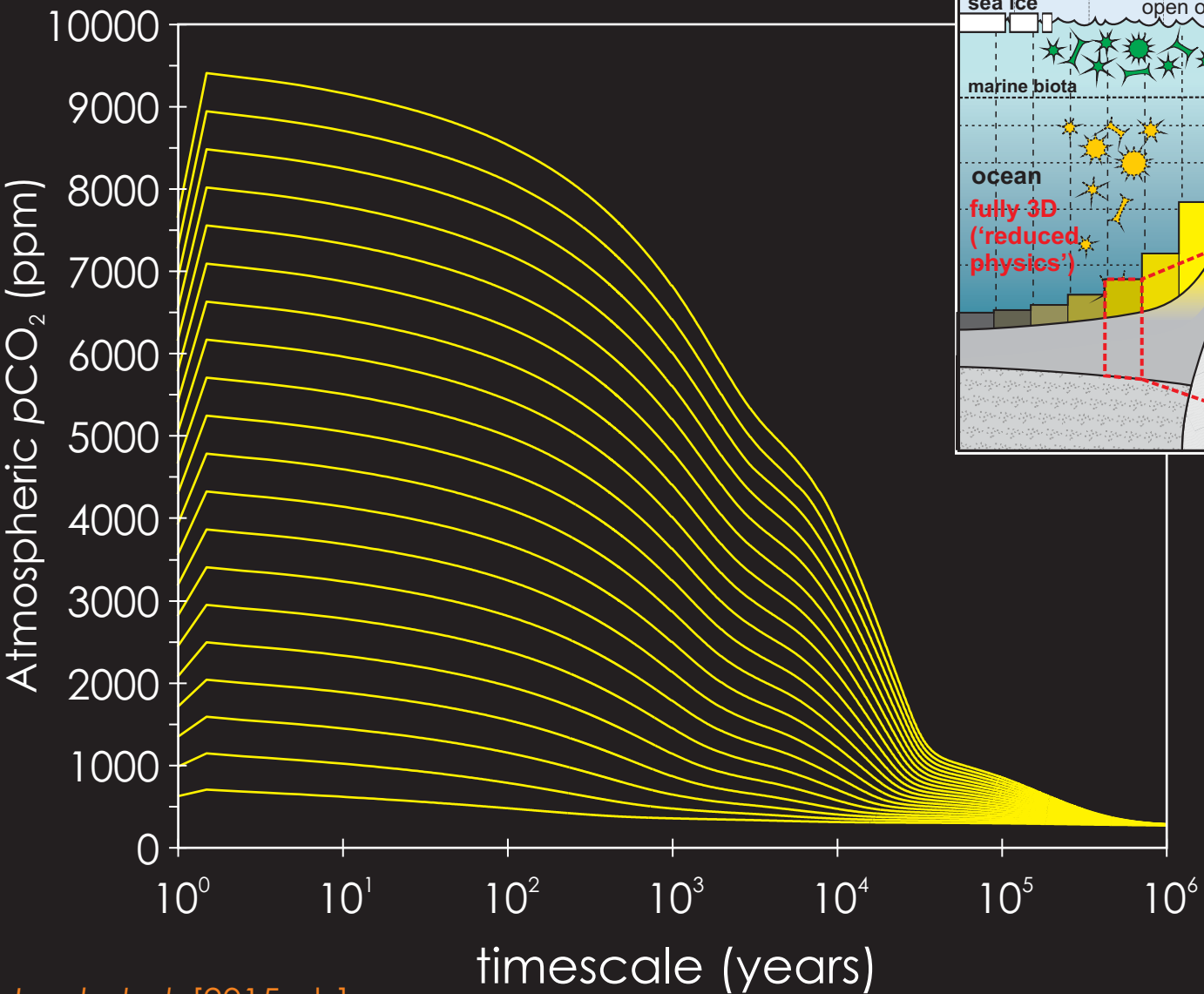
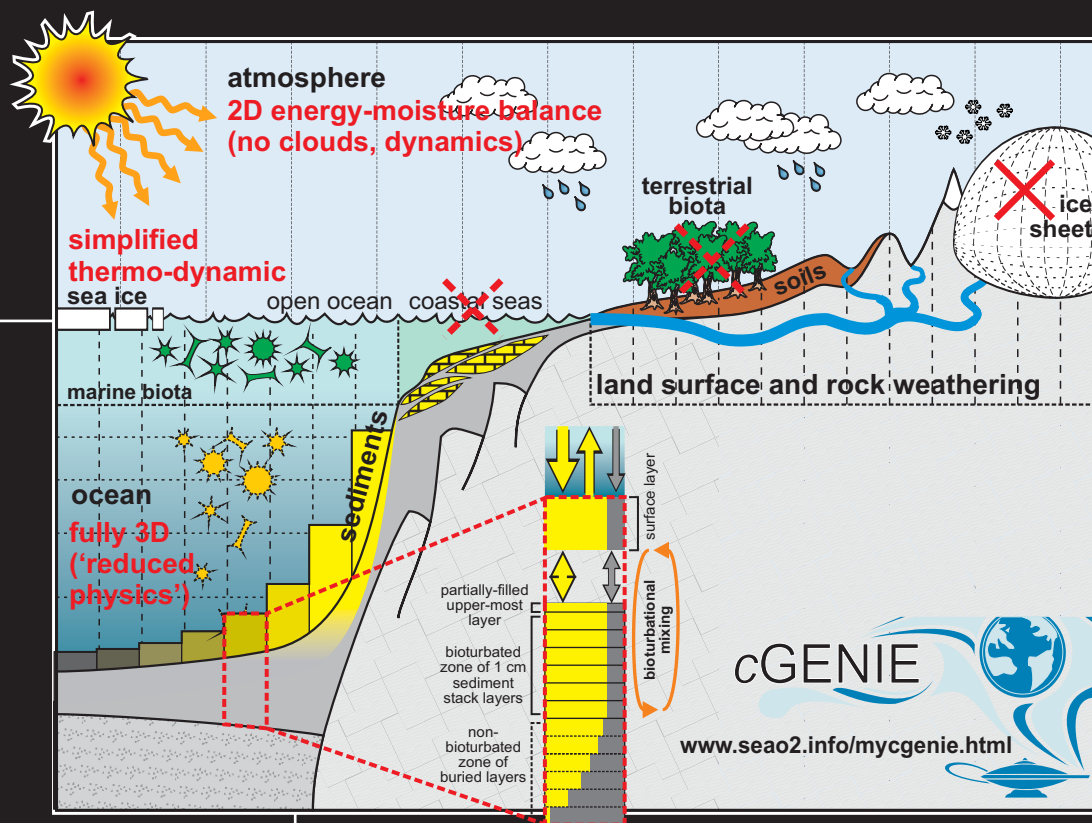
(1) Series of 1 Myr Earth system model experiments. CO<sub>2</sub> emissions from 1,000 to 20,000 PgC (GtC). Release interval: 1 yr.



# Impulse response function analysis of the 'long tail' of CO<sub>2(excess)</sub>



(1) Series of 1 Myr Earth system model experiments. CO<sub>2</sub> emissions from 1,000 to 20,000 PgC (GtC). Release interval: 1 yr.

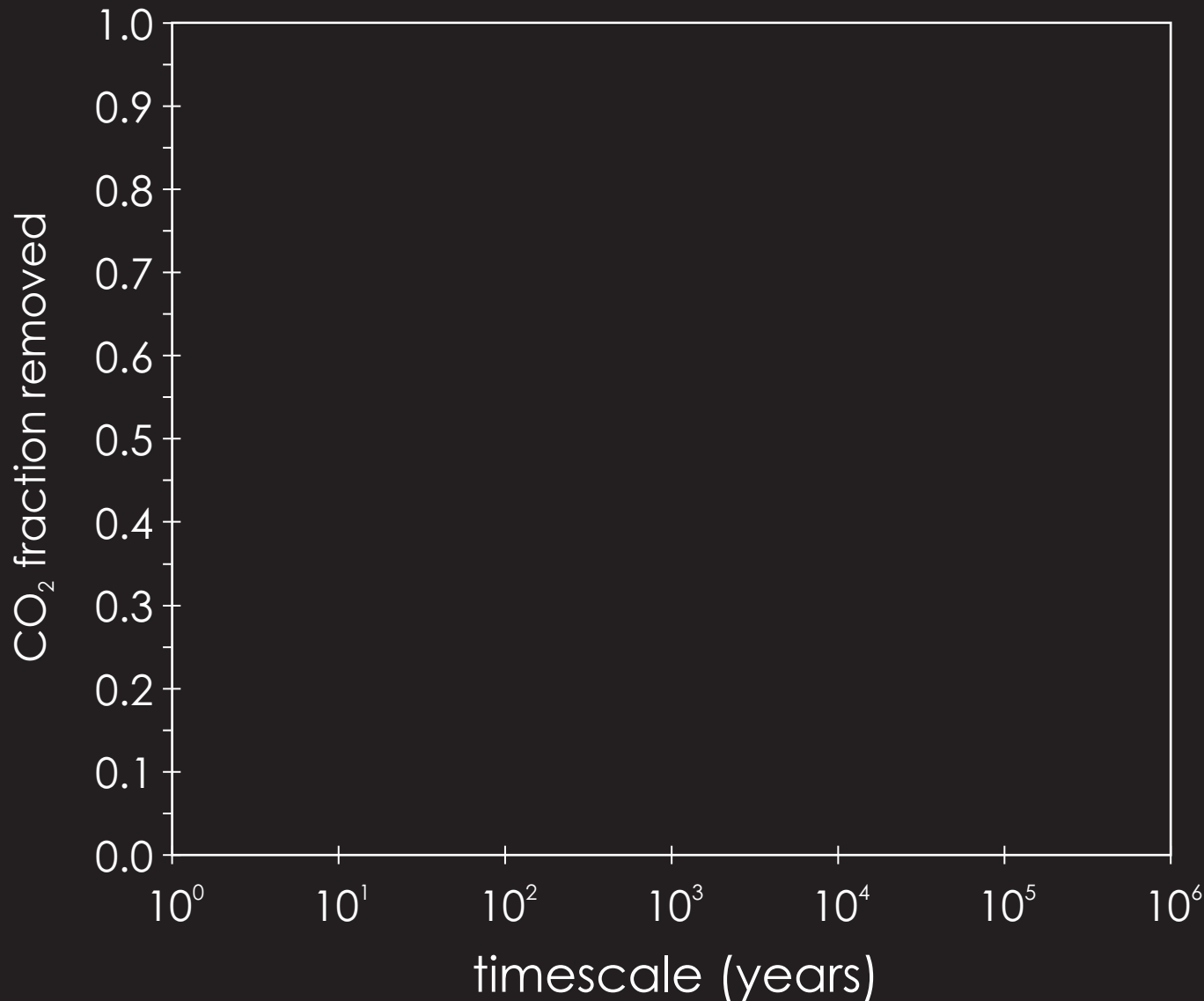


(2) Fit each CO<sub>2</sub> decay curve with a series (4 optimal) of exponentials. Extract the fraction of CO<sub>2</sub> and time-scale associated with each.

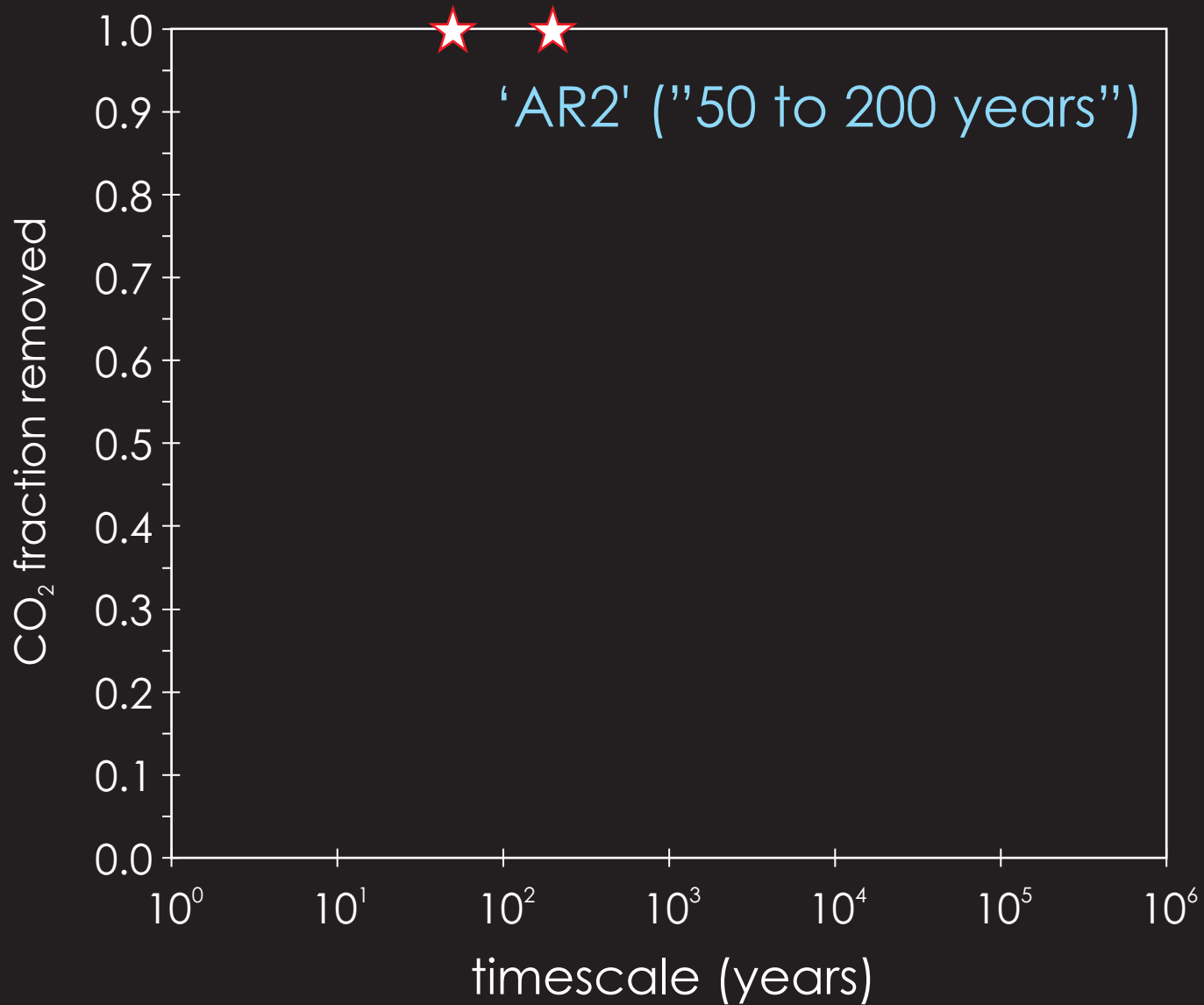
(The resulting empirical model can be used in place of a mechanistic model for projecting the long-term fate of carbon release.)



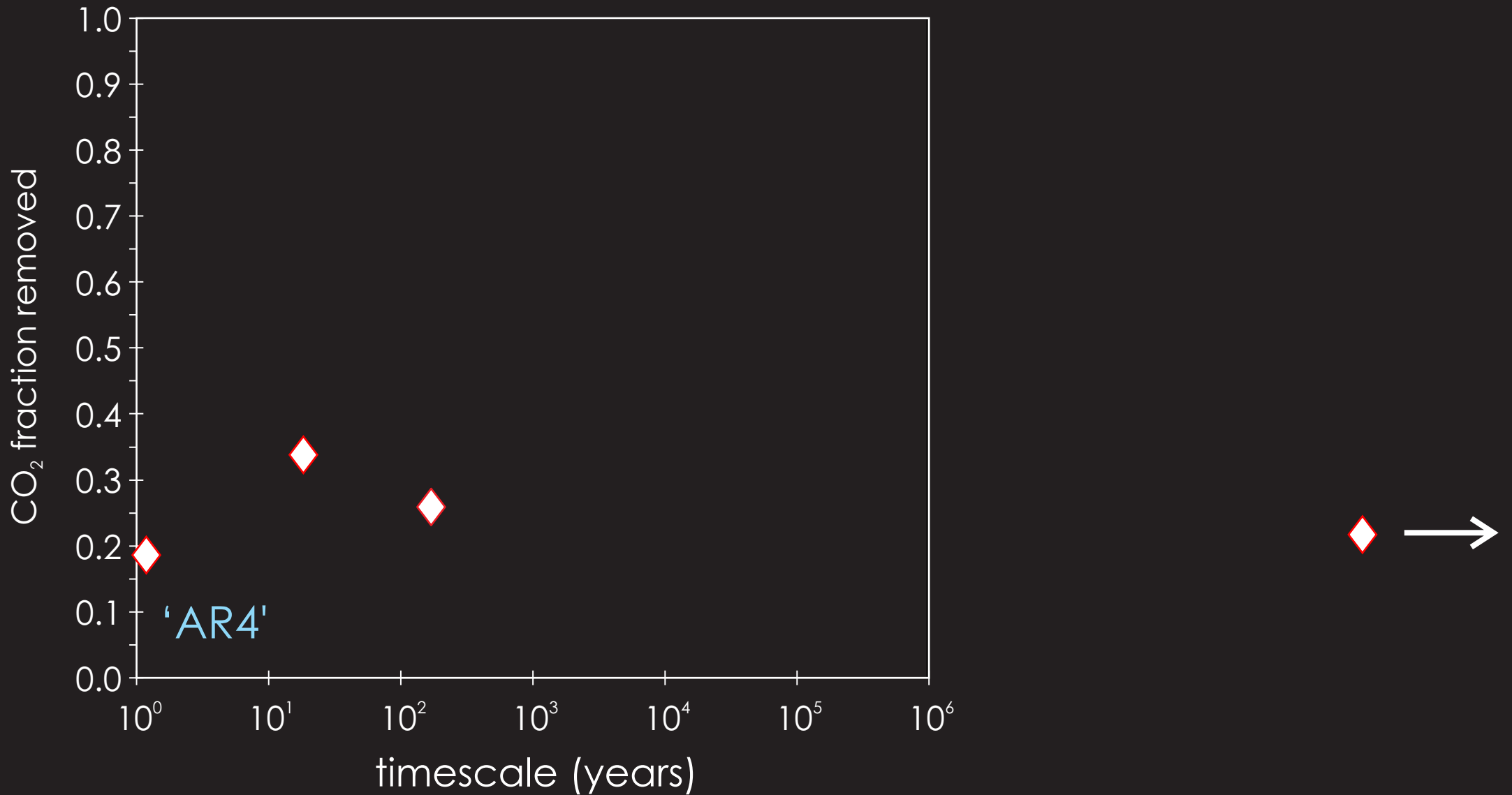
Cross-plot of the fraction of total  $\text{CO}_2$  emissions to the atmosphere removed by a particular process (carbon sink), vs. the characteristic (e-folding) time-scale of that process ( $\log_{10}$  scale).



# Impulse response function analysis of the 'long tail' of $\text{CO}_2(\text{excess})$

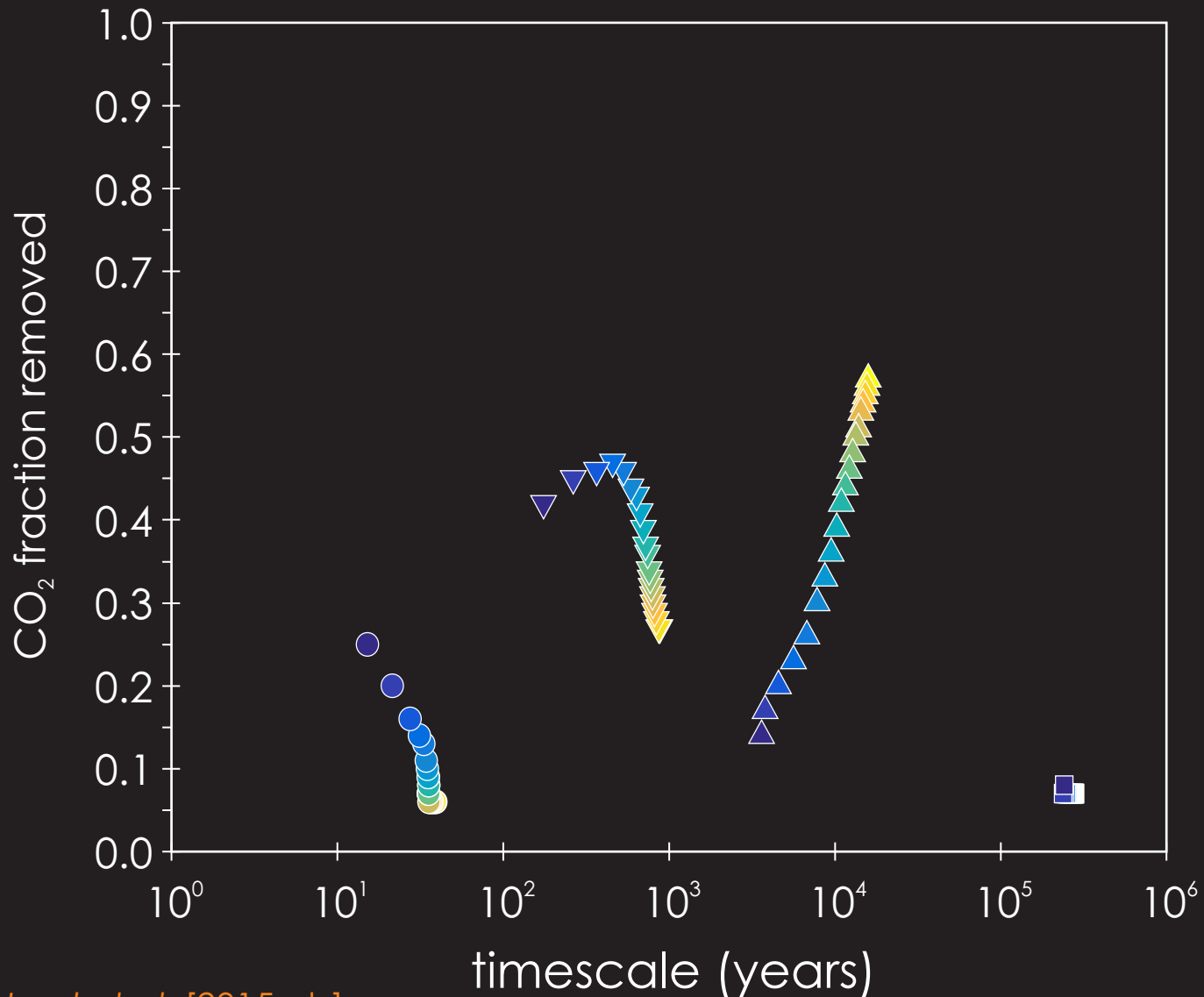


# Impulse response function analysis of the 'long tail' of $\text{CO}_{2(\text{excess})}$



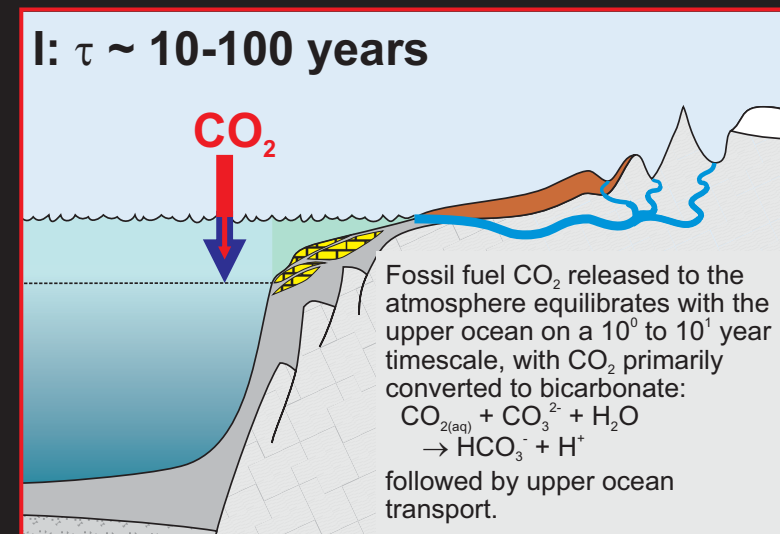
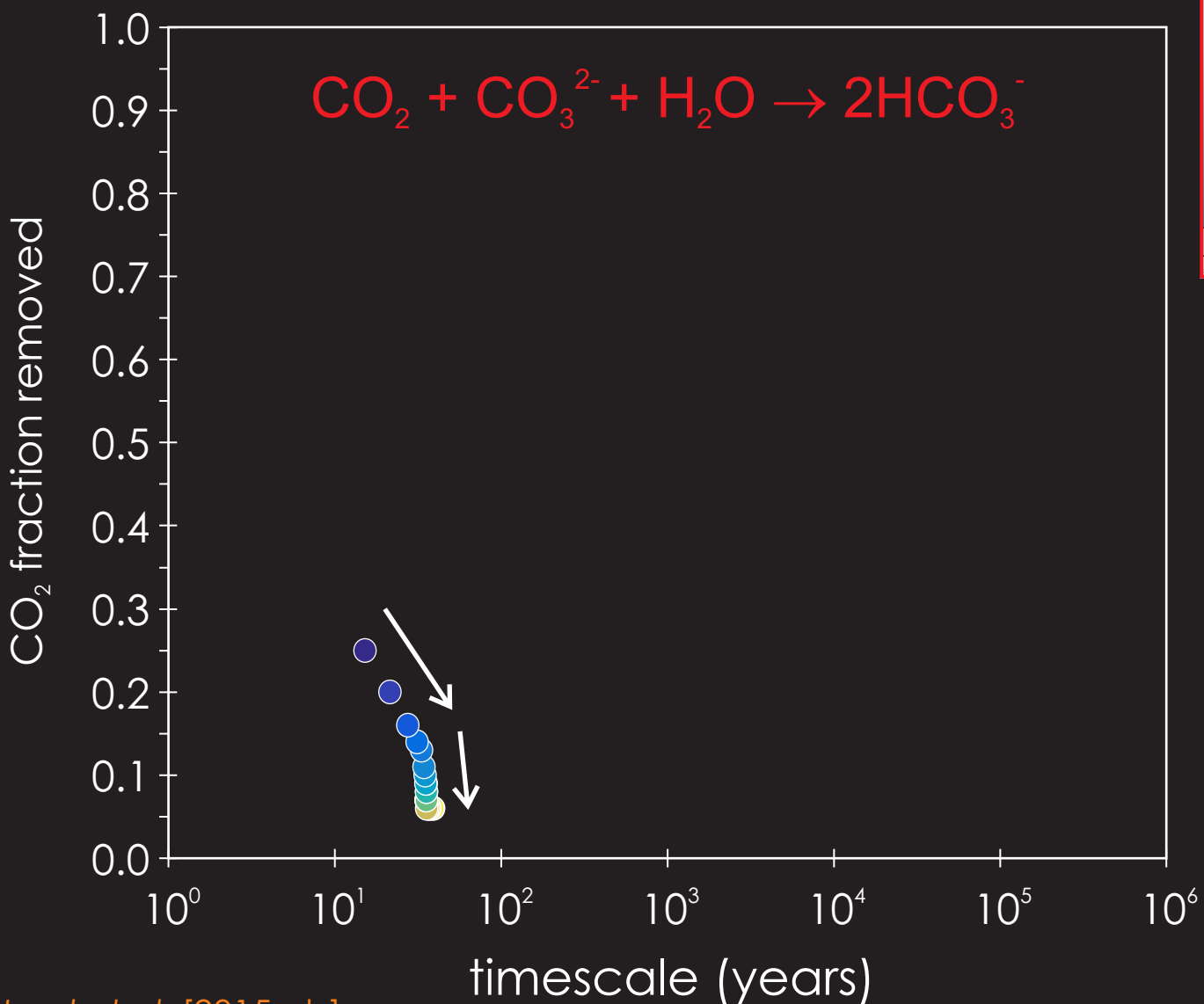


Response of fraction of  $\text{CO}_2$  removed vs. the characteristic time-scale, as a function of total emissions, ranging from 1,000 PgC (dark blue) to 20,000 PgC (yellow).



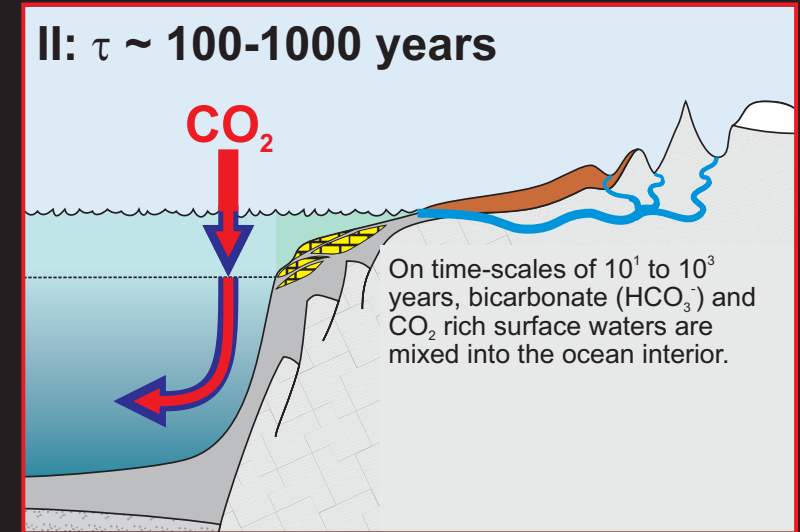
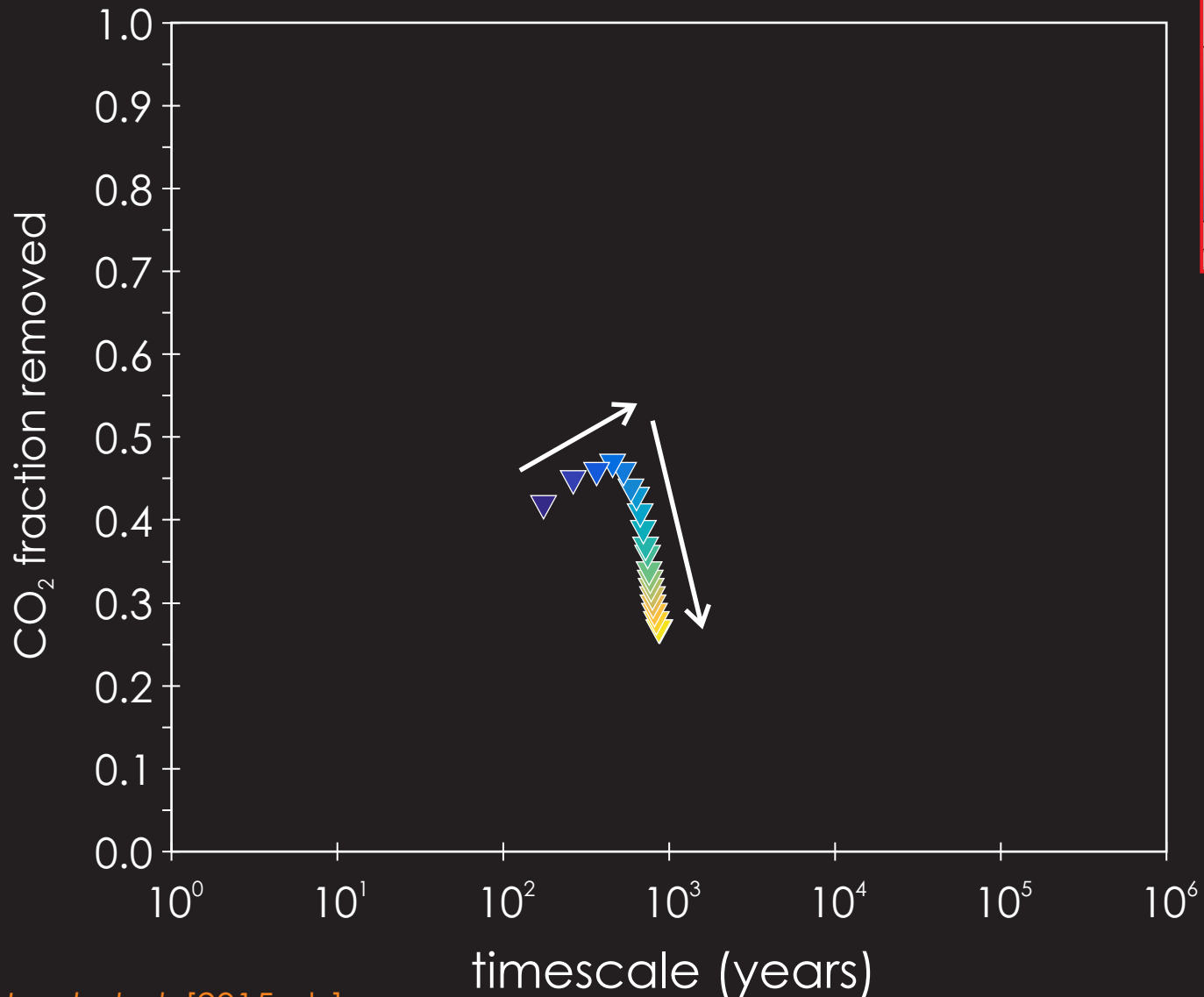


Depletion of mixed layer carbonate buffer;  
ocean stratification and reduced surface  
mixing. Warming and reduced  $\text{CO}_2$  solubility.





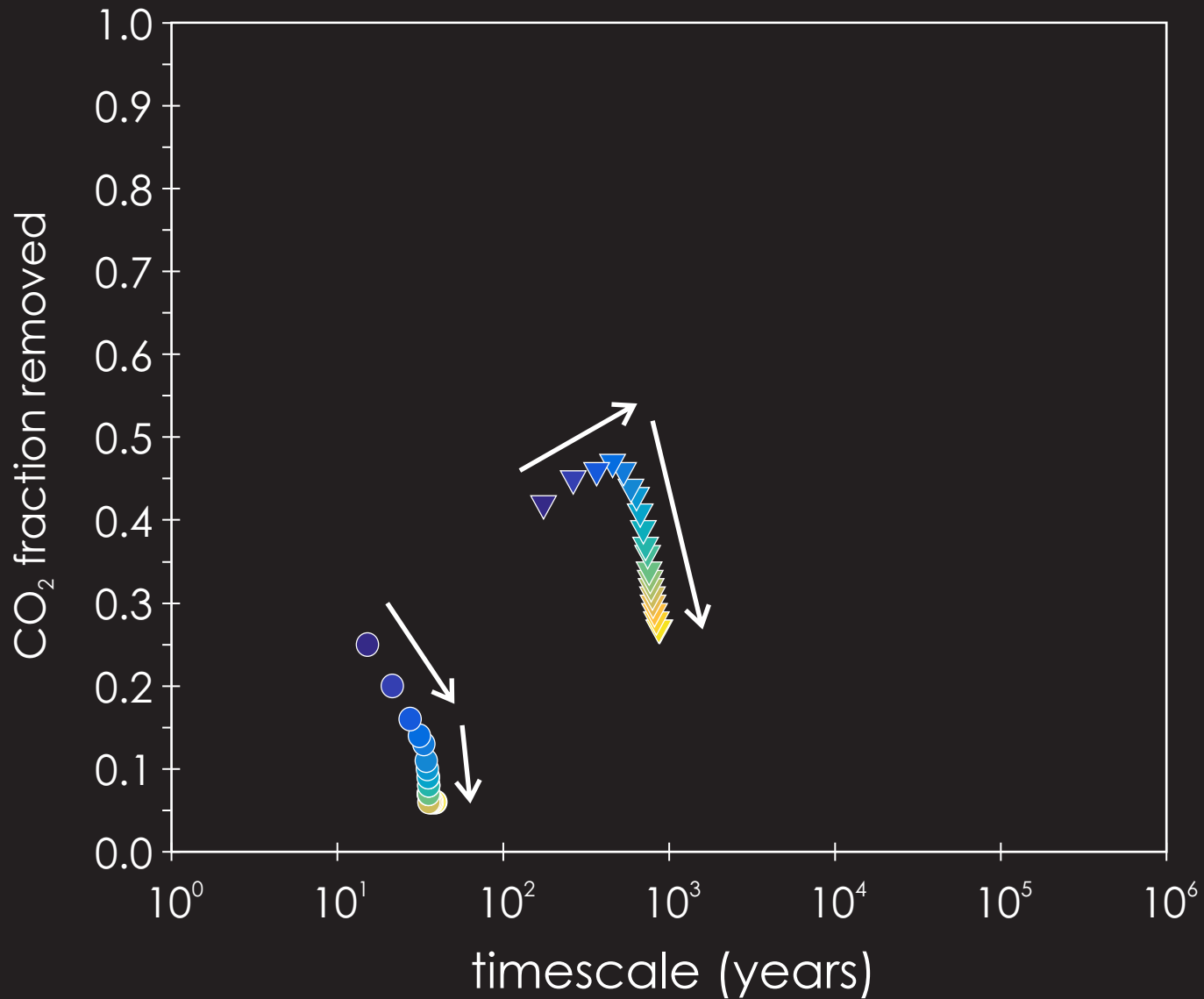
Ocean stratification and collapse of the AMOC  
(in this particular model).  
Threshold reached @  $\sim 4000 \text{ PgC}$ ?



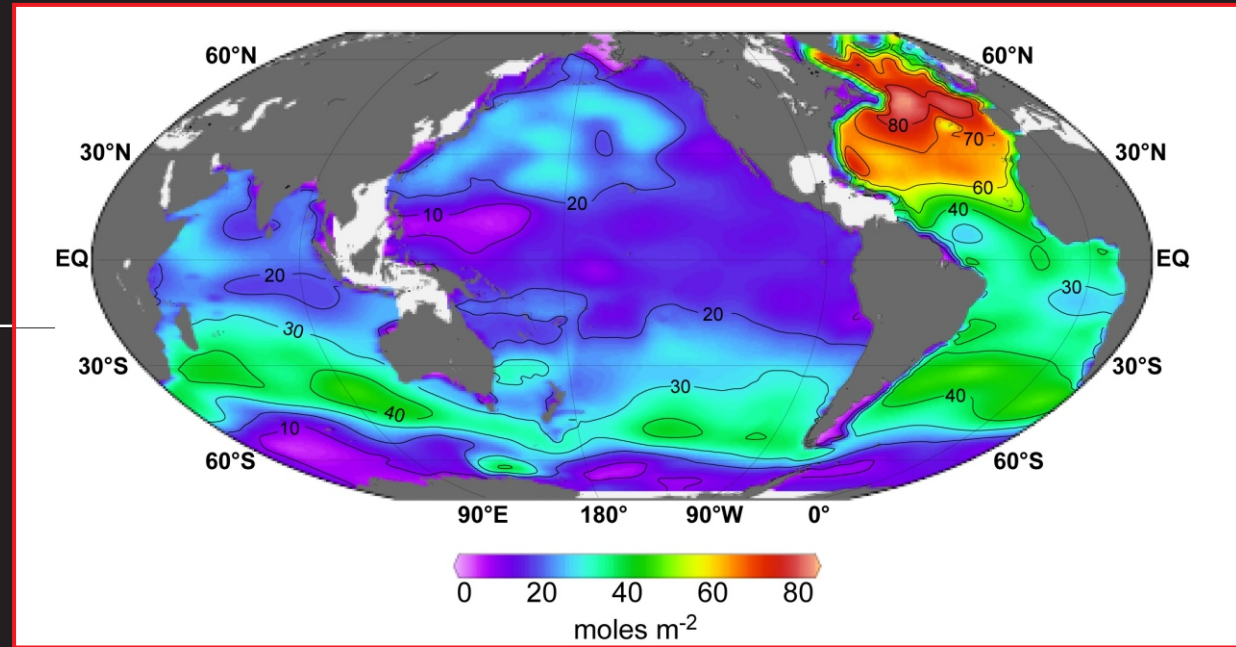
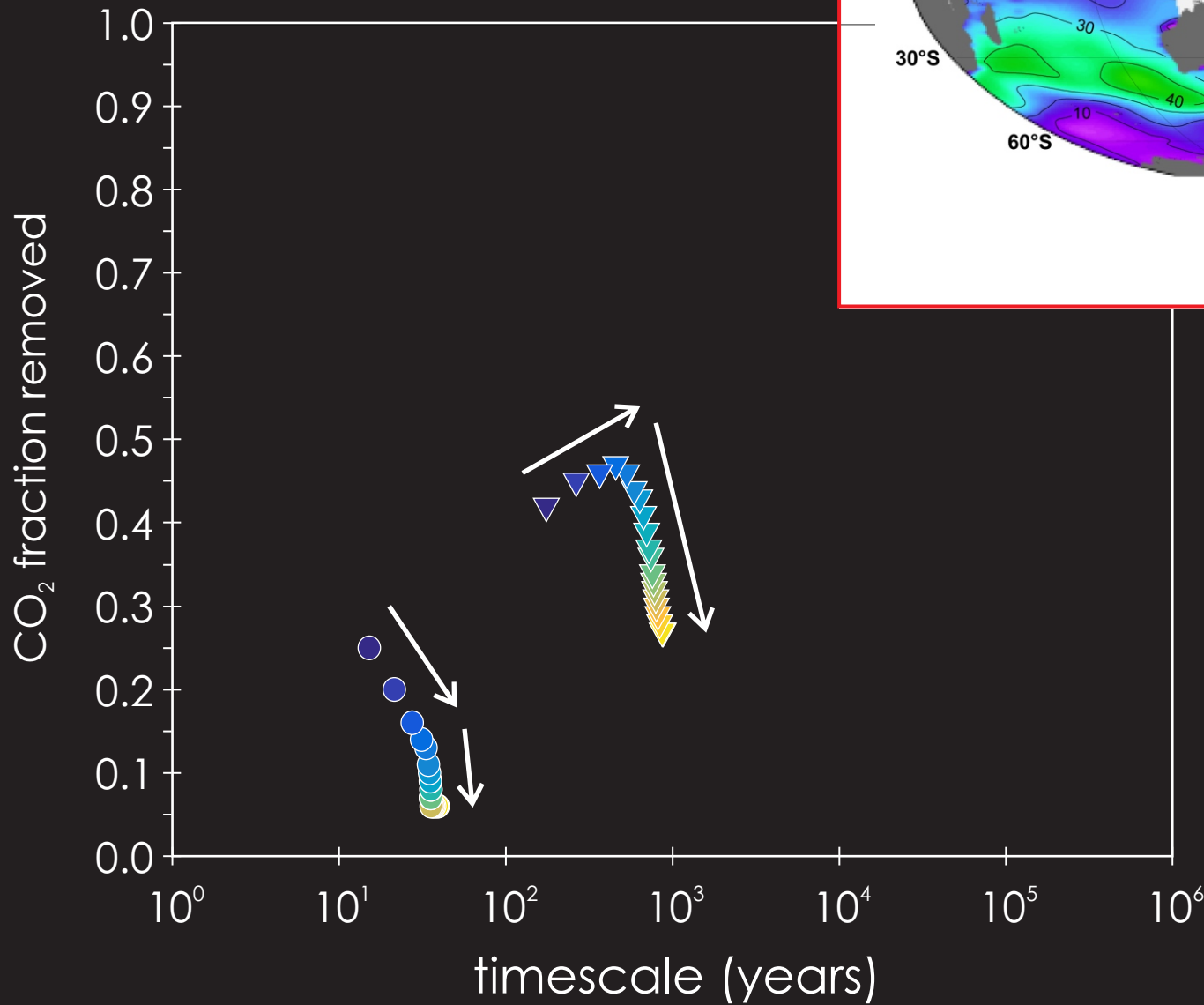




evidence?

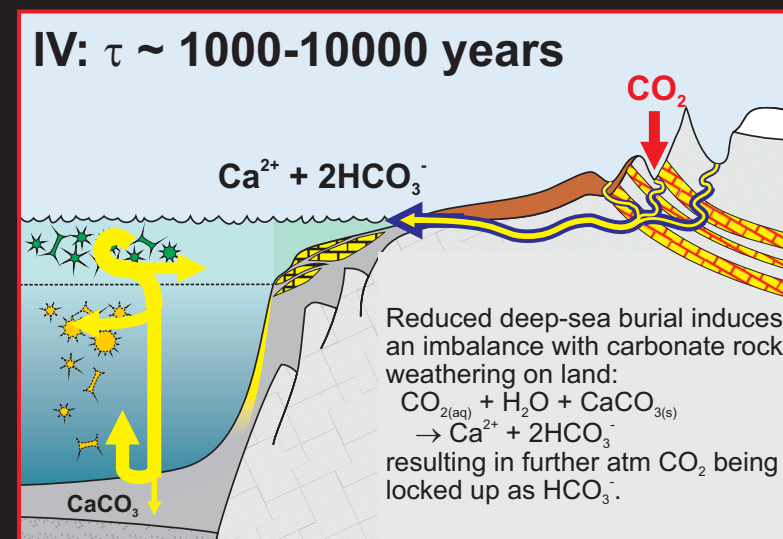
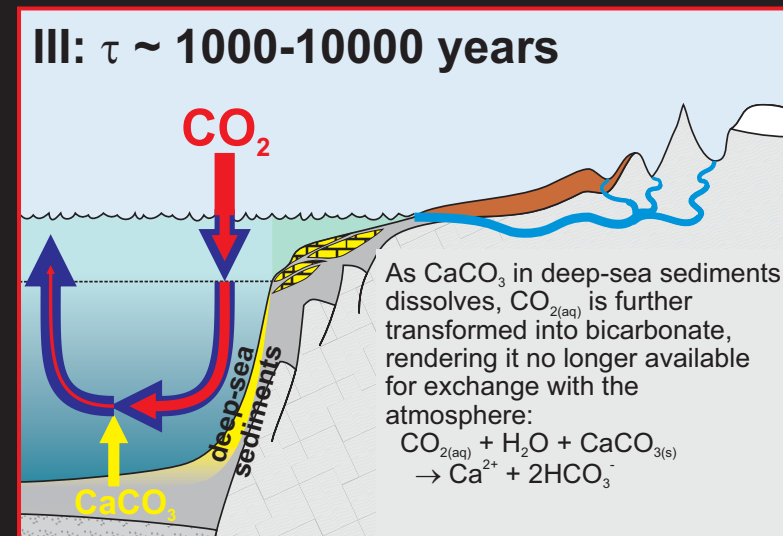
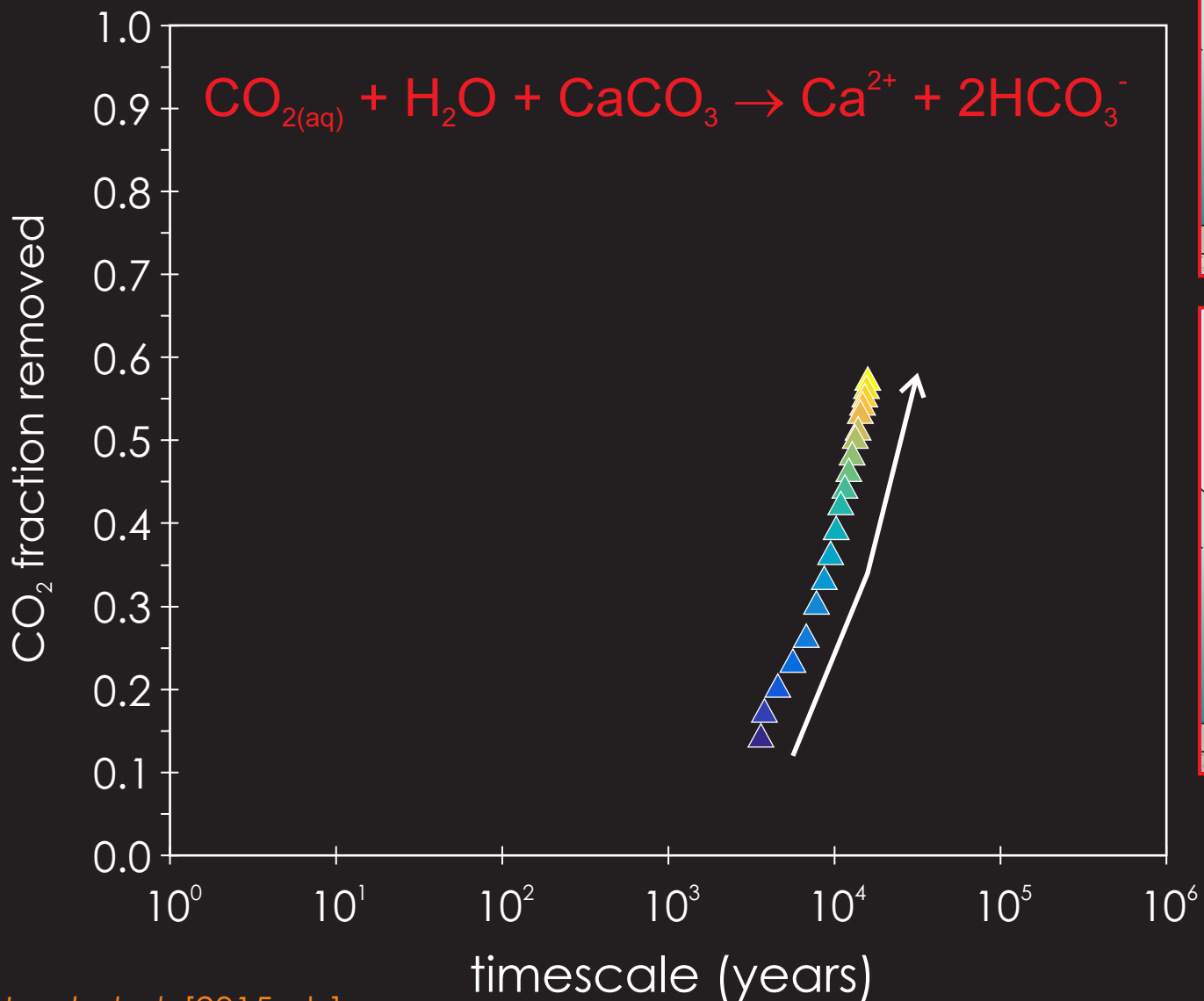


# Impulse response function analysis of the 'long tail' of $\text{CO}_2(\text{excess})$



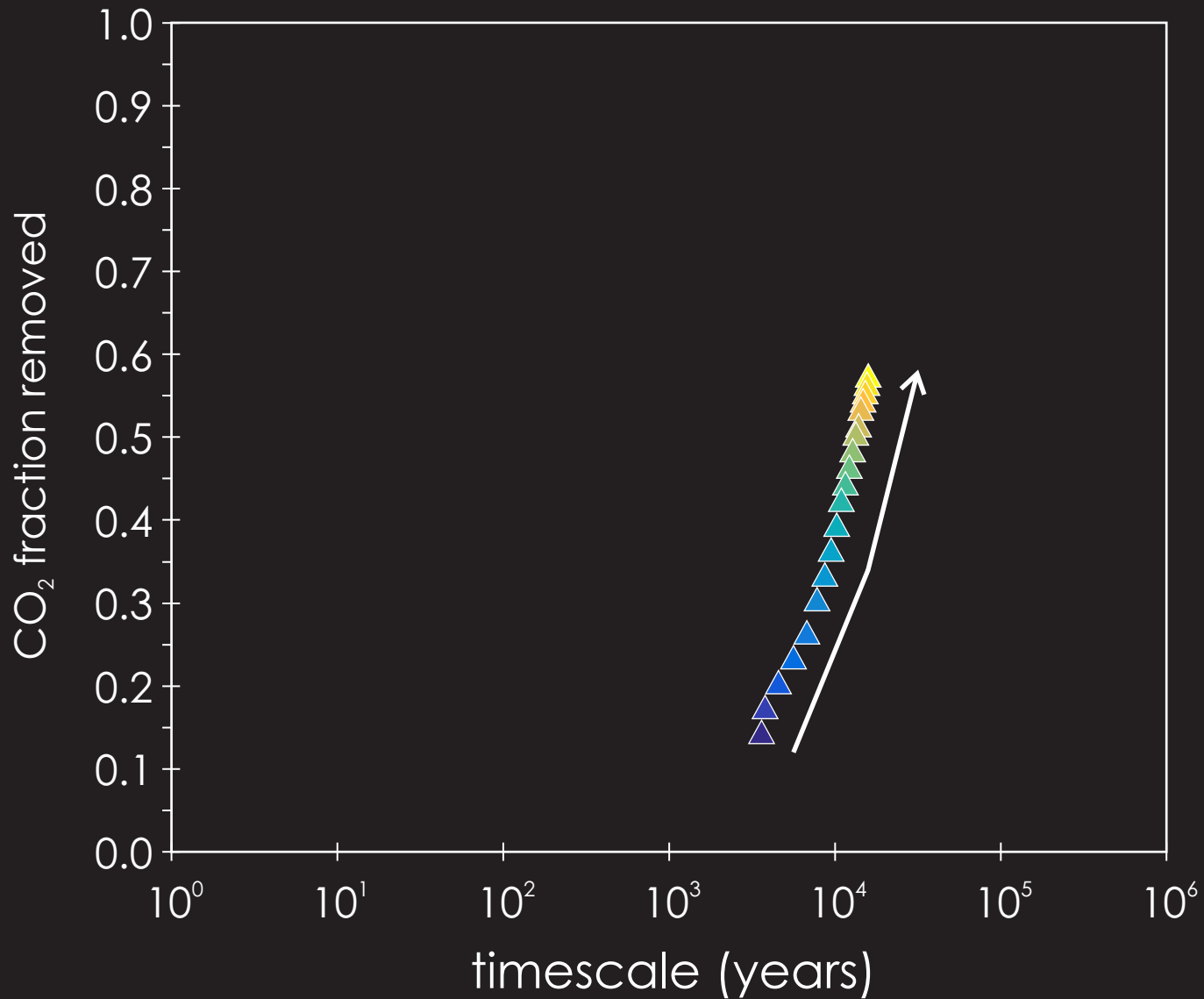


Geologic  $\text{CO}_2$  removal via carbonate rocks and marine sediments – occurring on an increasing protracted time-scale.





evidence?

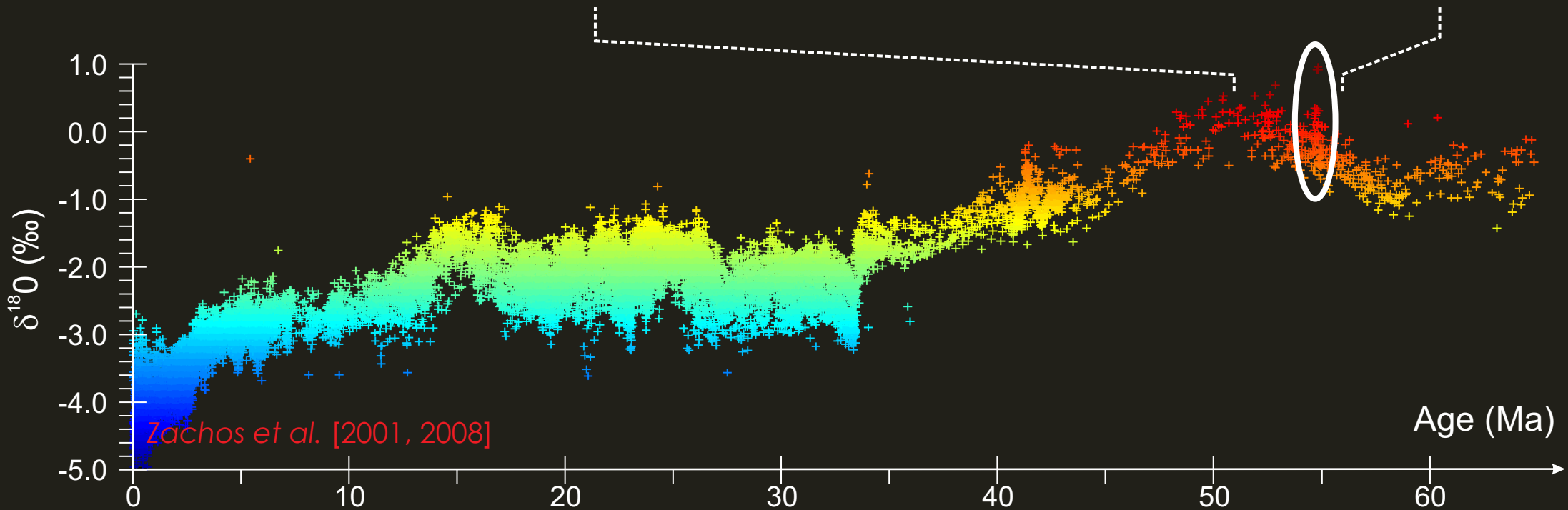
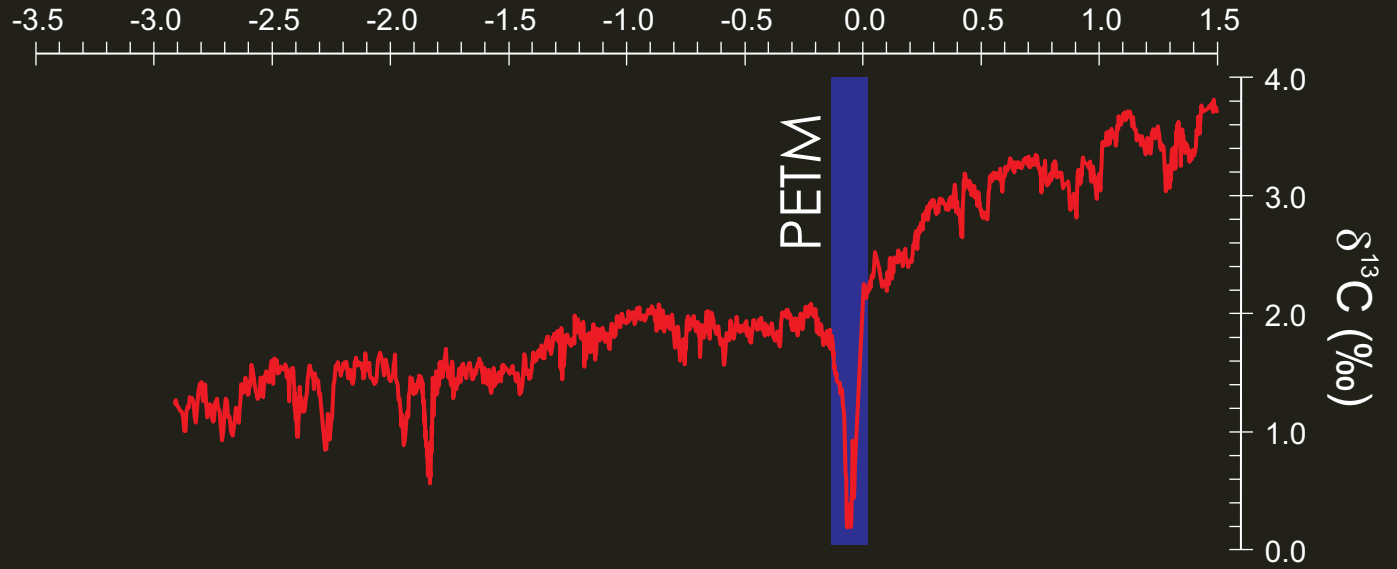


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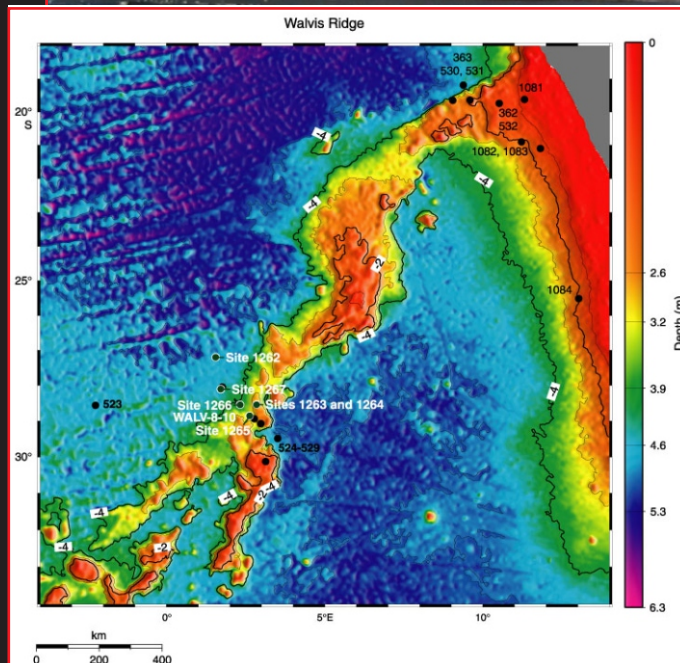
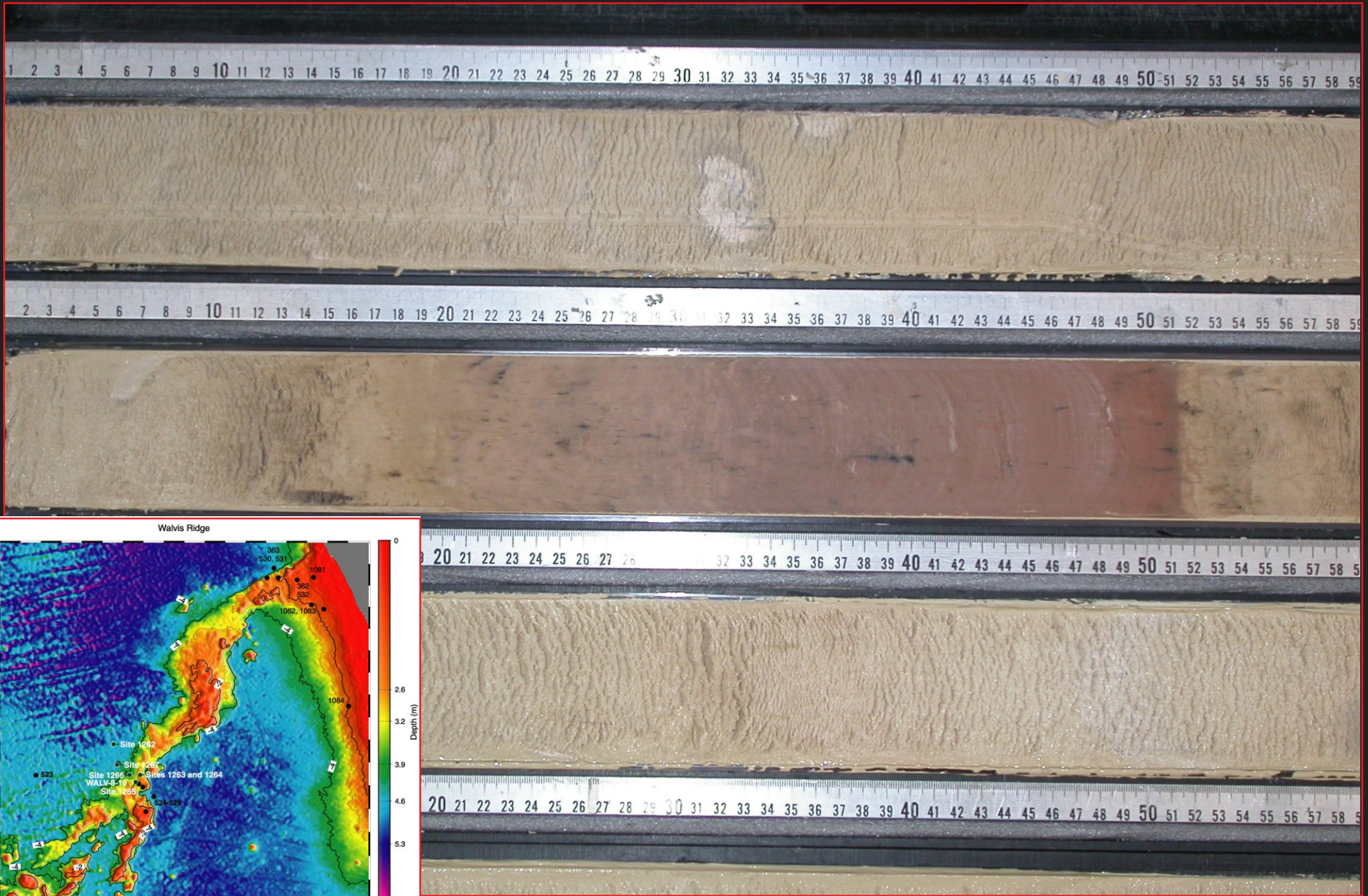
Zachos et al. [2010]  
Lunt et al. [2011]

Age relative to the PETM (Ma)



Zachos et al. [2001, 2008]

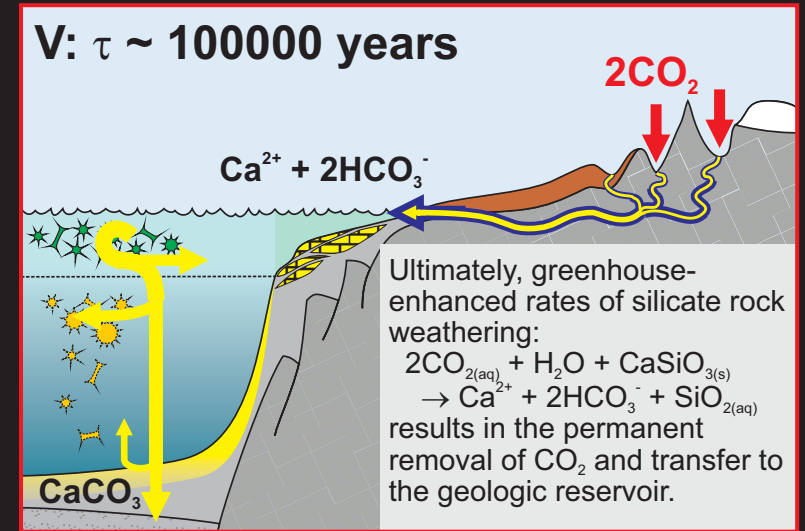
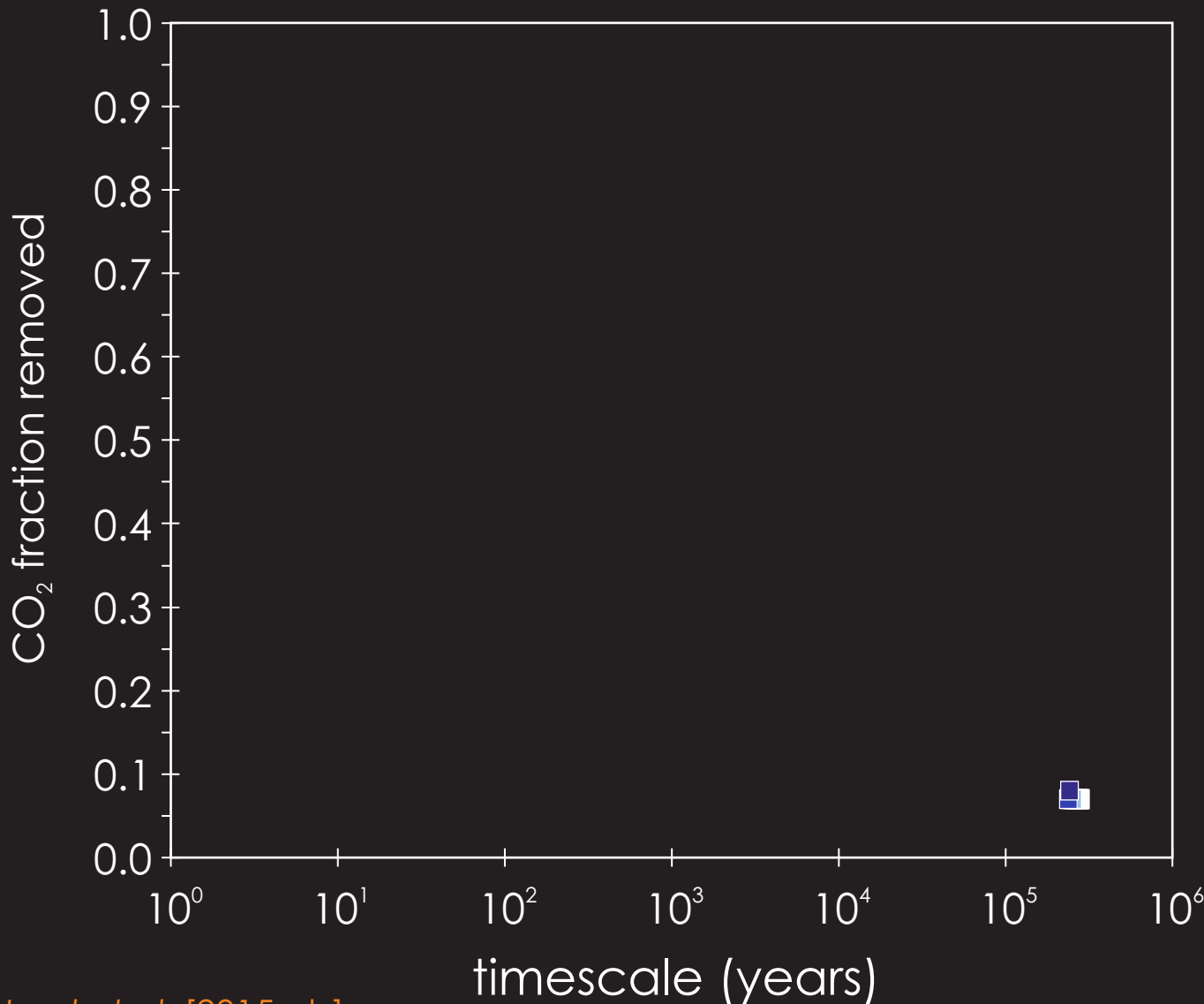
# Impulse response function analysis of the 'long tail' of $\text{CO}_2(\text{excess})$



Sediments spanning the Palaeocene-Eocene boundary from ODP Leg 208 (Walvis Ridge)  
Picture courtesy of Dani Schmidt (University of Bristol)

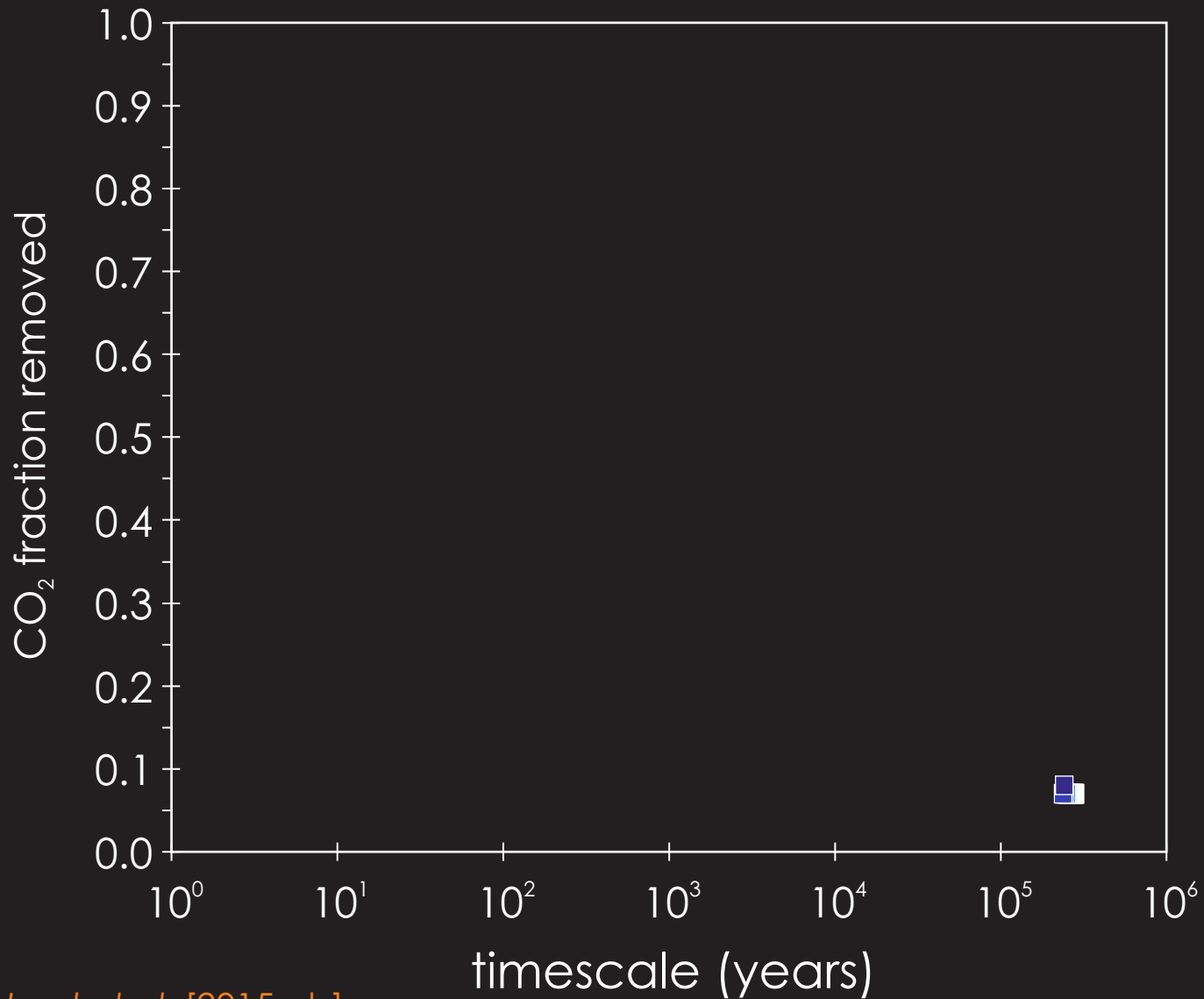


Silicate weathering (no time-scale response!).



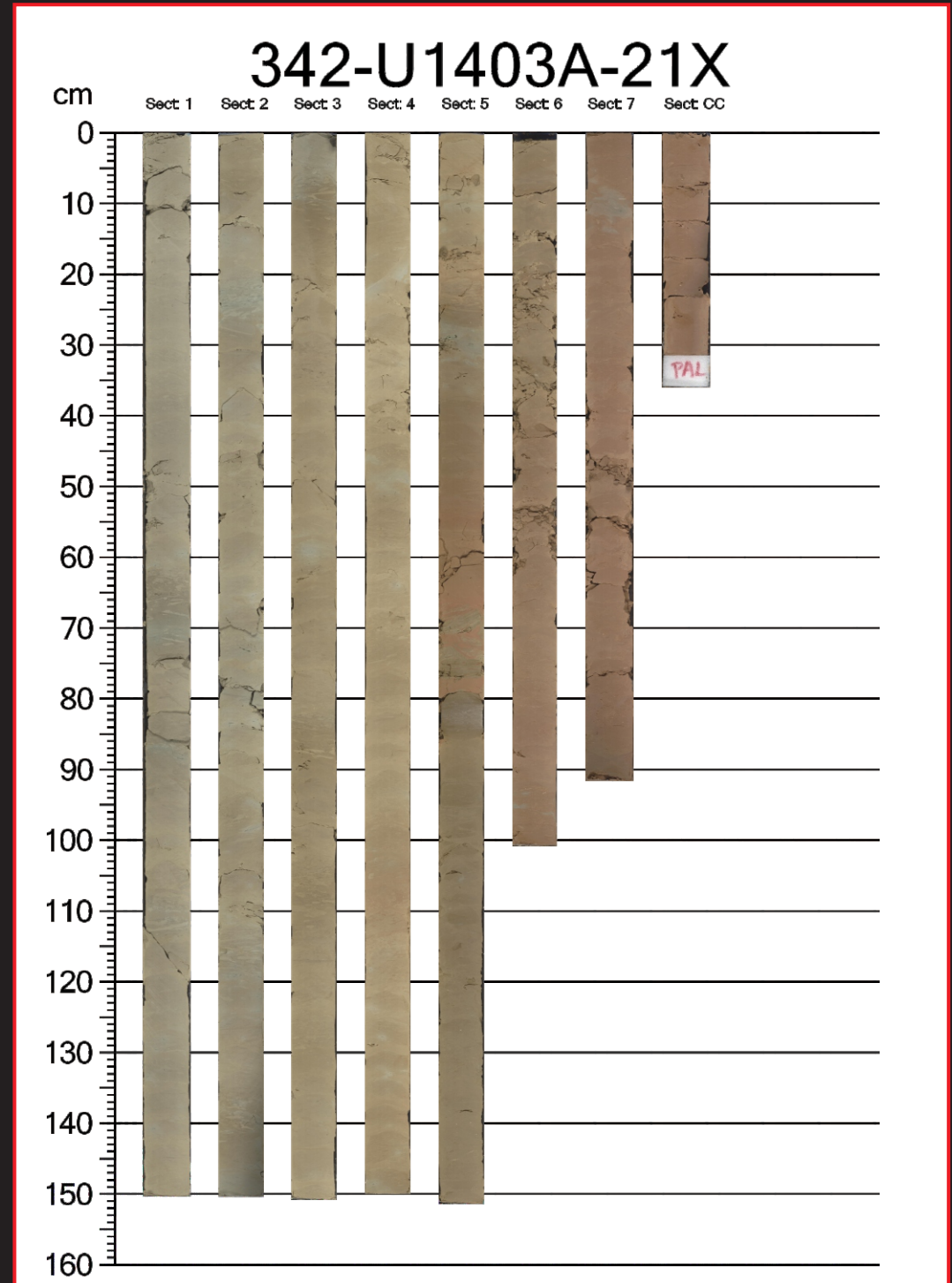
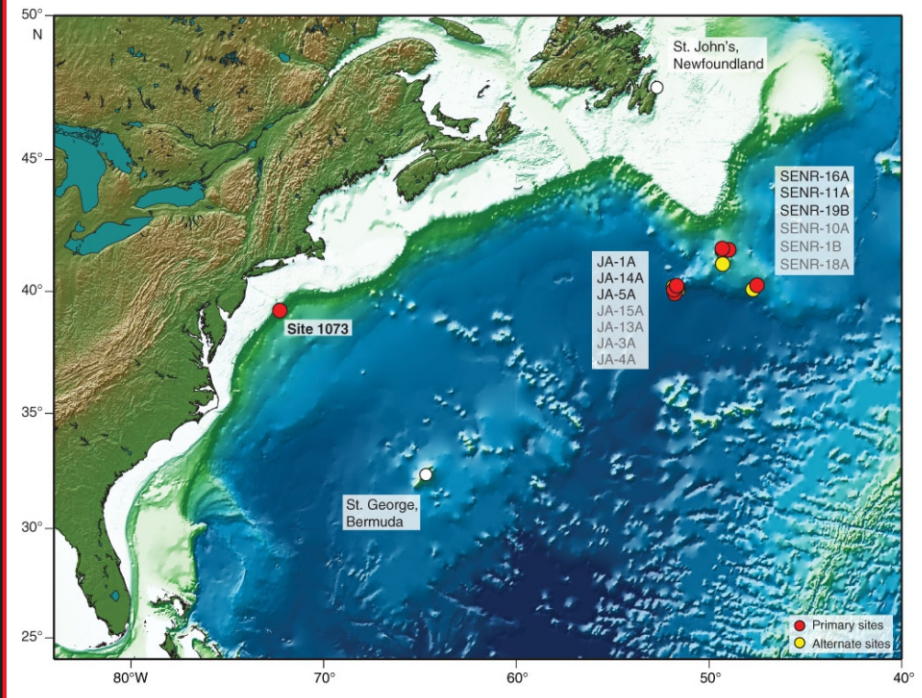


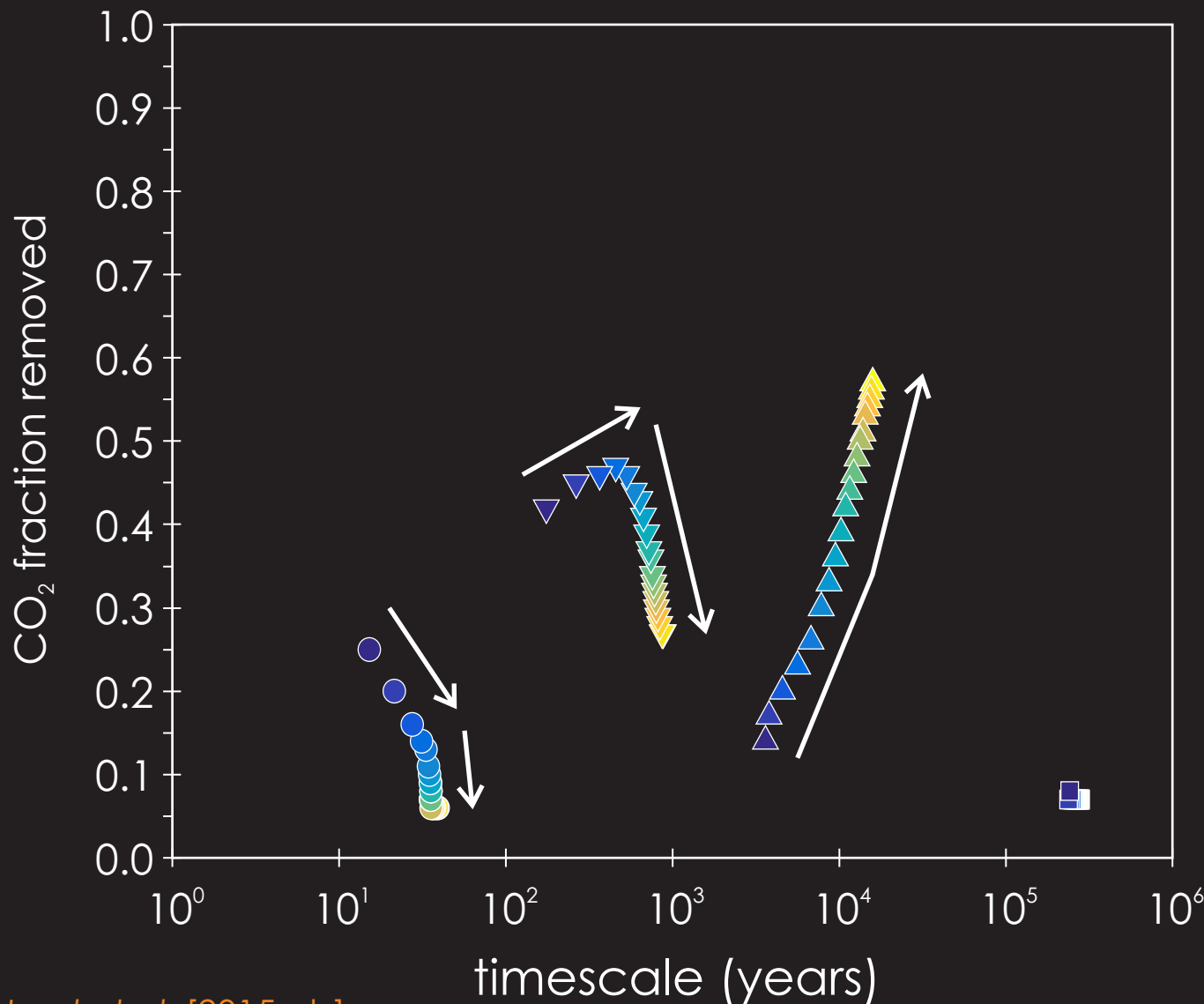
evidence?





# Impulse response function analysis of the 'long tail' of $CO_{2(excess)}$





With increasing total  $\text{CO}_2$  emissions, the response time of all sinks (bar silicate weathering) lengthen, and the shorter time-scale two weaken at the expense of the ~10,000 year  $\text{CaCO}_3$  burial process.

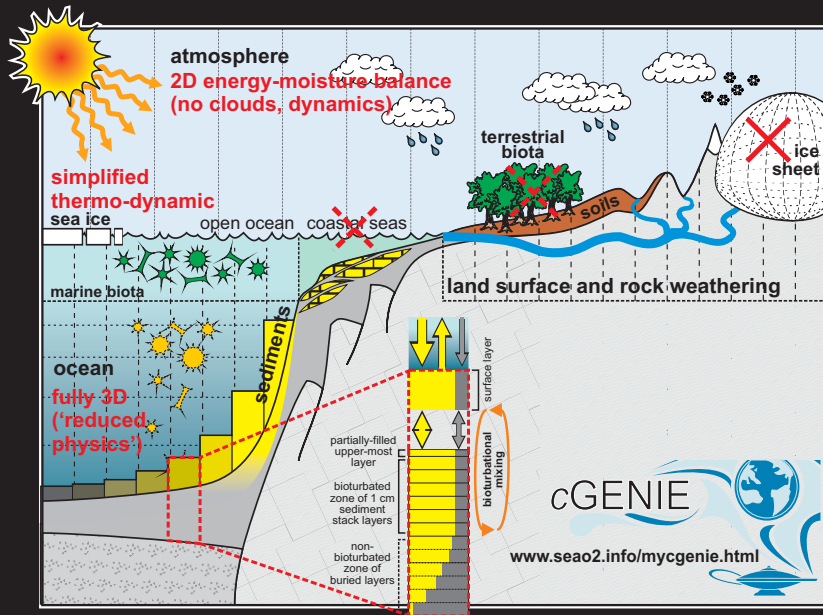
Elevated atmospheric  $p\text{CO}_2$  hence becomes more persistent as the main short-term  $\text{CO}_2$  feedbacks weaken.

The majority of carbon removal beyond ~10,000 PgC is removed only on time-scales exceeding 10,000 years.

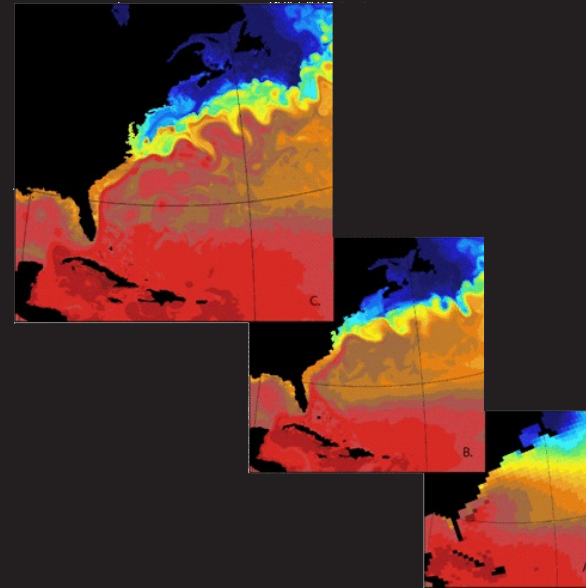
# Melting Antarctica



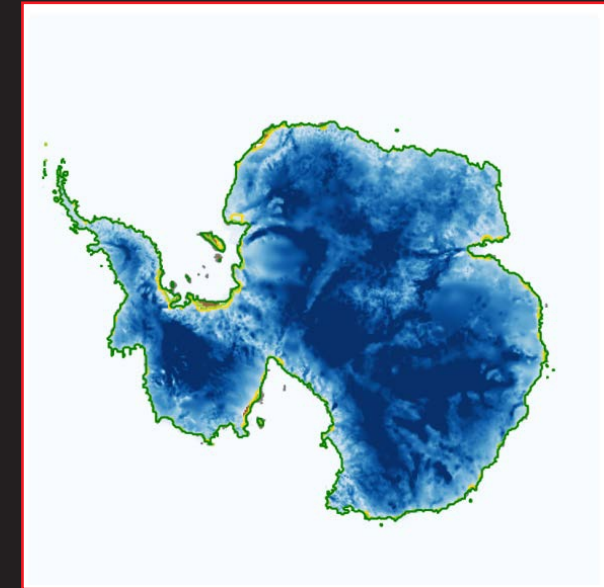
# Melting Antarctica



Earth system model  
(CO<sub>2</sub> and mean SST trajectories)

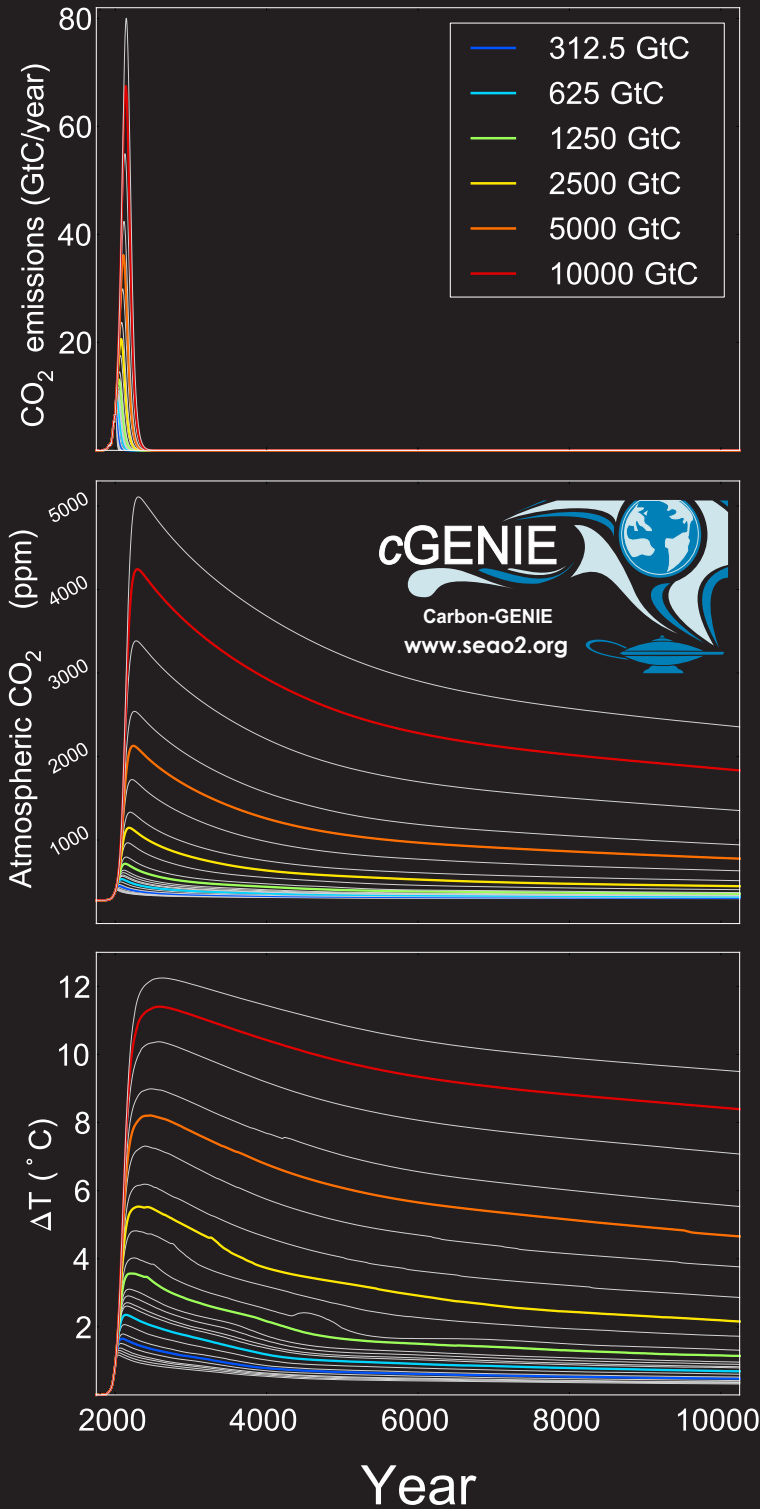


Downscaling  
(SO SST and regional climate)

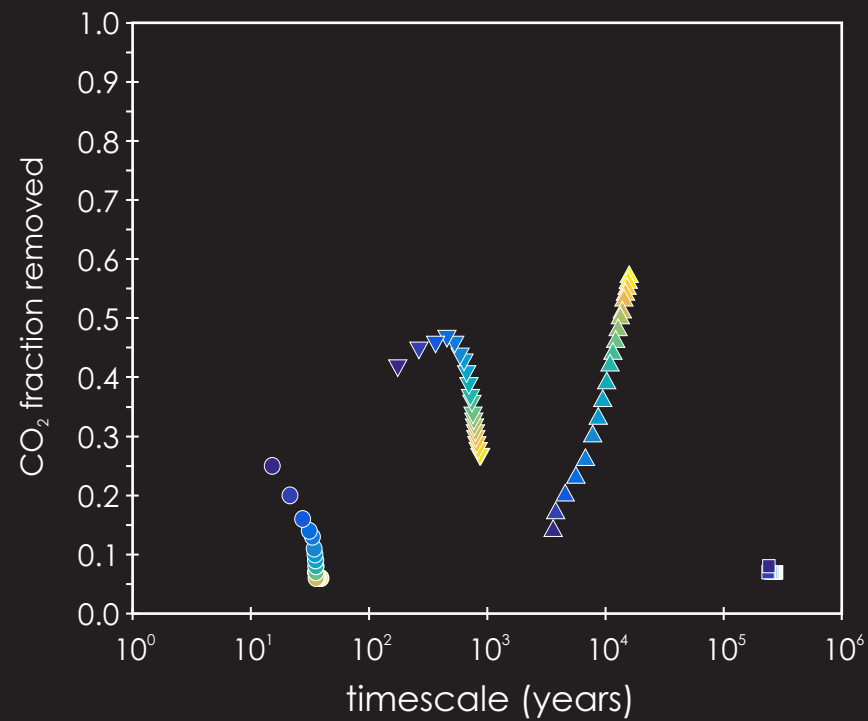
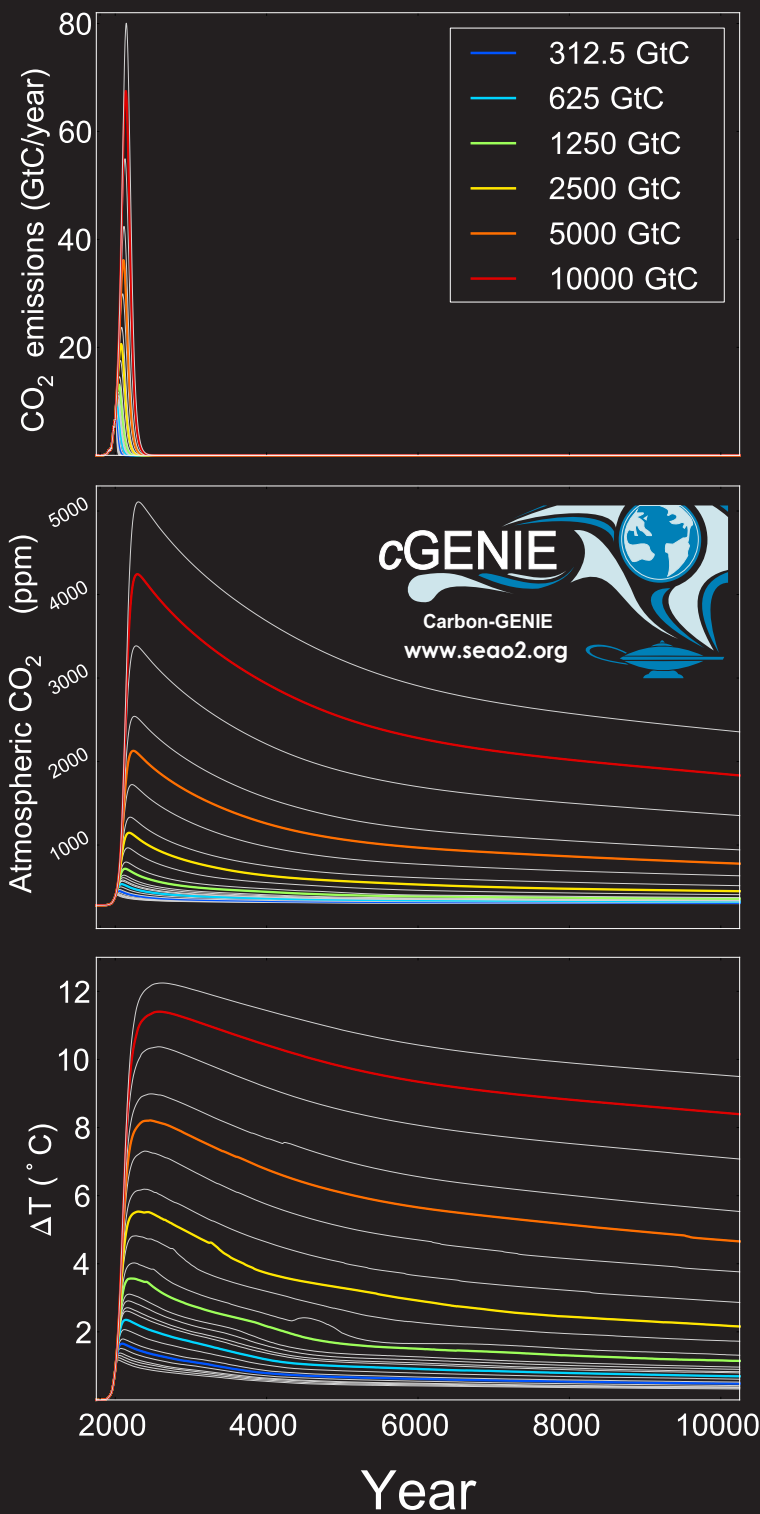


Ice sheet model

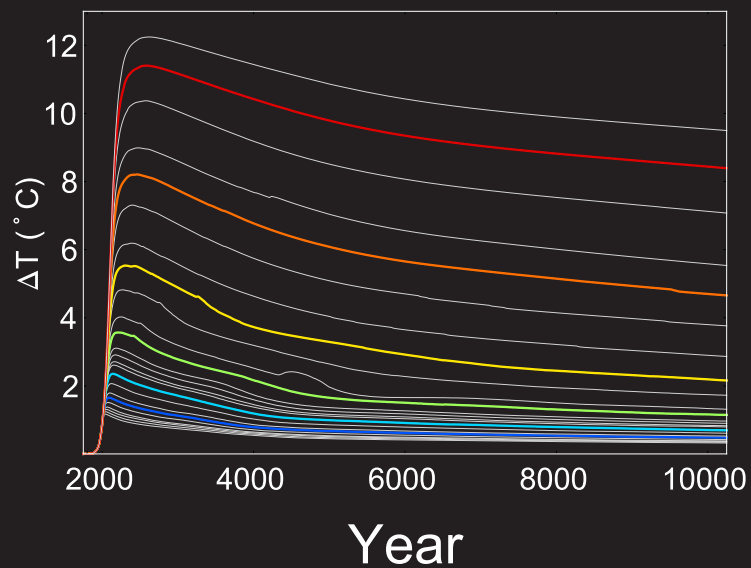
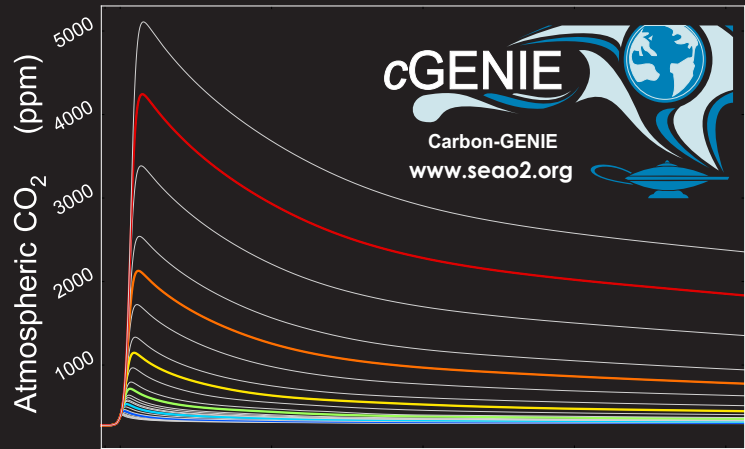
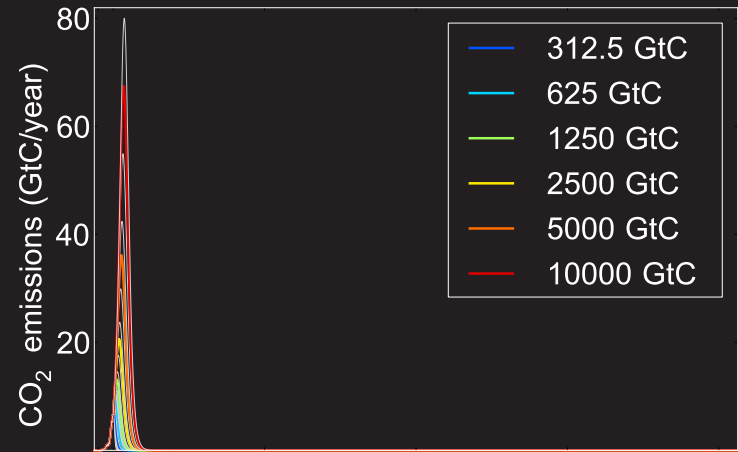
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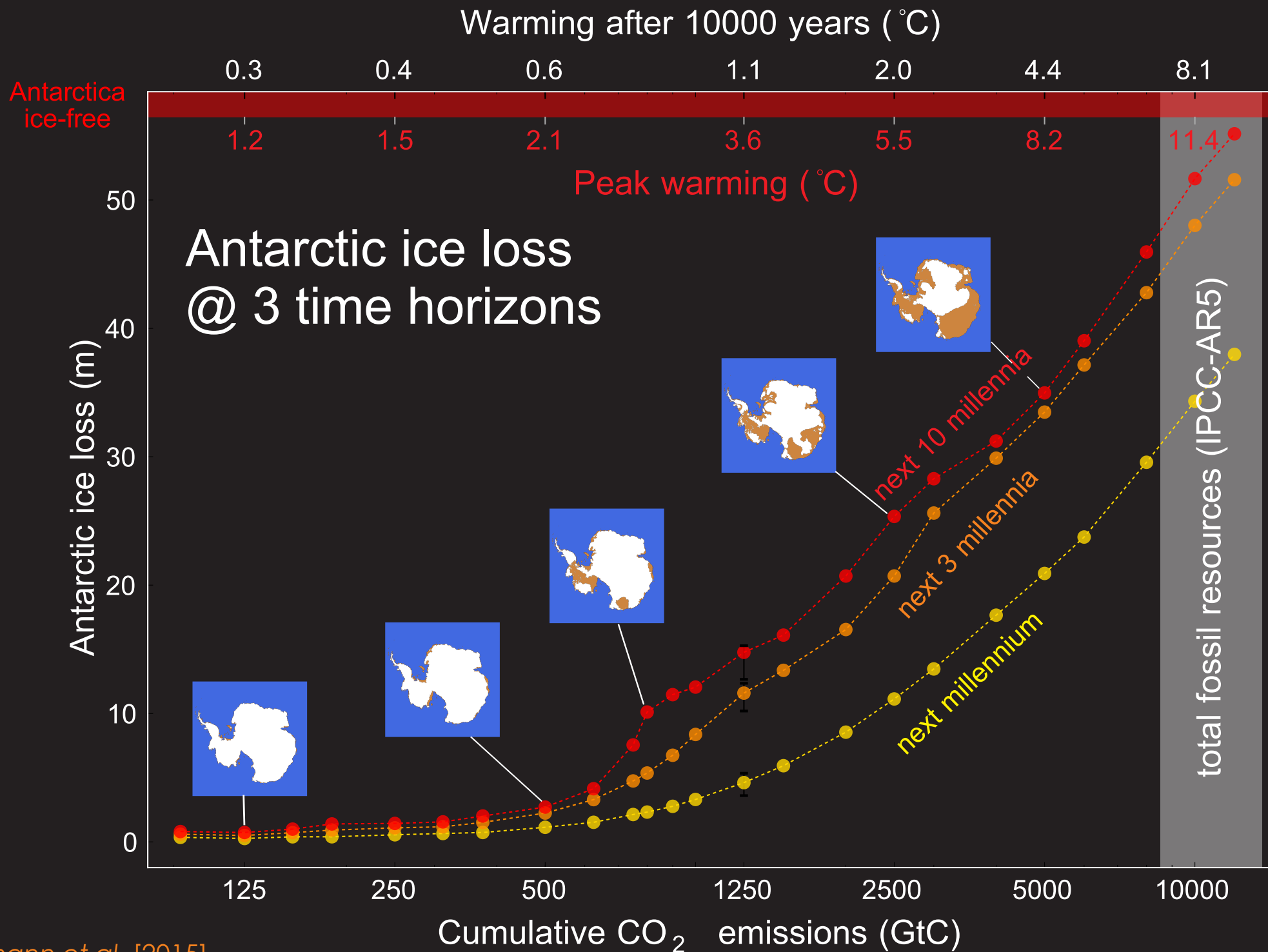


# Melting Antarctica



↔  $\Delta F \propto \ln(C/C_0)$

# Melting Antarctica





# Enhanced weathering (CO<sub>2</sub> removal geoengineering)



# Enhanced weathering (CO<sub>2</sub> removal geoengineering)



# Enhanced weathering (CO<sub>2</sub> removal geoengineering)





## granite ≈

SiO<sub>2</sub> = 72%

...

CaO = 1.8%

...

MgO = 0.7%

...

## basalt ≈

SiO<sub>2</sub> = 50%

...

CaO = 10%

...

MgO = 10%

...

# Enhanced weathering (CO<sub>2</sub> removal geoengineering)



~ plagioclase + pyroxene (+olivine)



# Enhanced weathering (CO<sub>2</sub> removal geoengineering)



~ olivine + pyroxene

Harzburgite

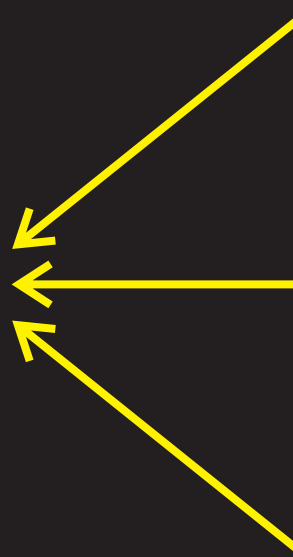


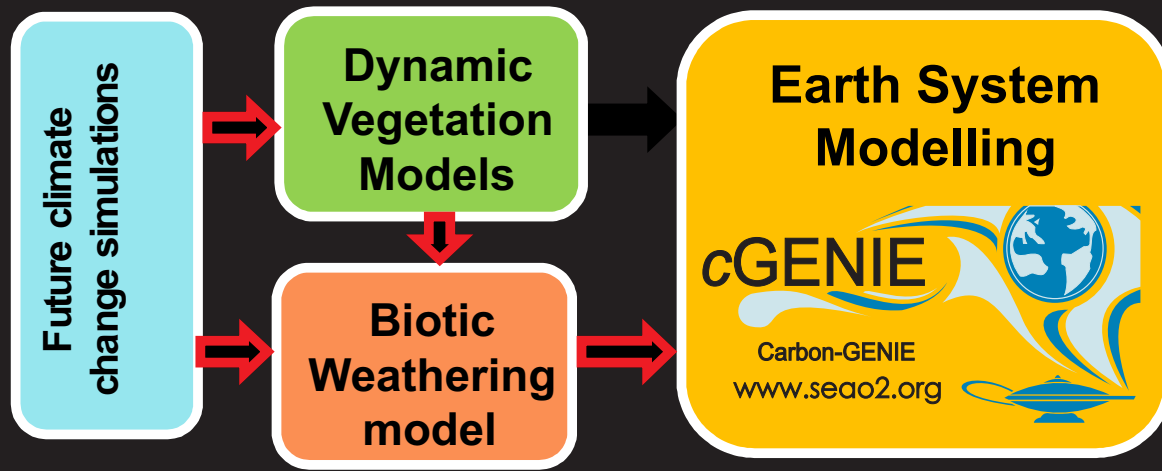
~ plagioclase + pyroxene (+olivine)



>90% olivine:  $(Mg^{+2}, Fe^{+2})_2SiO_4$

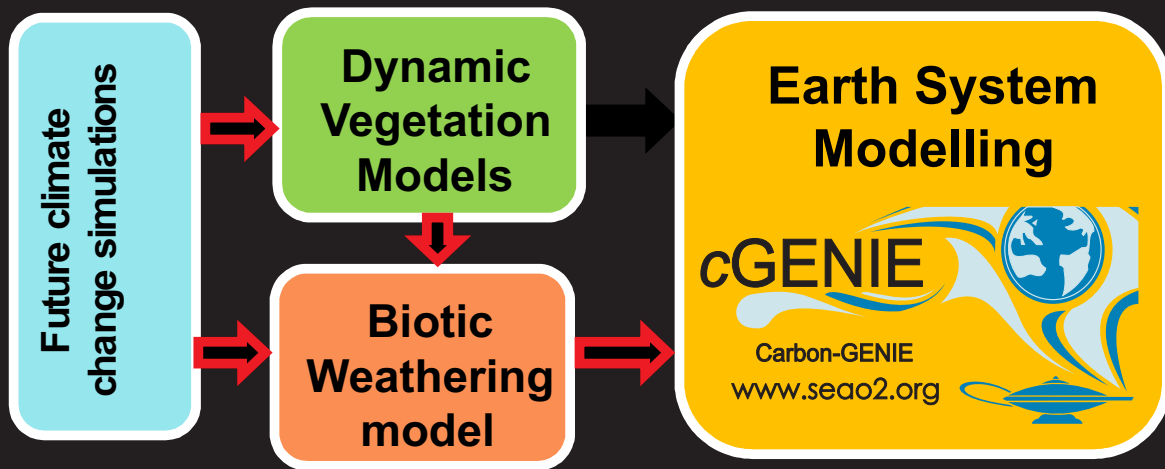
Dunite



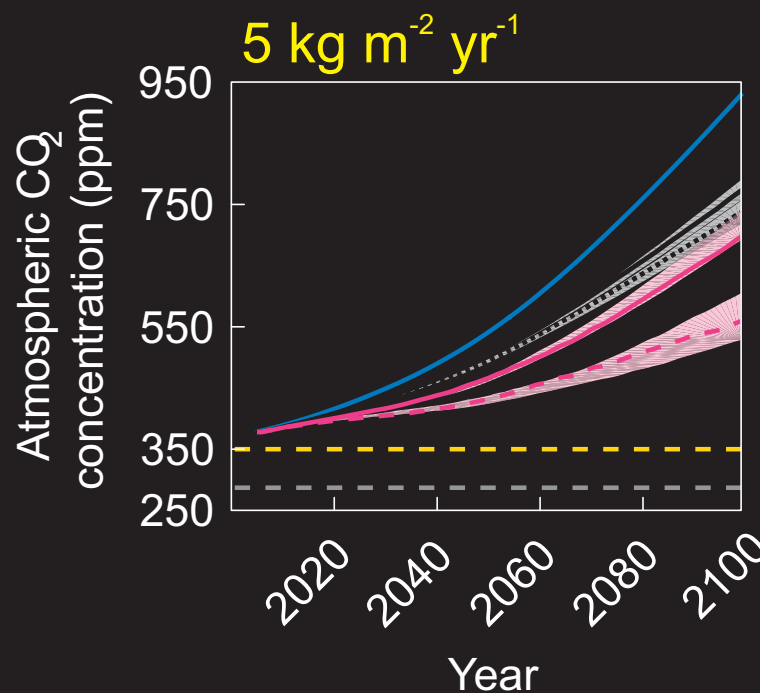
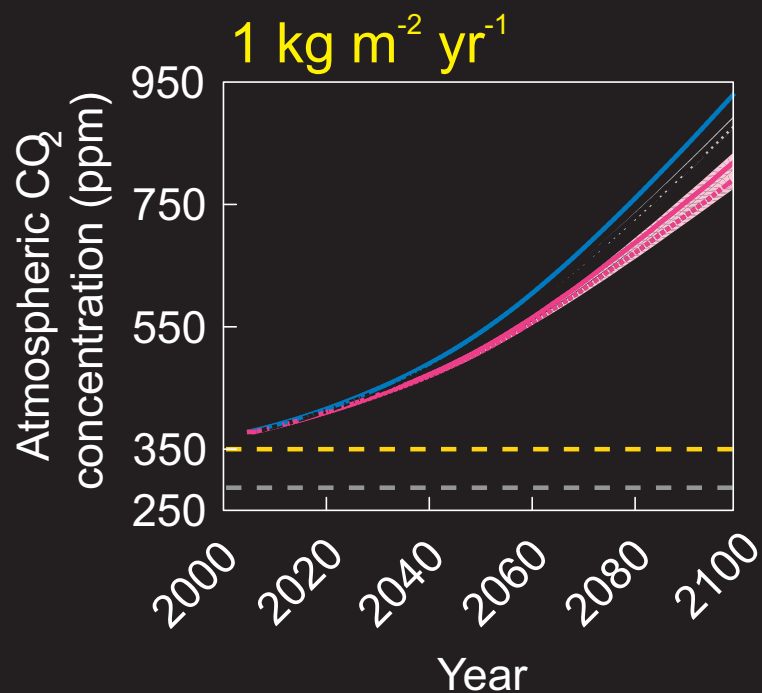


*Taylor et al. [2015]  
(Nature Climate Change)*

# Enhanced weathering (CO<sub>2</sub> removal geoengineering)



Taylor et al. [2015]  
(Nature Climate Change)

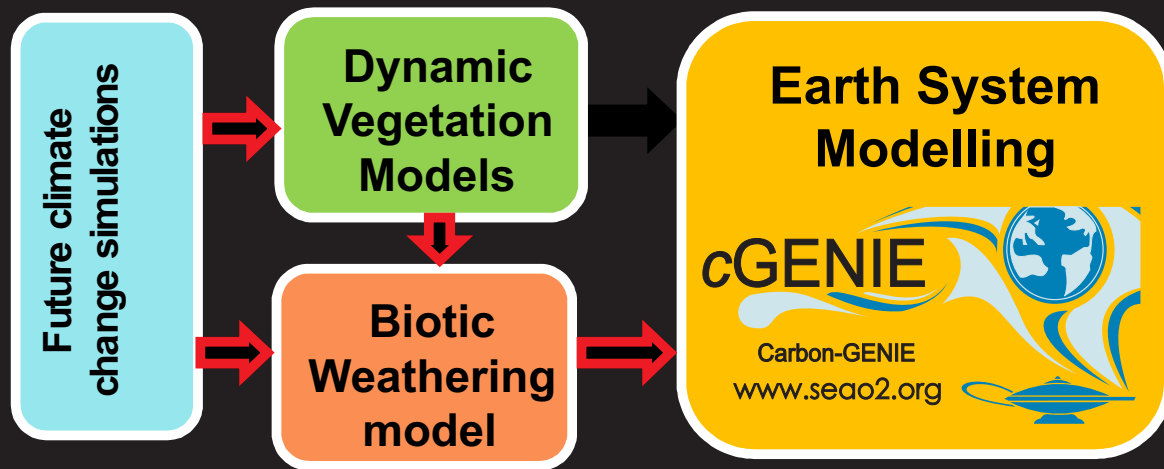


Ground rock application over 20×10<sup>3</sup> km<sup>2</sup> of tropical climatic 'hotspots'

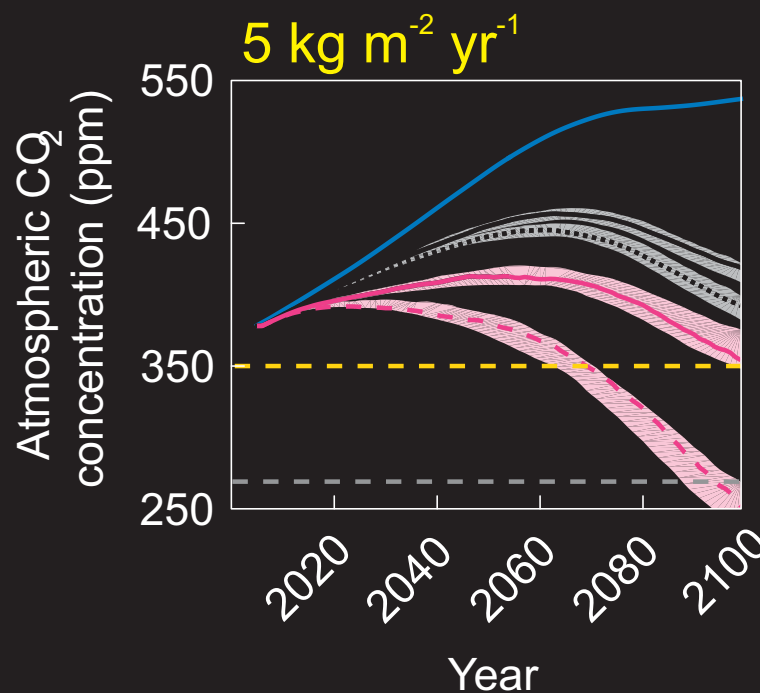
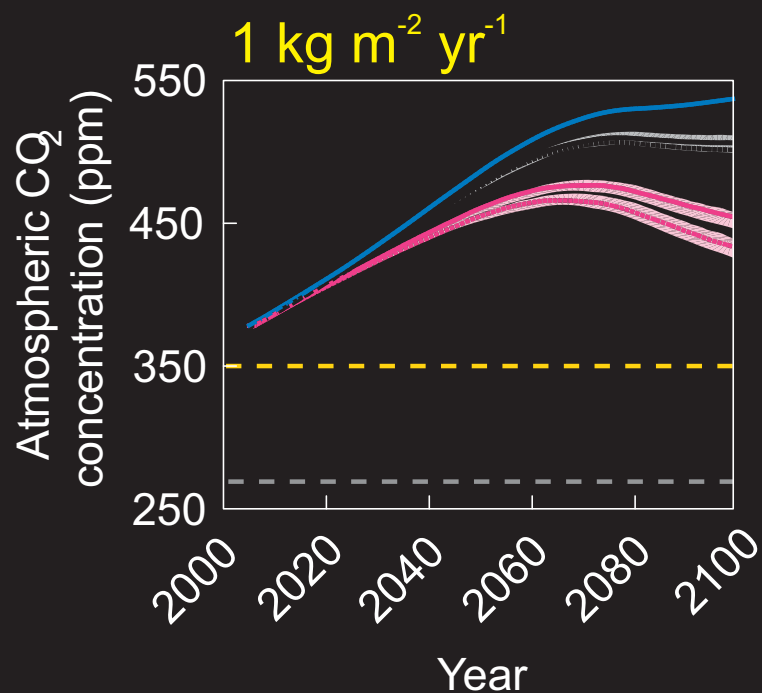
- RCP scenario
- Basalt 10 cm
- Basalt 30 cm
- Harzburgite 10 cm
- Harzburgite 30 cm
- - - 350 ppm CO<sub>2</sub>
- . . Pre-industrial CO<sub>2</sub>



# Enhanced weathering ( $\text{CO}_2$ removal geoengineering)



Taylor et al. [2015]  
(Nature Climate Change)



Ground rock application over  $20 \times 10^3 \text{ km}^2$  of tropical climatic 'hotspots'

- RCP scenario
- Basalt 10 cm    — Harzburgite 10 cm    - - - 350 ppm CO<sub>2</sub>
- Basalt 30 cm    — Harzburgite 30 cm    - - - Pre-industrial CO<sub>2</sub>



Current global oil  
consumption =  
90,136×10<sup>3</sup> barrels per  
day

$$\begin{aligned} 1.0 \text{ barrel} &= 159 \text{ l} \\ &= 159 \times 10^3 \text{ cm}^3 \end{aligned}$$

$$\begin{aligned} \Rightarrow \text{ oil consumption} \\ &= 5.23 \times 10^{15} \text{ cm}^3 \text{ year}^{-1} \\ &= \mathbf{5.23 \text{ km}^3 \text{ year}^{-1}} \end{aligned}$$

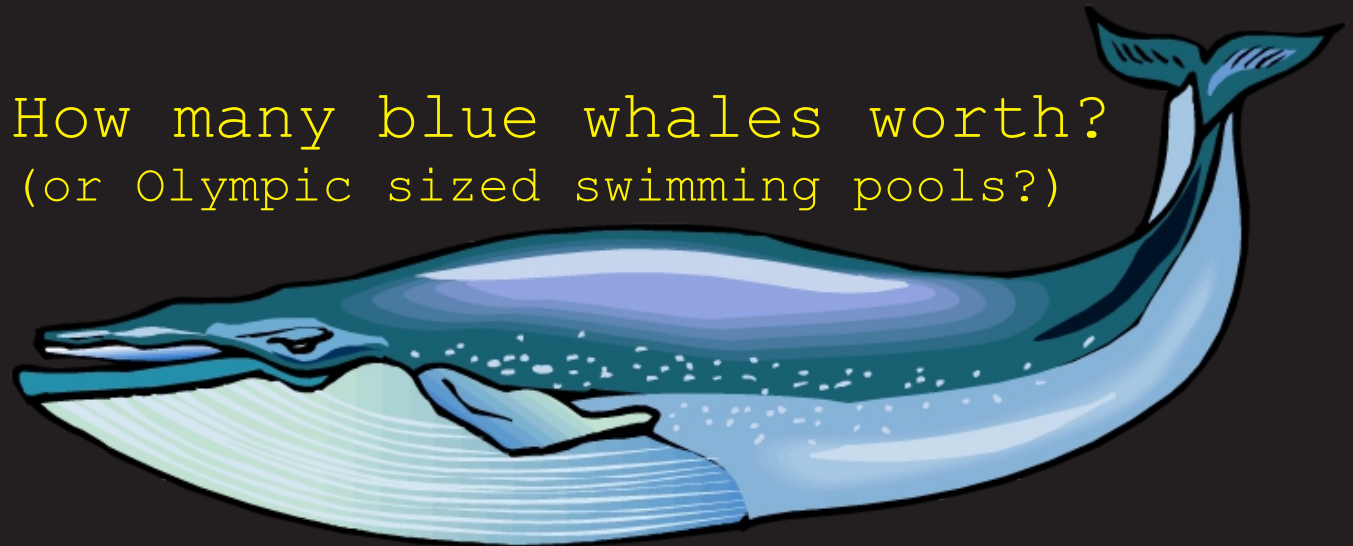


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⇒ oil consumption  
= 5.23×10<sup>15</sup> cm<sup>3</sup> year<sup>-1</sup>  
= **5.23 km<sup>3</sup> year<sup>-1</sup>**

How many blue whales worth?  
(or Olympic sized swimming pools?)





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= **5.23 km<sup>3</sup> year<sup>-1</sup>**

How many Yosemite Valleys?  
(equivalent volume)





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Yosemite Valley  
(Wikipedia):

1,200m deep × 1,600m  
across, 12.0 km long

⇒

$$\begin{aligned} \text{volume} &= 1.2 \times 1.6 \times 12.0 \\ &= \mathbf{23.0 \text{ km}^3} \end{aligned}$$

How many Yosemite Valleys?  
(equivalent volume)





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