Colonization of the terrestrial environment 2016



Colonization of the terrestrial environment in silico

Andy Ridgwell

University of California – Riverside University of Bristol

+ 'T' Davies-Barnard, Paul Valdes (University of Bristol)+ Ryan Pavlick (JPL)

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RNC official's plagiarism defense for speech is 'My Little Pony'

Trump camp: Ax won't fall over speech

Ryan: 'Holy Moly! This explains everything'

Bergen: America the weak?























Terrestrial weathering can be (approximately equally) divided into carbonate (CaCO₃) and calcium-silicate ('CaSiO₃') weathering:

(1) $2CO_{2(aq)} + H_2O + CaSiO_3 \rightarrow Ca^{2+} + 2HCO_3^{-} + SiO_2$

(2) $CO_{2(aq)} + H_2O + CaCO_3 \rightarrow Ca^{2+} + 2HCO_3^{-}$

Ultimately, the (alkalinity: Ca²⁺) weathering products must be removed through carbonate precipitation and burial in marine sediments:

(3) $Ca^{2+} + 2HCO_3^{-} \rightarrow CO_{2(aq)} + H_2O + CaCO_3$

It can be seen that in (2) + (3), that the CO_2 removed (from the atmosphere) during weathering, is returned upon carbonate precipitation (and burial). In (1) + (3) (silicate weathering) CO_2 is permanently removed to the geological reservoir. This CO_2 must be balanced by mantle (/volcanic) out-gassing on the very long term.







Furthermore, the rate of silicate weathering should scale with climate. Hence the **silicate weathering feedback** is formed:

higher $pCO_2 \Rightarrow$ higher temperatures \Rightarrow higher rates of weathering \Rightarrow lower pCO_2



(1) Series of 1 Myr Earth system model experiments. CO₂ emissions from 1,000 to 20,000 PgC (GtC). Release interval: 1 yr.





(2) Fit each CO_2 decay curve with a series (4 optimal) of exponentials. Extract the fraction of CO_2 and time-scale associated with each.

(The resulting empirical model can be used in place of a mechanistic model for projecting the long-term fate of carbon release.)

 10^{6}

The 'long tail' of $CO_{2(excess)}$

Response of fraction of CO₂ removed vs. the characteristic time-scale, as a function of total emissions, ranging from 1,000 PgC (dark blue) to 20,000 PgC (yellow).





The 'long tail' of $CO_{2(excess)}$



Depletion of mixed layer carbonate buffer; ocean stratification and reduced surface mixing. Warming and reduced CO₂ solubility.







Ocean stratification and collapse of the AMOC (in this particular model). Threshold reached @ ~4000 PgC?









The 'long tail' of $CO_{2(excess)}$

III: τ ~ 1000-10000 years

Geologic CO₂ removal via carbonate rocks and marine sediments – occurring on an increasing protracted time-scale.











Sediments spanning the Palaeocene-Eocene boundary from ODP Leg 208 (Walvis Ridge) Picture courtesy of Dani Schmidt (University of Bristol)

20 21 22 23 24 25 26 27 28 29 30 31 32 33 34 35 36 37 38 39 40 41 42 43 44 45 46 47 48 49 50 51 52 53 54 55 56 57 58 5



4.6









The 'long tail' of $CO_{2(excess)}$













weathering «

soil properties rock mineralogy

 $\otimes f($







weathering f(soil properties
rock mineralogy)Solution</t

hydrology temperature pCO_2 $\land f$ (plant productivity plant (roots) type







weathering ∝

soil properties rock mineralogy



 $\bigotimes f\left(\begin{array}{c} hydrology \\ temperature \\ pCO_2 \end{array}\right) \bigotimes f\left(\begin{array}{c} plant productivity \\ plant (roots) type \end{array}\right)$







weathering «

soil properties rock mineralogy

\otimes **f**(













































In GCMs: Ocean circulation becomes an emergent rather than prescribed property of the system.

















3500

Jena Diversity-Dynamic Global Vegetation Model (JeDi-DGVM) [*Pavlick et al.*, 2013, *Biogeosciences*]

Period Era

Eon

Ng Pg

Cenozoic

Κ

100

M

J

200

Phanerozoic

300

Table C2Summary of the functional trait parameters.

Parameter	Description
t ₁	Growth response time to moisture conditions
t ₂	Growth response time to temperature conditions
t ₃	Critical temperature for growth
t ₄	Germination fraction
t ₅	Allocation to reproduction
t ₆	Allocation to aboveground growth
t ₇	Allocation to belowground growth
t ₈	Allocation to storage
t9	Relative allocation to aboveground structure
t ₁₀	Relative allocation to belowground structure
t ₁₁	Turnover time of structural pools
t ₁₂	Turnover time of leaf and fine root pools
t ₁₃	Senescence response time to productivity condition
t ₁₄	Relative senescence aboveground
t ₁₅	Plant nitrogen status

Neoproterozoic

1000

Cm

500

Time (Ma)

S

D

400

Paleozoic

Mesoproterozoic

2000

2500

3000





Time (Ma)



3000

3500

pre-industrial climate, *p*CO₂ == 280 ppm **modern** net primary productivity [gC m⁻² day⁻¹]



Time (Ma)



3500

pre-industrial climate, $pCO_2 == 280$ ppm modern functional richness [growth strategies]













weathering «

Na Pa

Cenozoic

Period

Era

Eon

0

Κ

J

Phanerozoic

300

soil properties rock mineralogy



Neoproterozoic

1000

Cm

500

Time (Ma)

Mesoproterozoic

3000

3500

 $\bigotimes_{p \in O_2} f\left(\begin{array}{c} \text{hydrology} \\ \text{temperature} \\ p \in O_2 \end{array}\right) \bigotimes_{p \in O_2} f\left(\begin{array}{c} \text{plant productivity} \\ \text{plant type (roots)} \end{array}\right)$

x8 CO₂ climate, pCO₂ == 1680 ppm **Ypresian** net primary productivity [gC m⁻² day⁻¹]



Paleozoic





Thanks to the funders ...

