Dynamics of the Cretaceous Oceans: A Numerical Recipe



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European Research Council



Outline

 "Take one whole fresh super-continent and break into pieces. Pick out the rock phosphate and place to one side.
Immerse the continental fragments in seawater until the shelves and interior seaways are thoroughly flooded."

* "Add a pinch of CO₂ and heat gently."

"While the ocean is warming and de-oxygenating, gradually stir in the phosphate that was put aside earlier.
Keep stirring and adding CO₂ and phosphate until a thick black carbon crust suddenly forms. Remove the crust.
Repeat to create as many carbon layers as possible before the cake starts to cool and the ocean re-oxygenates."

* The primary question is then ...

continental topology/ topography

temperature/ stratification

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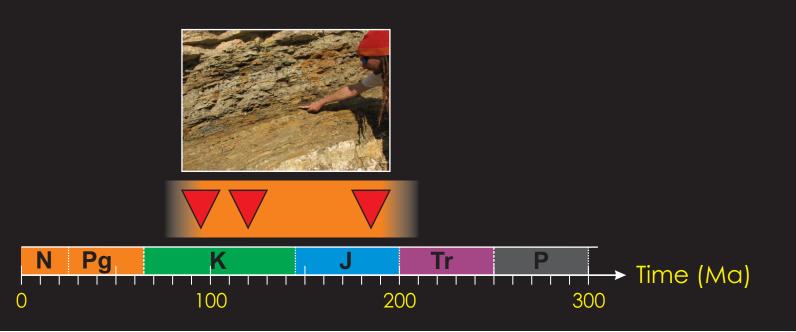
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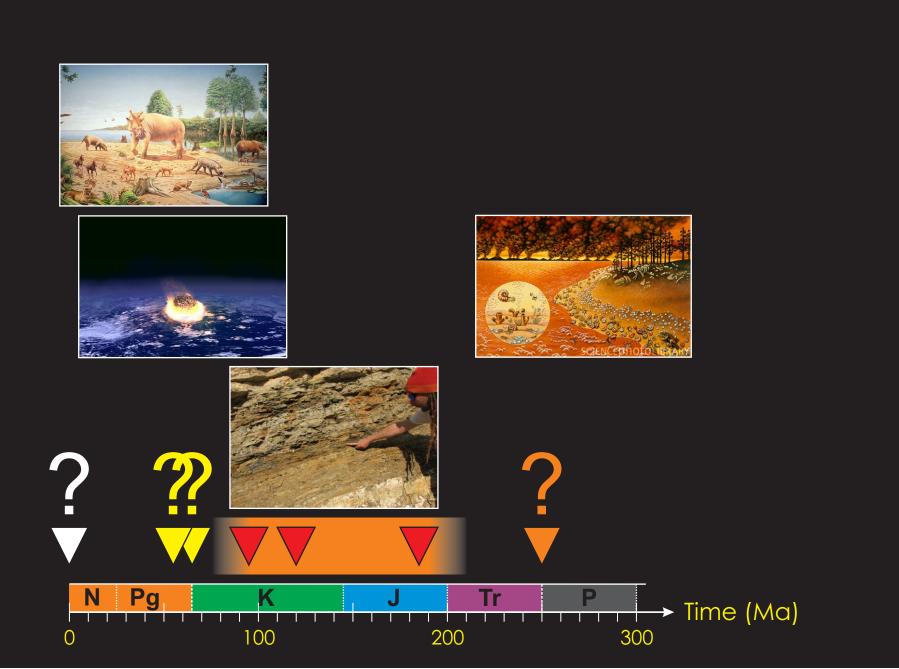


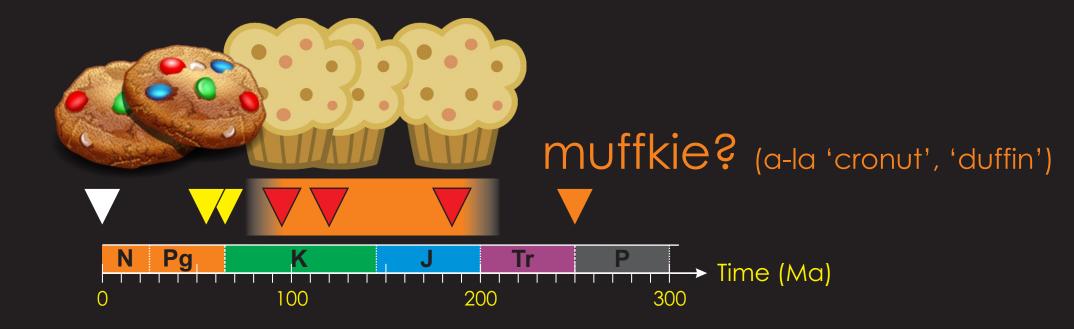
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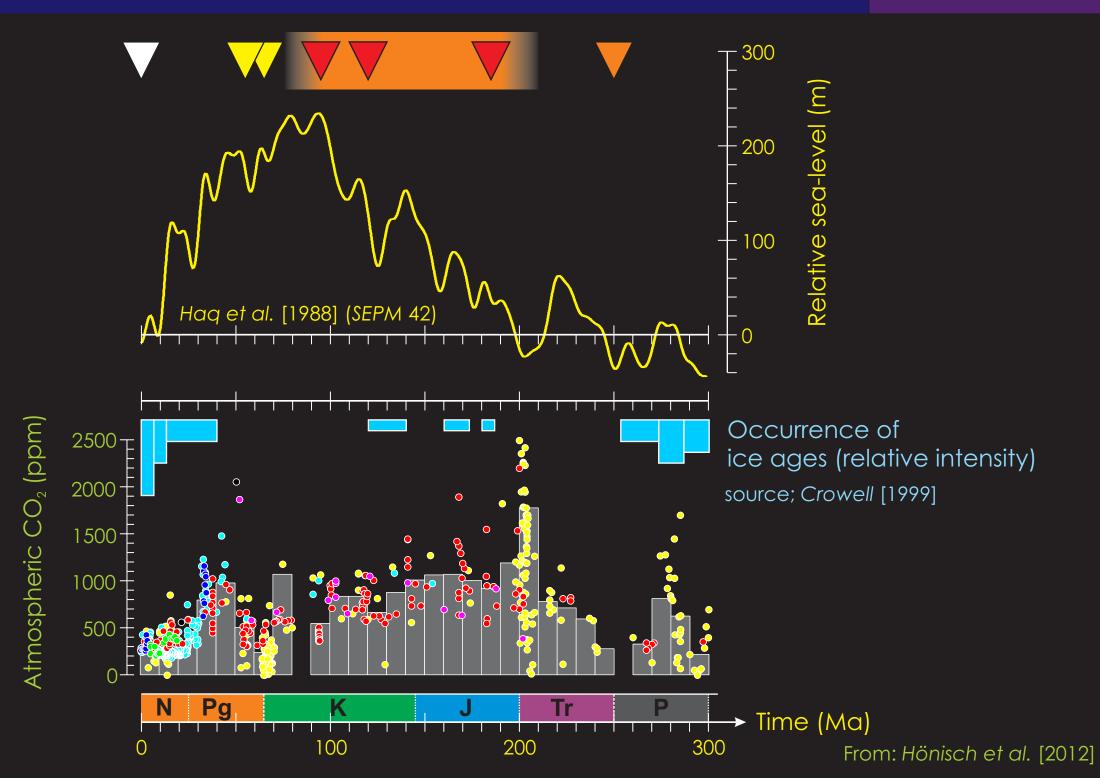


What is the 'recipe' for OAE (occurrence)?

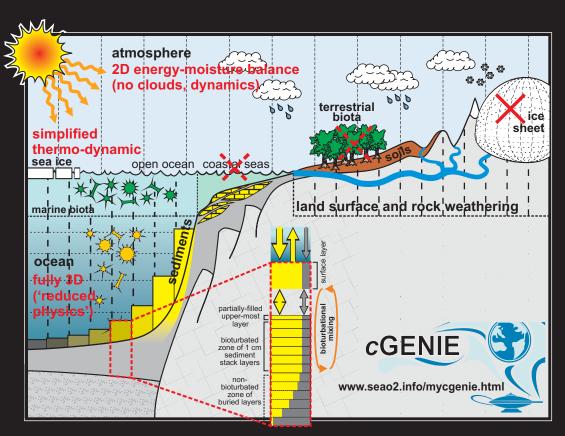




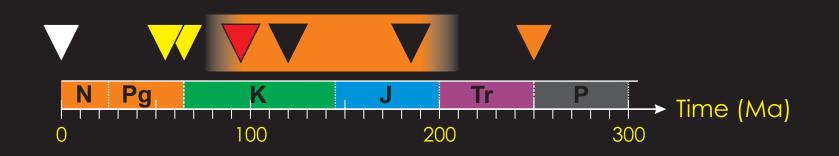
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Carbon cycle and oxygen perturbations



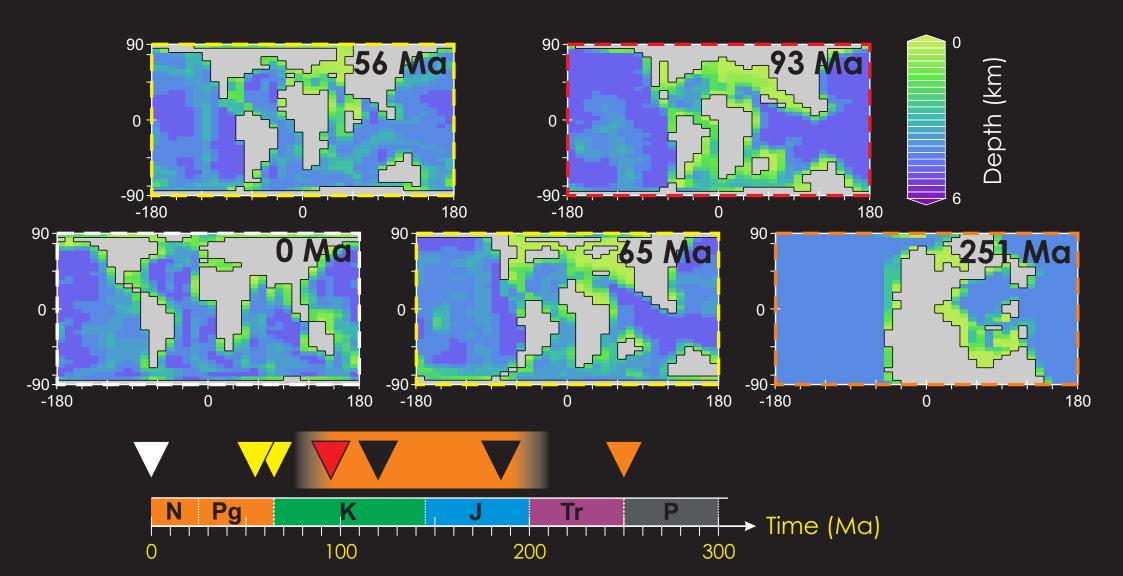
'cGENIE' Earth system model ('of Intermediate Complexity')



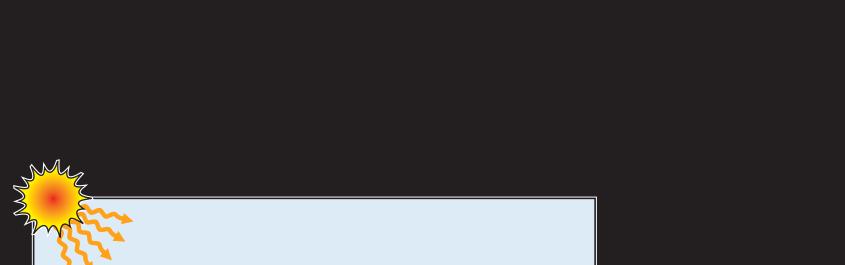
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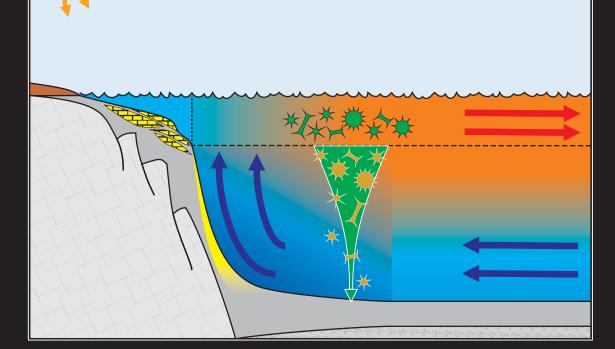
ocean bathymetry & continental configuration

#endrainbow



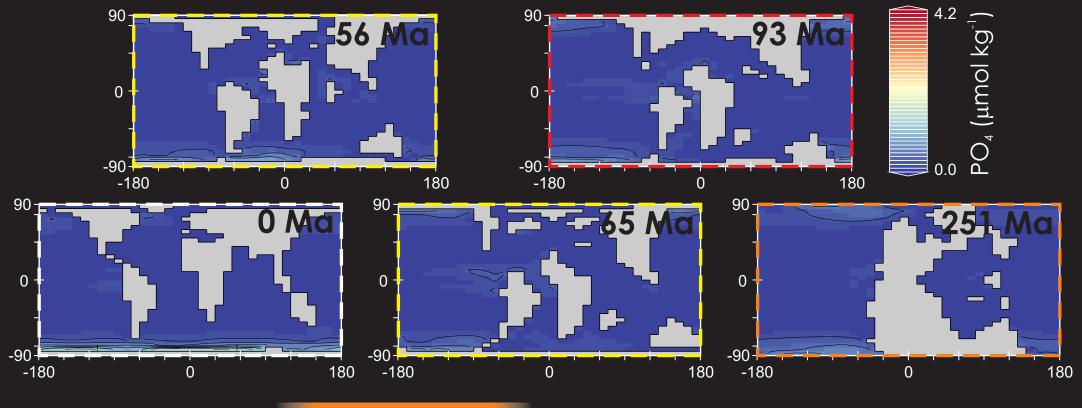
'estuarine circulation'

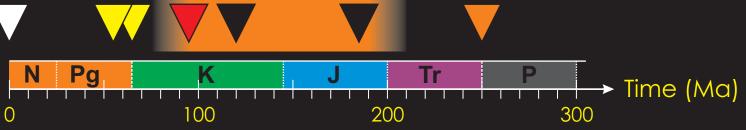




ocean surface [PO₄]

Biological export calculated by restoring nutrient concentrations ro zero. No 'spices' or 'flavourings' (e.g. iron) considered.

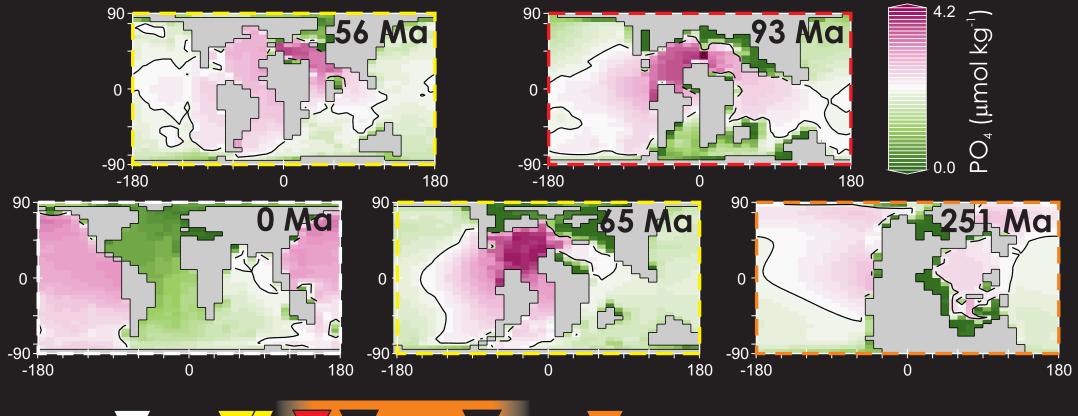


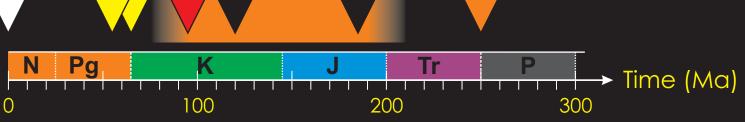


mean water column $[PO_4]$

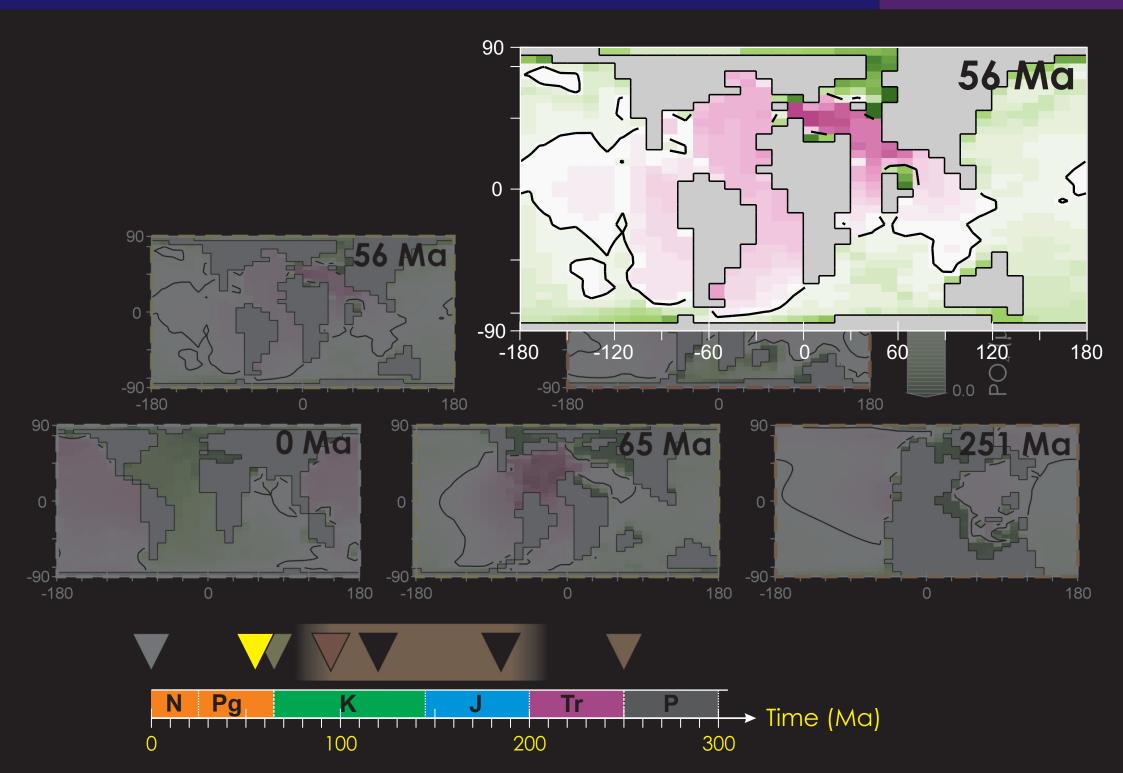
A measure of nutrient-trapping.

Contour represents concentrations equal to the prescribed whole ocean mean (2.1 μ mol kg⁻¹).

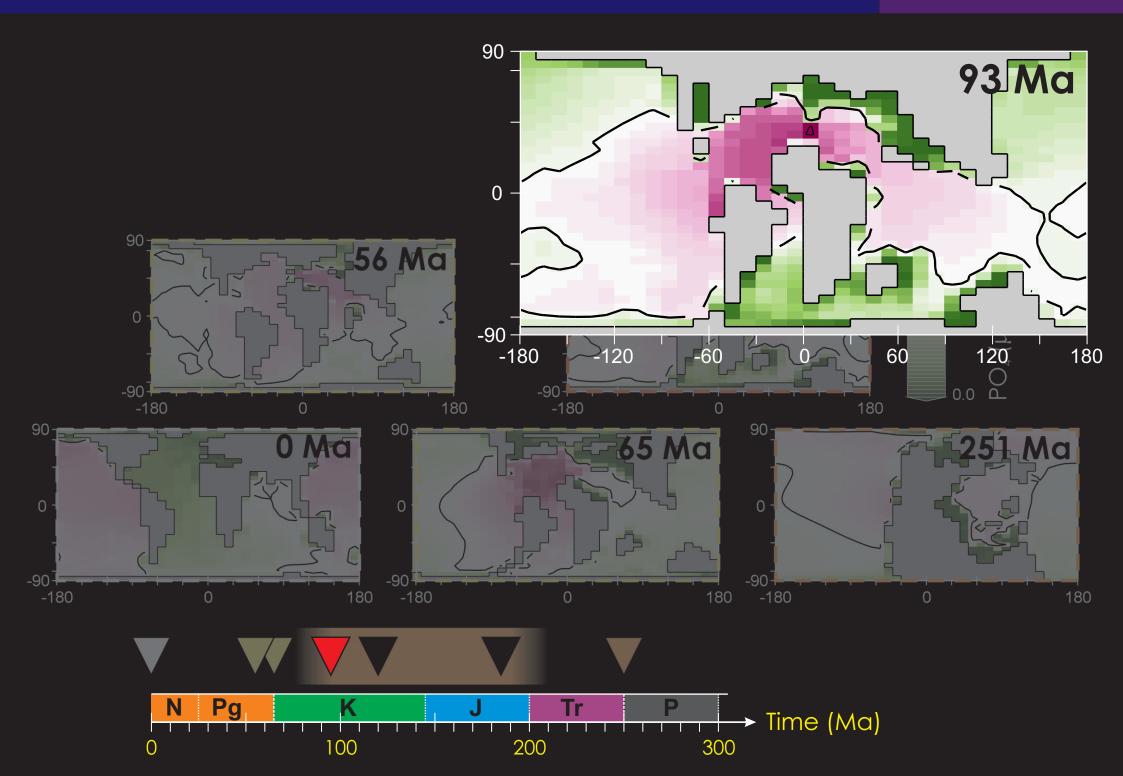




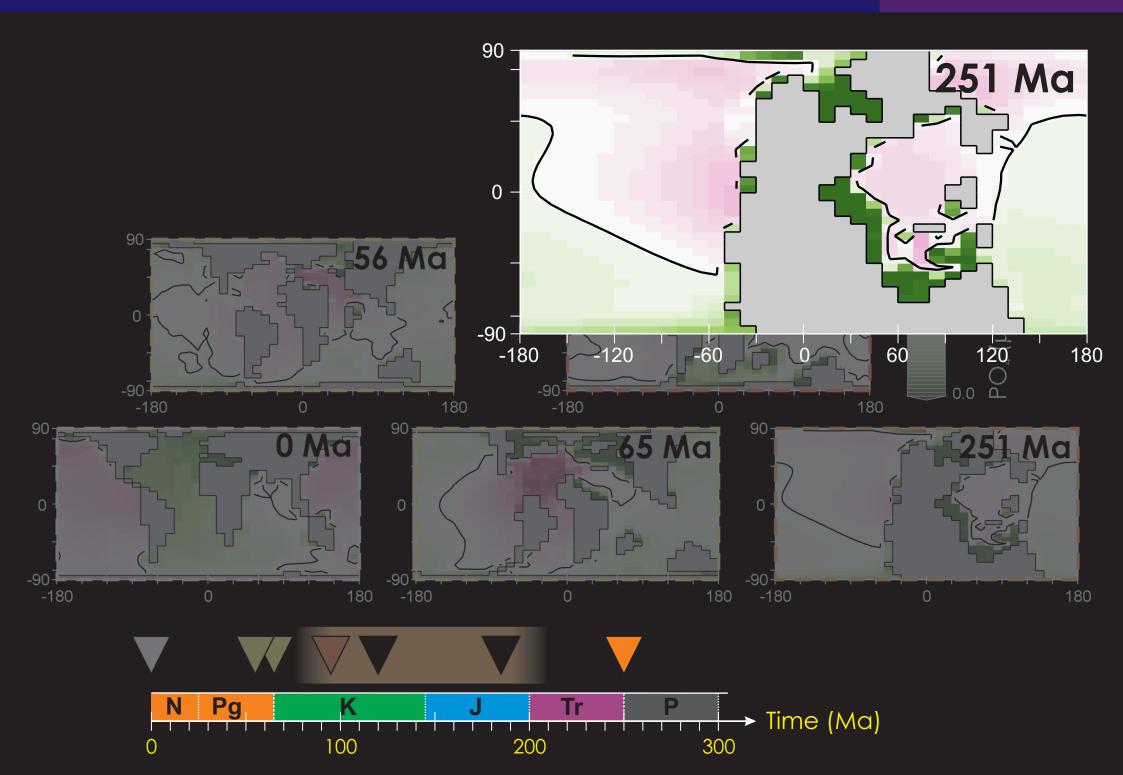
Carbon cycle and oxygen perturbations

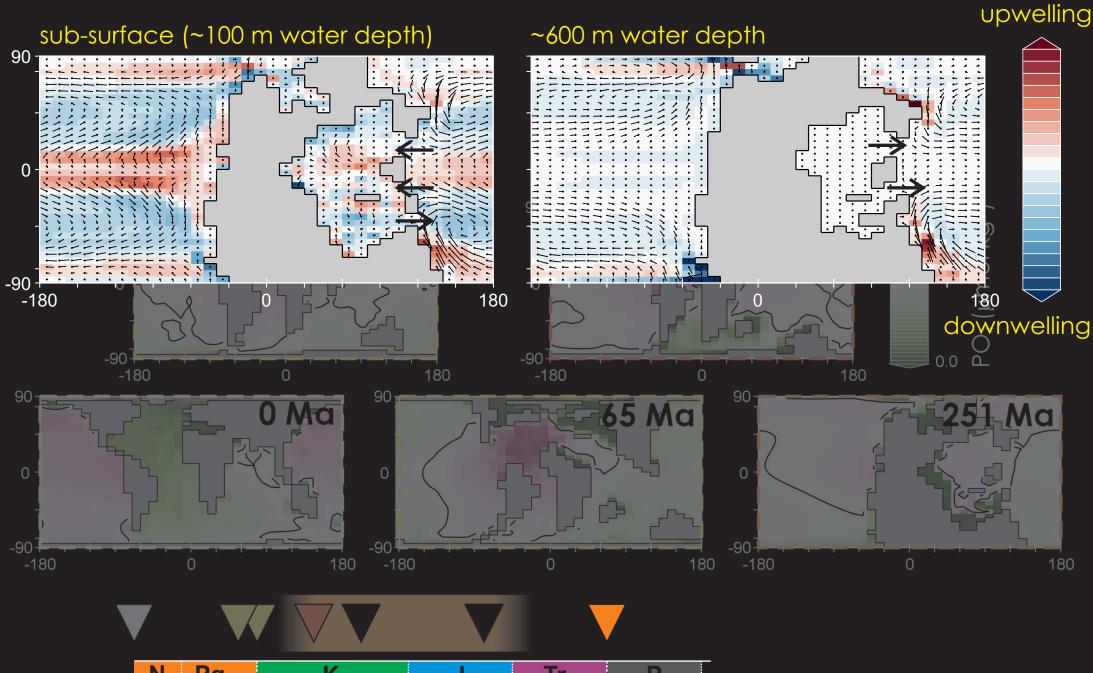


Carbon cycle and oxygen perturbations

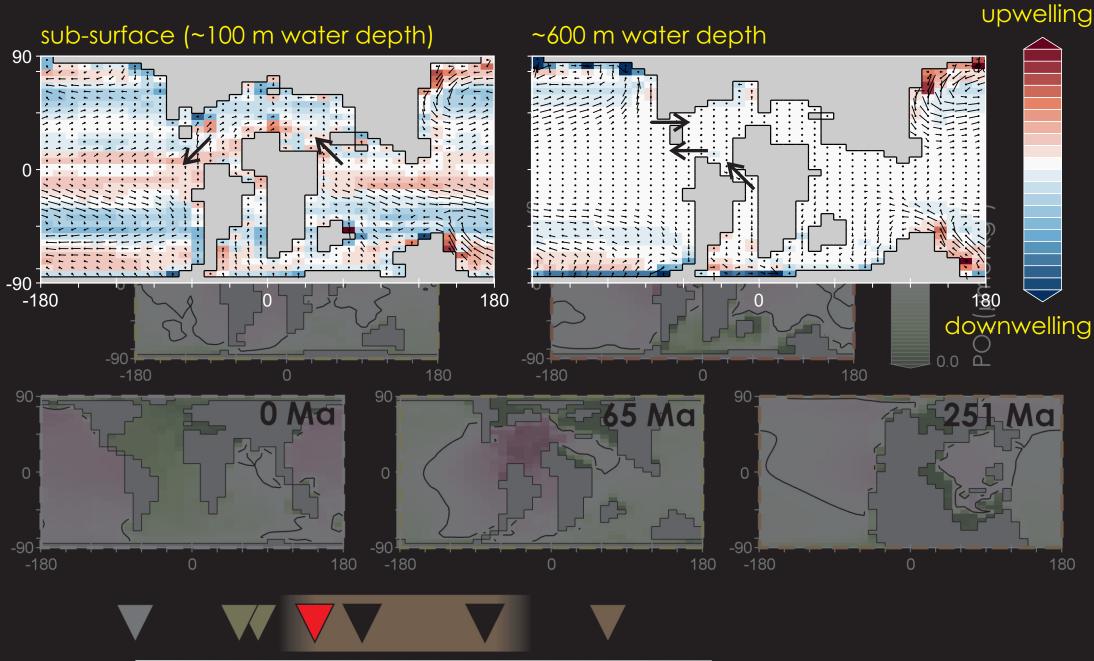


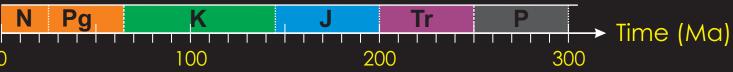
Carbon cycle and oxygen perturbations

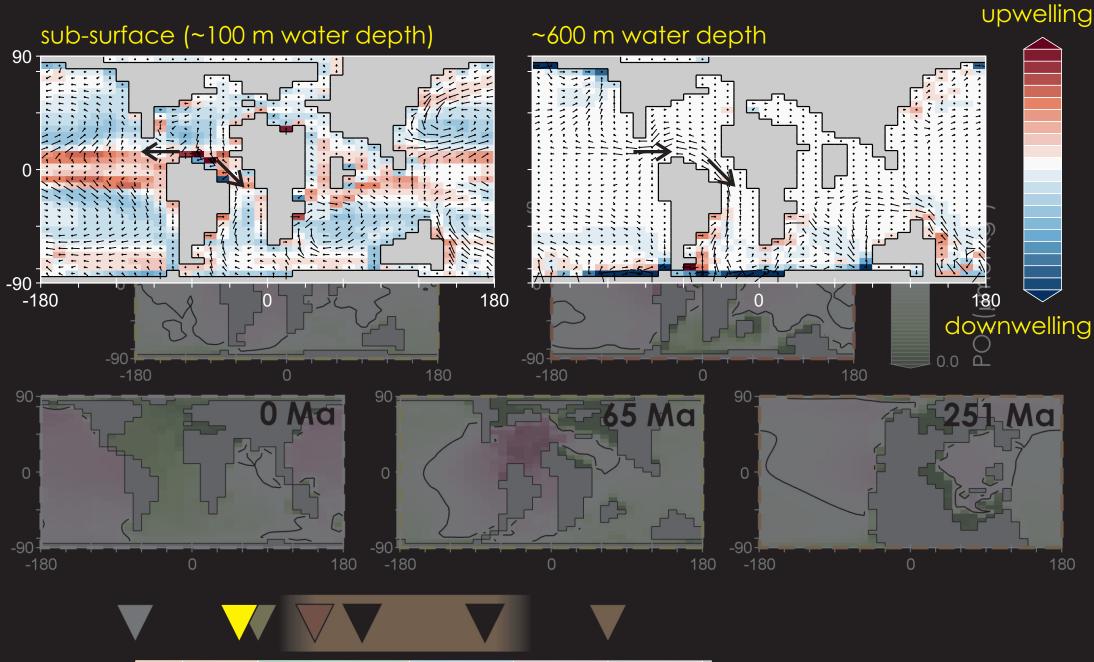














What is the 'recipe' for OAE (occurrence)?

* An 'estuarine-like' circulation, bringing in relatively nutrient-rich water at depth and removing relatively nutrient-depleted waters at the surface, will tend to lead to the trapping of nutrients and hence regional anoxia.

* Conversely, a circulation pattern in which water is net exported at depth will tend to act against the occurrence of regional OAE-like conditions.

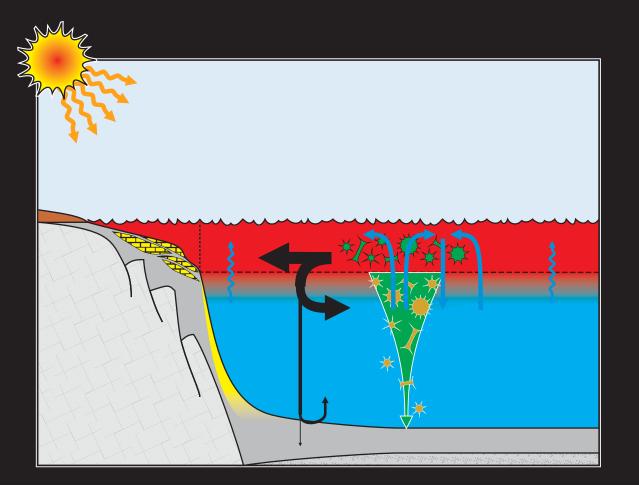
* However, the late Permian Tethys appears to have had an anti-estuarine circulation.

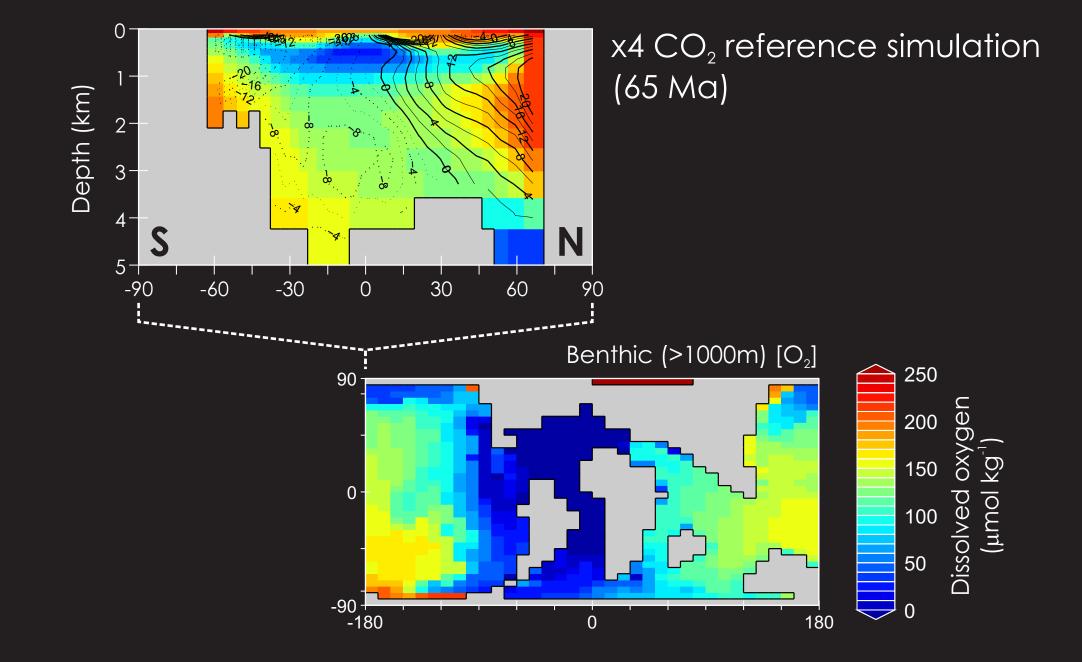


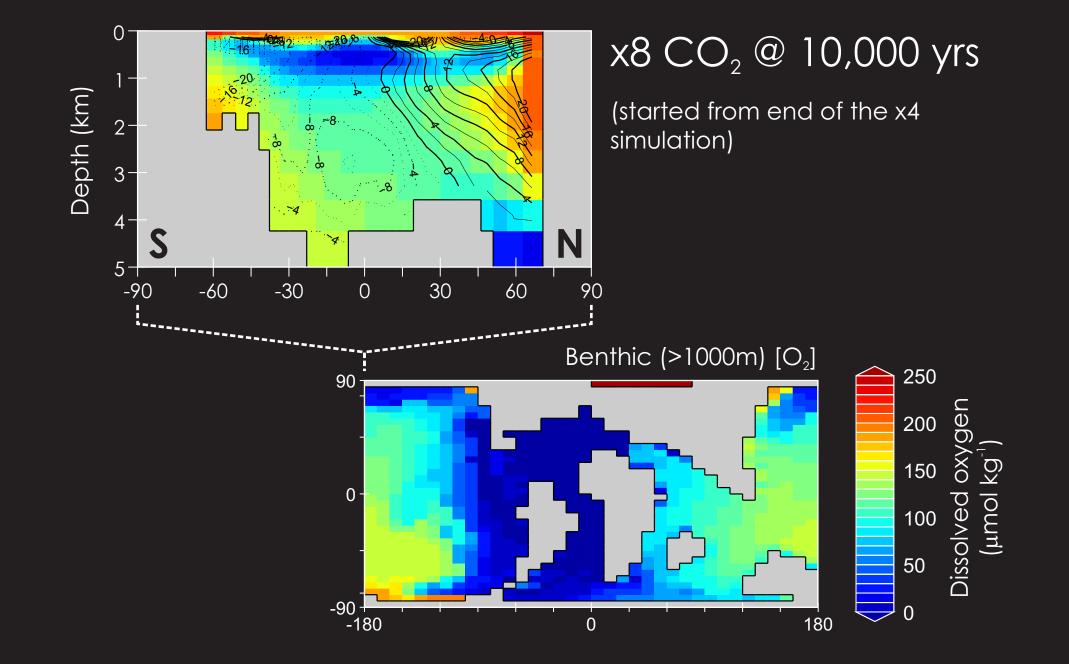


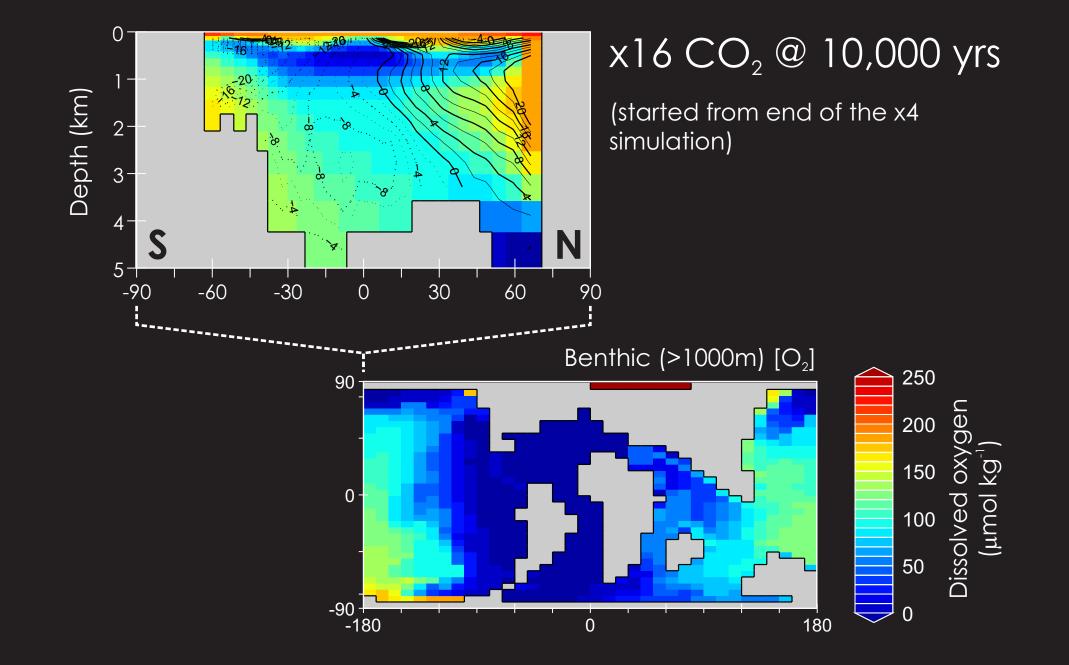
muffkie?

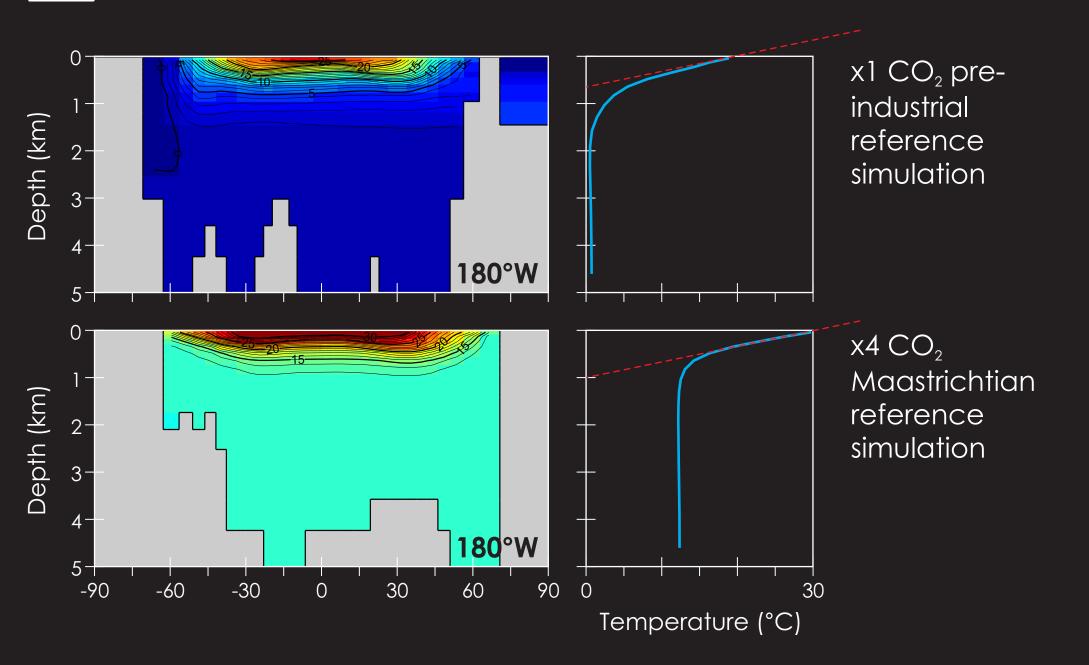
('stratified' || 'sluggish' || 'stagnant')







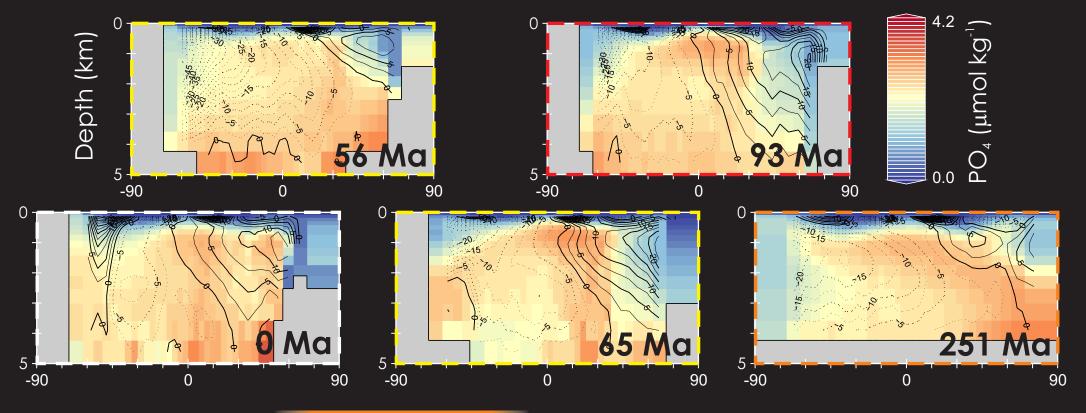


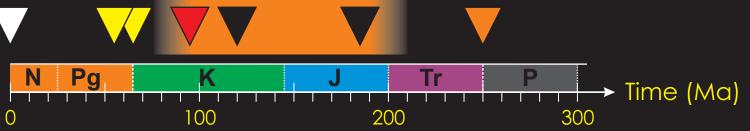


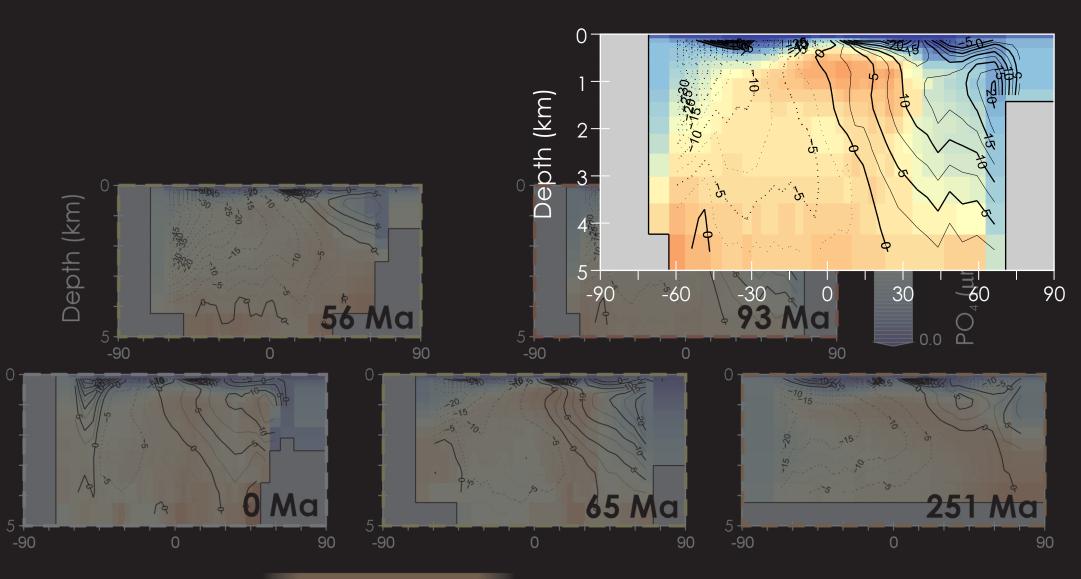
zonal mean latitude-depth [PO₄] distribution

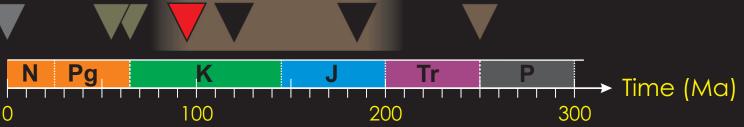
A measure of the partitioning of PO_4 and hence oxygen consumption, in the water column.

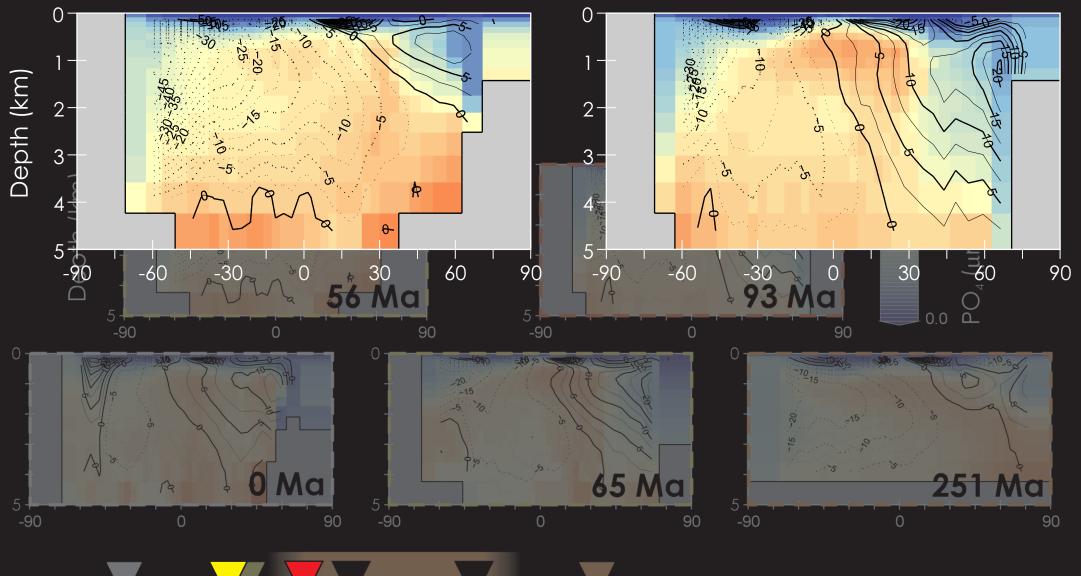
Contour overlay is the global mean overturning stream-function.

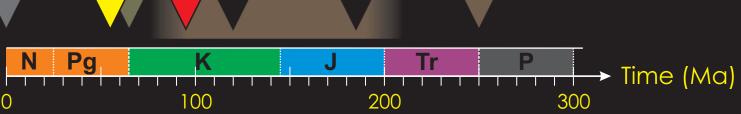


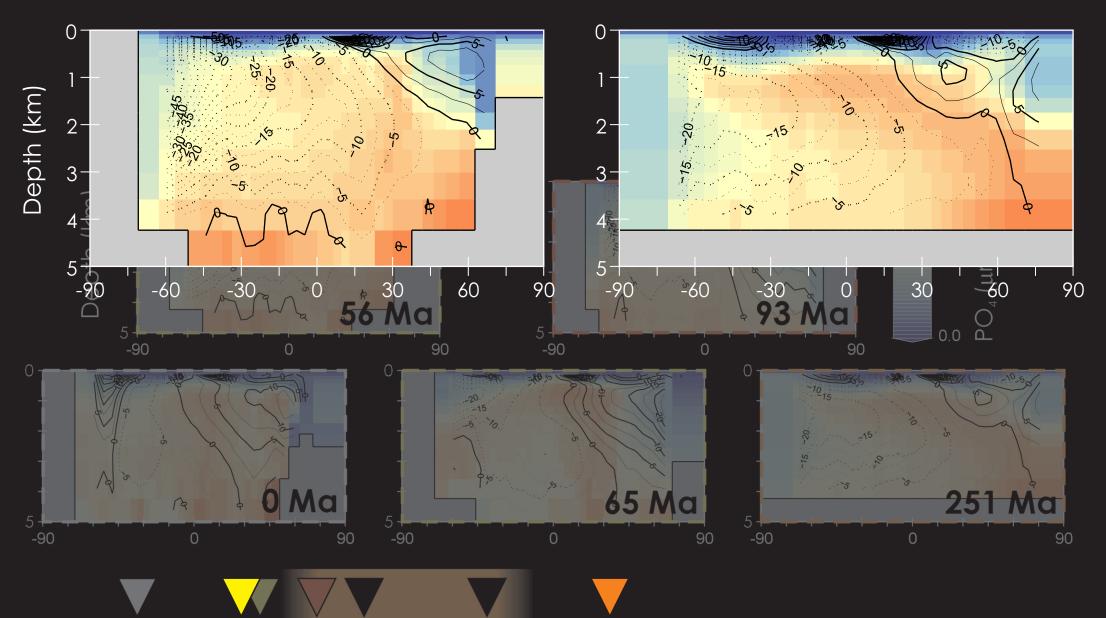


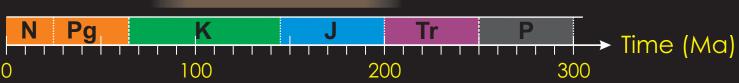












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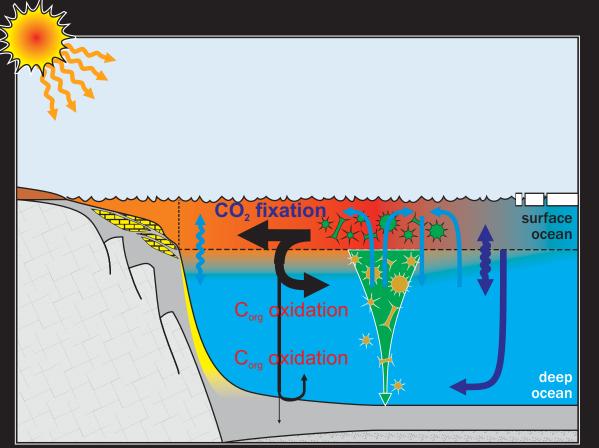
 \ast Strong (vigorous) ventilation of the deep ocean, rather than acting against the tendency towards OAE-like conditions, may actually be a pre-requesite as PO_4 is more rapidly returned to shallow and intermediate depths .

★ Conversely, weak ventilation and PO₄ trapping in the deep ocean may tend to act against the occurrence of OAE-like conditions.







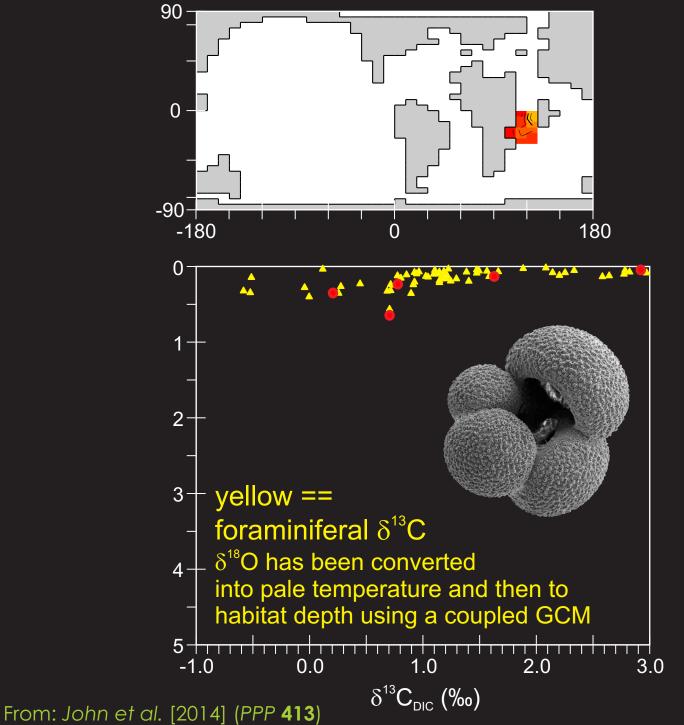


Bacterial metabolism and hence rate of dedregation of settling particulate organic matter (POM) should be temperate sensitive (e.g. Q_{10} ca. ~2).

It is reasonable to posit that a warmer ocean interior will have an on-average shallower depth of POM remineralization and hence more rapid nutrient recycling.

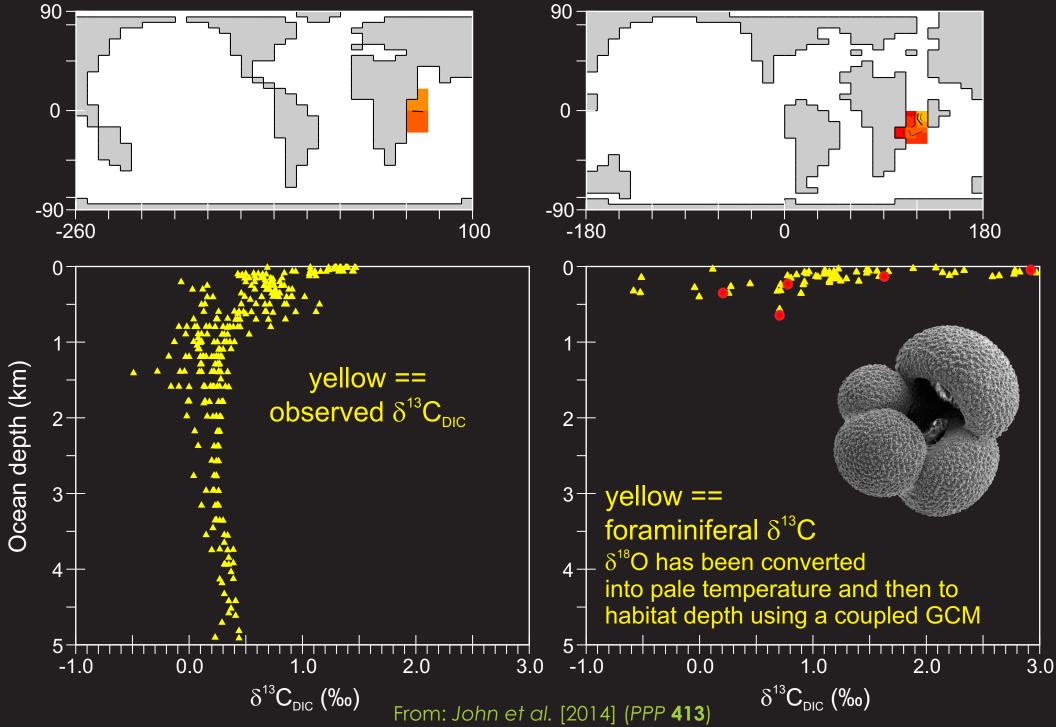
Evidence for this?

Planktic foraminiferal δ^{13} C from early Eocene Tanzania



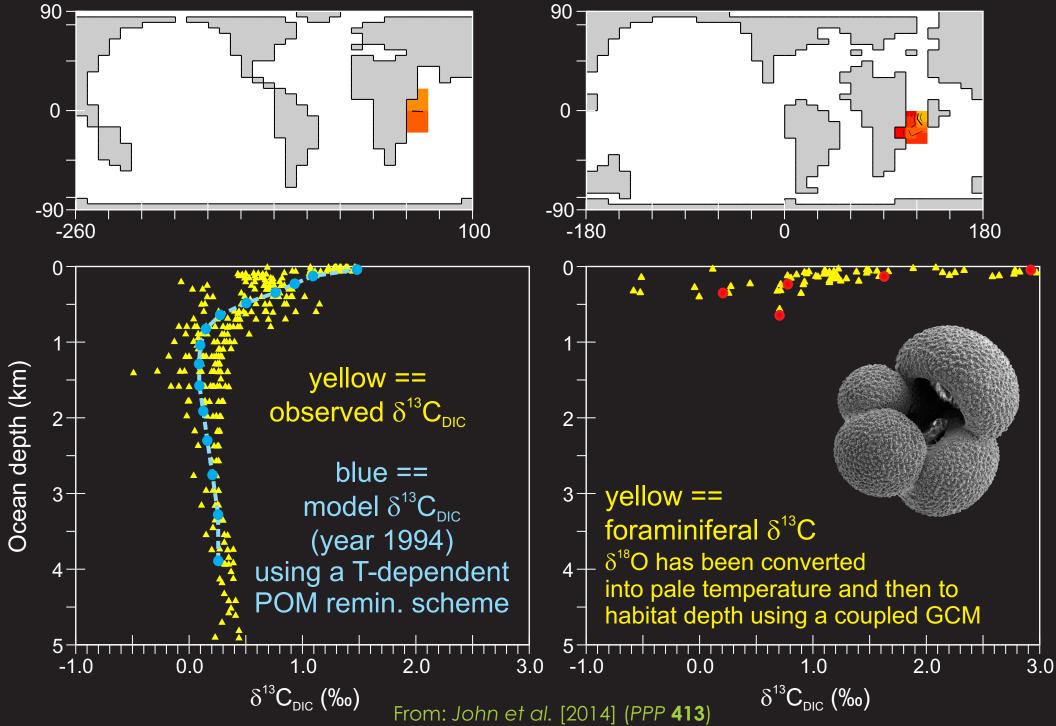
Open ocean $\delta^{13}C_{\text{DIC}}$ adjacent to modern Tanzania





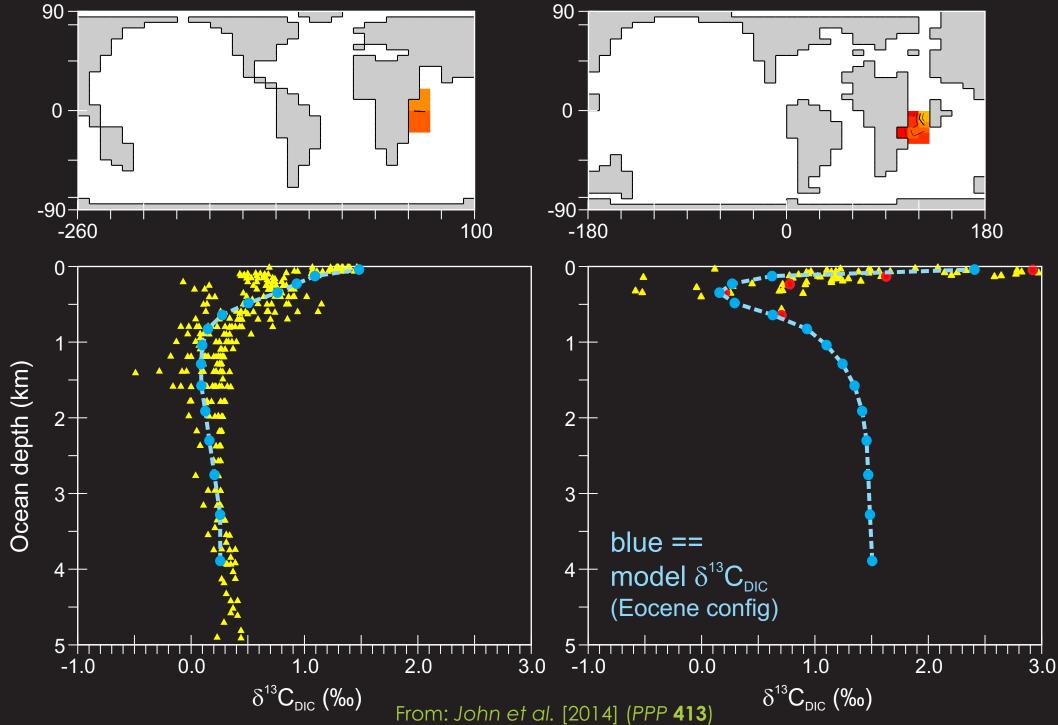
Open ocean $\delta^{^{13}}C_{_{\text{DIC}}}$ adjacent to modern Tanzania





Open ocean $\delta^{13}C_{DIC}$ adjacent to modern Tanzania

Planktic foraminiferal δ^{13} C from early Eocene Tanzania



Conclusions

* The mid Cretaceous had a tendency towards the occurrence of OAE-like conditions because of:

(i) nutrient trapping in the North Atlantic region,

(ii) vigorous ventilation of the deep ocean that helped recycle nutrients back towards the surface,

(iii) warm ocean temperatures and a shallow recycling depth-scale.

* The Paleocene-Eocene did not experience the occurrence of OAE-like conditions because of poorer deep ocean ventilation. However, otherwise, ocean circulation and temperature were relatively favourable.

* The end Permian appears anomalous in that the Tethys should have been nutrient poor. However, the Panthalassic ocean may have been well ventilated at depth with a pan-global concentration of nutrients at relatively shallow depths, then aided by progressive greenhouse warming. (i.e. not a situation of 'spreading' out from the Tethys but occurring truely globally) (analogous to the relationship of the PETM to the smaller Paleocene-Eocene hyperthermals?)

* Or ... differences in atmospheric pO_2 , etc etc ... ?



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